Zhengshi Wang^{1,2}, Shan Ren², Ning Huang^{1,2}*

1 Key Laboratory of Mechanics on Disaster and Environment in Western China (Lanzhou University), The Ministry of Education of China, Lanzhou, China, 2 Department of Mechanics, School of Civil Engineering and Mechanics, Lanzhou University, Lanzhou, China

Abstract

PLOS ONE

Saltation is an important geological process and the primary source of atmospheric mineral dust aerosols. Unfortunately, no studies to date have been able to precisely reproduce the saltation process because of the simplified theoretical models used. For example, sand particles in most of the existing wind sand movement models are considered to be spherical, the effects of the sand shape on the structure of the wind sand flow are rarely studied, and the effect of mid-air collision is usually neglected. In fact, sand grains are rarely round in natural environments. In this paper, we first analyzed the drag coefficients, drag forces, and starting friction wind speeds of sand grains with different shapes in the saltation process, then established a sand saltation model that considers the coupling effect between wind and the sand grains, the effect of the mid-air collision of sand grains, and the effect of the sand grain shape. Based on this model, the saltation process and sand transport rate of non-spherical sand particles were simulated. The results show that the sand shape has a significant impact on the saltation process; for the same wind speed, the sand transport rates varied for different shapes of sand grains by as much as several-fold. Therefore, sand shape is one of the important factors affecting wind-sand movement.

Citation: Wang Z, Ren S, Huang N (2014) Saltation of Non-Spherical Sand Particles. PLoS ONE 9(8): e105208. doi:10.1371/journal.pone.0105208

Editor: Vanesa Magar, Centro de Investigacion Científica y Educacion Superior de Ensenada, Mexico

Received April 21, 2014; Accepted July 19, 2014; Published August 29, 2014

Copyright: © 2014 Wang et al. This is an open-access article distributed under the terms of the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

Data Availability: The authors confirm that all data underlying the findings are fully available without restriction. All relevant data are within the paper and its Supporting Information files.

Funding: This work is supported by the National Key Basic Research and Development Program ("973" Program) (2009CB421304), the Innovative Research Groups of the National Natural Science Foundation of China (11121202), the National Natural Science Foundation of China (11172118, 41371034, 91325203), and the National Key Technologies R & D Program of China (2013BAC07B01). The funders had no role in study design, data collection and analysis, decision to publish, or preparation of the manuscript.

Competing Interests: The authors have declared that no competing interests exist.

* Email: huangn@lzu.edu.cn

Introduction

Saltation plays a key role in various geological processes, including wind erosion, sediment transport, and the formation of sand dunes [1]. In-depth studies of sand saltation, which is the main manifestation of sand movement near the bed surface [2],[3], are considered to be essential to comprehensively understanding wind erosion and desertification processes and, furthermore, to preventing sand disasters[4]. Many scholars have carried out studies to simulate the saltation process [5–14]. However, significant discrepancies still exist between the simulated results and measurements [1]. Current models of saltation make many simplifications because of the complexity of sand saltation. For example, most models assume that sand particles are spherical, and the effect of mid-air collision is ignored.

In fact, spherical and sub-spherical sand grains generally comprise no more than 30% of the total sand [15]. Many scholars have studied sand shape in the desert [16–20]. For example, Sagga [16] and Juan [17] studied the roundness of sand grains from dunes in Saudi Arabia and Mexico, respectively. Sagga found that the sands of the inter-dunes are more rounded than are those of the adjacent dunes and that sand particles tend to be less rounded with increasing dune height. Juan proposed that aeolian and marine processes might generate different grain sizes with different compositions. Cheng et al. analyzed aeolian sandy physical characteristics and measured the roundness, average particle size and size distribution of desert sand north of Beijing. They indicated that these parameters have an important influence on the interactional force of aeolian sand [18].

Researchers have investigated the mechanical properties of nonspherical particles for over half century. Barton studied the slip correction factors for non-spherical bodies in continuum flow in 1973 [21]. Then, Gavze and Shapiro investigated particles in a shear flow near a solid wall considering the impact of nonspherical grains [22]. In recent years, more studies have been conducted in this field. For example, Yin et al. developed a model of motion of cylindrical particles in a non-uniform flow [23]. Xu et al. discussed drag and lift forces of rotational non-spherical particles [24]. Zastawny et al. considered the drag and lift force and torque coefficients for non-spherical grains in flows [25].

Based on the above studies of the motion of particles in a shear flow, our study first proposes a relationship between sand shape and the threshold friction velocity of wind. Then, a sand saltation model, which takes into consideration the coupling effect between wind and sand grains, the effect of the mid-air collision of sand grains, and the effect of the shape of sand grains, is established, and the saltation process of non-spherical particles is simulated. Finally, the effects of different shapes of non-spherical sand grains on wind-sand movement are analyzed in detail in this paper.

Methods

1. Analysis of sand shapes

Blott and Pye [26] applied Beckman-Coulter to analyze natural sand particles, including desert dune and coastal dune sand

particles, and found that the average roundness Ψ of natural sand is in the range of 0.82–0.86. The average roundness Ψ is defined as the ratio of the surface area A_v of a sphere to the surface area A_p of a sand particle with an equivalent volume, that is

$$\Psi = A_v / A_p = \pi d_v^2 / A_p \tag{1}$$

where d_{v} is the equivalent diameter of the sphere with the same volume as that of the sand particle.

Let $S_i(a,b,c)$ represent non-spherical particles of different shapes, where a, b and c are their size parameters, respectively. In this study, we examined the effects of four types of nonspherical particles, including ellipsoid, cube, cylinder and frustum, on wind-sand movement.

- 1) Ellipsoid-shaped sand particles: particles with a shape approximate to an ellipsoid are simplified to an ellipsoid, with a, b and c indicating the lengths of the long axis, mid axis, and short axis, respectively. We selected two types of spheroids, $S_1(l,l,2l/5)$ and $S_2(l,l/3,l/3)$, respectively.
- 2) Cube-shaped sand particles: particles shaped like cubes are reduced to a regular cube $S_2(l,l,l)$.
- 3) Cylinder-shaped sand particles: particles shaped like cylinders are simplified to a regular cylinder, with a and b indicating the diameter of the bottom and c indicating the height. Two cuboids were selected, $S_4(3l/5,3l/5,l)$ and $S_5(l,l,3l/5)$; that is, their bottom diameters are 3l/5 and l, respectively, while their heights are l and 3l/5, respectively.
- 4) Frustum-shaped sand particles: irregular frustum-shaped sand particles are reduced to a regular frustum, wherein a, b and c indicate the upper end diameter, the lower end diameter, and the height, respectively. The regular frustum $S_6(l/2, l, l)$ was selected with a, b and c corresponding to l/2, l, and l, respectively.

The symbol l in the above 1) to 4) is a constant. To consider the effects of different shapes, three shape parameters are introduced [27] as follows:

- 1. The aspect ratio, $E = d_{\parallel}/d_{\perp}$, where d_{\parallel} and d_{\perp} are the maximum diameters or sizes of the horizontal and vertical directions of the non-spherical particles (Table 1);
- 2. The ratio of the surface areas, $A_{surf}^* \equiv A_{surf} / (\pi d_v^2)$; A_{surf} is the surface area of non-spherical grains. By comparison with formula (1), $A_{surf}^* = 1/\Psi$ (Table 1);
- The ratio of the projected areas, A^{*}_{proj} ≡ 4A_{proj}/πd²_v; A_{proj} is the projected area of the non-spherical grains (Table 1).

2. Drag coefficient of non-spherical grains

The flow field can be divided into three parts according to the Reynolds number [28]: a creeping field of flow ($Re \ll 1$), an intermediate field of flow and a Newton field of flow (ca. 2000 < Re < 300000).

Under the actions of the three fields of flow, the drag coefficient of spherical particles often is [2],

$$C_D = 24/\text{Re} + 6/(1 + (\text{Re})^{1/2}) + 0.4$$
 (2)

where the Reynolds number $\operatorname{Re} = D_p V_r / v$.

To consider the impact of particle shape, the modified coefficient of shape is defined as [27],

$$C_{shape} = C_{D,shape} / C_{D,sphere} \tag{3}$$

where $C_{D,shape}$ is the drag coefficient of a non-spherical sand particle, and $C_{D,sphere}$ is the drag coefficient of a spherical sand particle.

Under different fields of flow, the modified coefficient of particle shape will be different. The following is an analysis of the shapemodified coefficients of non-spherical particles subject to three types of fields of flow:

The particles in the creeping flow field [27], Re <<1:

$$C_{shape}^{1} = \sqrt{A_{proj}^{*}} / 3 + 2\sqrt{A_{surf}^{*}} / 3 \tag{4}$$

The particles in the Newton flow field [27], $2000 < R_e < 300000$:

$$C_{shape}^{2} = 1 + 1.5(A_{surf}^{*} - 1)^{1/2} + 6.7(A_{surf}^{*} - 1), E \le 1$$
 (5)

$$C_{shape}^{2} = 1 + 0.7(A_{surf}^{*} - 1)^{1/2} + 2.4(A_{surf}^{*} - 1), E > 1$$
 (6)

The particles in the intermediate flow field [27]: $C_D^* = C_D / C_{shape}, R_{ep}^* = C_{shape}^2 R_{ep} / C_{shape}^1$, where C_D^* and R_{ep}^* are the substitute of the intermediate flow field. Some scholars have obtained the following empirical formula [27]:

	S ₁	S ₂	S ₃	S ₄	S ₅	S ₆
Ψ	0.853	0.849	0.806	0.847	0.850	0.842
A^*_{surf}	1.1729	1.1779	1.2407	1.1806	1.1765	1.1876
A^*_{proj}	1.8420	0.4808	0.8271	1.0728	0.5429	1.0931
Ε	2/5	3	1	3/5	5/3	1

Table 1. Parameters of sand particles with different shapes.

where S_1 and S_2 are two different ellipsoid-shaped sand particles, S_3 represents cube-shaped sand particles, S_4 and S_5 represent two types of cylinder-shaped sand particles, and S_6 represents frustum-shaped sand particles.

doi:10.1371/journal.pone.0105208.t001



Figure 1. Schematic of analysis of the forces acting on a sand particle. Particles B and C are the supporting sand particles on the bed surface. doi:10.1371/journal.pone.0105208.g001

For sand particles with a circular cross-section:

$$C_D^* = 24 \left[1 + 0.15 (R_{ep}^*)^{0.687} \right] / R_{ep}^* + 0.42 / \left[1 + 42500 / (R_{ep}^*)^{1.16} \right] (7)$$

For sand particles with a non-circular cross-section:

$$C_D^* = 24[1 + 0.035(R_{ep}^*)^{0.74}] / R_{ep}^* + 0.42 / [1 + 33 / (R_{ep}^*)^{0.5}]$$
(8)

3. Starting friction speed of non-spherical particles

Before starting to move, sand particles on the bed surface are mainly subject to the gravity W, the drag force F_D , the force of supporting the bed F_N , and the cohesive force F_{μ} among the sand grains, among which the gravity and the drag force are given as follows [11]:

$$W = \pi d_v^3 (\rho_p - \rho)g/6 \tag{9}$$

$$F_D = C_D \rho A V_r^2 / 2 \tag{10}$$

where A is the maximum area of the particle cross-section. A will adopt different values for different sand shapes; V_r is the relative speed between the sand grains and the gas flow; μ is the aerodynamic coefficient of viscosity; ρ_P is the density of the sand particles; ρ is the density of air; and g is the acceleration due to gravity.

Here, we focus on dry sand particles. Thus, for large particles, the cohesive force among the particles is very small and can be ignored [2]; that is, $F_{\mu} = 0.0$.

Fig. 1 shows a schematic of the force analysis among three nonspherical particles on the bed surface, in which particles B and C are the supporting sand particles of the bed surface. Relative to Point O_1 , the sand particle A is subject to the moment of force. According to the moment balance principle, as the sand particles satisfy formula (11), they will leave the bed surface:

$$F_D L_D \ge W L_W \tag{11}$$

where $L_W = l \cos \theta$ and $L_D = l \sin \theta$.

From inequality (10), the critical condition for the sand particles to start saltation is:

$$F_D \sin \theta = W \cos \theta \tag{12}$$

Substituting formulas (9) and (10) into (12), we obtain the starting speed of flow as follows:

$$u_c = \sqrt{\pi d_v^{3}(\rho_p - \rho)g/(3C_D\rho A\tan\theta)}$$
(13)

For a given sand particle shape, u_c is only related to θ , which is a random variable, depending on the acting point of the supporting particle against the saltation particle.

4. Wind field equation

Before sand particles start to move near the ground's surface, the motion of air flow can be described by the Navier-Stokes equation:

$$\rho \,\partial \vec{u} / \partial t + \rho \vec{u} \cdot \nabla \vec{u} = -\nabla p + \nabla \cdot \vec{\tau} + \rho \vec{g} \tag{14}$$

where \vec{u} is the velocity of wind, t is time, \vec{g} is the gravitational acceleration, $\vec{\tau}$ is the Reynolds' tangential stress, and p is the pressure.

When sand particles leave the bed and move into the saltation layer, the blocking effect of the particles against the wind field will change their distribution, and this blocking effect can be equivalent to adding a correction term, the volume force, to the right of the Navier-Stokes equations:

$$\rho \,\partial \vec{u} / \partial t + \rho \vec{u} \cdot \nabla \vec{u} = -\nabla p + \nabla \cdot \vec{\tau} + \rho \vec{g} + \vec{F}_x \tag{15}$$

Letting the direction of air flow movement be in the x-axis and the direction perpendicular to the bottom be the y-axis, $\vec{u} \cdot \nabla \vec{u} = 0$, $\partial/\partial y \gg \partial/\partial x$, $\partial/\partial y \gg \partial/\partial z$. Then, we obtain the motion differential equation of the wind field component in the x direction of in the saltation layer of the wind sand flow.

$$\rho \partial u / \partial t = \partial \tau / \partial y + F_x(y) \tag{16}$$

If a two dimensional steady state is fully developed in the flow field, then the left term in Eq. (16) is equal to zero, that is, $\frac{\partial}{\partial t} = 0$; thus, Eq. (16) becomes

$$0 = d\tau/dy + F_x(y) \tag{17}$$

The blocking effect of saltating particles against the wind field changes with height, and thus the stress acting in the field of the wind also changes with height. Integrating Eq. (17) with respect to height results in Eq. (18):

$$\tau(y) + \int_{y}^{y_{max}} F_x(y) dy = \tau_b$$
(18)

where y_{max} is the max height to which the particle can move upward, τ_b is the shear stress of the air flow from the blocking effect from the sand particle above y_{max} , τ_y is the air flow shear stress at height y, and $\int_{y}^{y_{\text{max}}} F_x(y) dy$ is the particle-loaded tangential stress [7], [11].

Based on the Prandtl mixing length, the Reynolds tangential stress becomes

$$\tau = \mu_t \, du/dy \tag{19}$$

In the saltation layer of the wind-sand flow, the use of the equivalent friction speed can obtain [7], [11]:

$$u_{*eff} = \sqrt{\tau/\rho} \tag{20}$$

$$\mu_t = \rho k u_{*eff} \tag{21}$$

where u_{*eff} is the equivalent friction speed in the saltation layer and varies with y, and μ_t is the viscosity coefficient of the turbulent flow. Bringing Eqs. (19)–(21) into Eq. (18), we obtain the twodimensional steady state flow field equation fully developed for wind sand flow [13]:

$$du/dy = \left[\tau_b - \int_y^{y_{\text{max}}} F_x(y) dy \middle/ \rho\right]^{\frac{1}{2}} \middle/ ky$$
(22)

In accordance with the logarithmic distribution profile that the wind field obeys

$$u = 5.75u_* \lg(y/y_0) \tag{23}$$

where $y_0 = d_v/30$ is the roughness of the bed surface.

The critical starting height is $y_c = d/2$. Bringing $u = u_c$ and $y = y_c$ into formula (23), we can obtain the critical starting friction speed u_{*t} of non-spherical sand particles.

5. Saltation equation

The analysis of sand grains subject to forces indicates that in the process of their movement, they experience only the gravity W and the drag force F_D , as shown in Formulas (9) and (10), respectively. The basic equations of sand saltation movement are:

$$m_p d^2 x / dt^2 = f_x = -\rho A C_D (\dot{x} - u) \sqrt{(\dot{x} - u)^2 + \dot{y}^2} / 2$$
 (24a)

$$m_p d^2 y / dt^2 = f_y - m_p g = -\rho A C_D \dot{y} \sqrt{(\dot{x} - u)^2 + \dot{y}^2} / 2 - m_p g (24b)$$

in which m_p is the mass of the sand particles, g is the gravitational acceleration, u is the wind speed in the x-direction, and \dot{x} and \dot{y} are the velocity components of the particles in the x-direction and y-direction, respectively.

The starting initial velocity distribution of saltating sand particles was obtained by Dong et al. in their experiments [29]:

$$f(v_p) = f(u_0, v_0) = f(u_0)f(v_0) = (f + (A / \omega \sqrt{\pi/2})e^{-2(u_0 - u_c / \omega)^2})(A_1 e^{-(v_0 - x_0)/t_1} + A_2 e^{-(v_0 - x_0)/t_2} + f_1)$$
(25)

where f, A, ω , u_c , f_1 , x_0 , A_1 , t_1 , A_2 , and t_2 are recession parameters.

6. Resistance of saltating sand groups to the wind field

The interaction between the wind flow and the saltation layer plays an important role in the saltation process [3]. After considering the sand shape, the resistance from sand particles experienced by the wind per volume can be described [13]:

$$F_{x}(y) = s(\pi d_{v}^{3} \rho_{p} / 6) \int f(v_{p})(a_{x\uparrow}(y) / |\dot{y}_{\uparrow}| + a_{x\downarrow}(y) / |\dot{y}_{\downarrow}|) dv_{p} \quad (26)$$

where $a_{x\uparrow}(y)$ and $a_{x\downarrow}(y)$ denote the accelerations of the particles moving upwards and downwards at height y, respectively, and \dot{y}_{\uparrow} and \dot{y}_{\downarrow} are the average speeds of the particles moving upwards and downwards at height y, respectively.

7. Calculations of the sand transport rate (STR) and the STR per width

Considering the mechanism of sand collision, the STRs of irregular particles can be divided into two STRs representing



Figure 2. The drag coefficient of sand particles as a function of the Reynolds number. The relationship of the drag coefficient of sand grains of different shape to the Reynolds number is shown. Here, the black solid line indicates spherical sand grains, S_1 and S_2 are two different ellipsoid-shaped sand particles, S_3 represents cube-shaped sand particles, S_4 and S_5 represent two types of cylinder-shaped sand particles, and S_6 represents frustum-shaped sand particles. doi:10.1371/journal.pone.0105208.g002

moving particles with collision and those without collision in air. The STR at height *y* considering the aerial collision mechanism of sand particles and the co-initial speed distribution of starting particles is described as [13], [30]:

$$Q_{c}(y) = \int_{v_{p\min}}^{v_{p\max}} sm_{p}f(v_{p})[(1-p_{\uparrow}(v_{p},y))\dot{x}_{\uparrow}(v_{p})/\dot{y}_{\uparrow}(v_{p})-(1-p_{\downarrow}(v_{p},y))\dot{x}_{\downarrow}(v_{p})/\dot{y}_{\downarrow}(v_{p})]dv_{p}$$

$$(27)$$

$$+ \int_{v_{p\min}}^{v_{p\max}} sm_{p}f(v_{p})[p_{\uparrow}(v_{p},y)(\dot{x}_{\uparrow}(v_{p})/\dot{y}_{\uparrow}(v_{p}))_{p}-p_{\downarrow}(v_{p},y)(\dot{x}_{\downarrow}(v_{p})/\dot{y}_{\downarrow}(v_{p}))_{p}]dv_{p}$$

Integrating Eq. (26) with respect to height, we obtain the STR per width of irregular particles [13], [30]:

$$Q_{mc} = \int_{0}^{y} \int_{v_{p\min}}^{v_{p\max}} sm_{p}f(v_{p})[(1-p_{\uparrow}(v_{p},y))\dot{x}_{\uparrow}(v_{p})/\dot{y}_{\uparrow}(v_{p})-(1-p_{\downarrow}(v_{p},y))\dot{x}_{\downarrow}(v_{p})/\dot{y}_{\downarrow}(v_{p})]dv_{p}$$

$$+ \int_{0}^{y} \int_{v_{p\min}}^{v_{p\max}} sm_{p}f(v_{p})[p_{\uparrow}(v_{p},y)(\dot{x}_{\uparrow}(v_{p})/\dot{y}_{\uparrow}(v_{p}))_{p} - p_{\downarrow}(v_{p},y)(\dot{x}_{\downarrow}(v_{p})/\dot{y}_{\downarrow}(v_{p}))_{p}]dv_{p}$$

$$(28)$$

where $p_{\uparrow}(v_p, y)$ is the probability of collision of rising particles at height y with an initial lifting-off velocity of v_p , $(\dot{x}_{\uparrow}(v_p)/\dot{y}_{\uparrow}(v_p))_p$ is the ratio of the horizontal velocity to the vertical velocity after the aerial collision of the rising particles at height y with an initial lifting-off velocity of v_p considering the aerial collision mechanism, $p_{\downarrow}(v_p, y)$ is the probability of collision of falling particles at heightywith a co-initial lifting-off velocity of v_p , and $(\dot{x}_{\downarrow}(v_p)/\dot{y}_{\downarrow}(v_p))_p$ is the ratio of the horizontal velocity to the vertical velocity after the aerial collision of falling particles at height y with a co-initial lifting off velocity v_p .

8. Criterion of wind sand flow up to dynamic equilibrium

The critical impact starting velocity of wind is approximately 80% of the starting wind velocity of a critical fluid [2]. From the starting wind speed of a critical fluid u_c obtained from Eq. (13), the critical impact starting shear stress considering the sand shape is

$$\tau_{ti} = 0.8\tau_{tf} = 0.8\rho u_c^2 \tag{29}$$

where τ_{ti} and τ_{tf} are the critical impact starting shear stress of the wind sand flow and the starting shear stress of a critical fluid [13].

9. Constants in the model

For the convenience of comparison, the equivalent diameters of various shapes of sand particles are assumed to be the same. We select two groups of sand particles in this article, $d_v = 0.25$ mm and 0.35 mm; the sand density is $\rho_p = 2650 \text{ kg/m}^3$, the air density is $\rho = 1.22 \text{ kg/m}^3$, and the dynamic viscosity coefficient of air is $v = 1.5 \times 10^{-5}$; and we assume that the initial wind speed profile obeys a logarithmic distribution.



Figure 3. The critical starting friction wind velocity for sand grains of different shape. (a). Comparison of the critical starting friction wind velocity of non-spherical particles with that of spherical particles (equivalent diameter $d_v = 0.25mm$). (b). Comparison of the critical starting friction wind velocity of non-spherical particles with that of spherical particles (equivalent diameter $d_v = 0.25mm$). (b). Comparison of the critical starting friction wind velocity of non-spherical particles with that of spherical particles (equivalent diameter $d_v = 0.35mm$). The black solid line indicates the spherical sand grains, S_1 and S_2 represent two different ellipsoid-shaped sand particles, S_3 represents cube-shaped sand particles, S_4 and S_5 are two types of cylinder-shaped sand particles, and S_6 represents frustum-shaped sand particles. doi:10.1371/journal.pone.0105208.q003



Figure 4. The drag force of different shaped particles with height. All the sand particles of different shapes take off with the same horizontal initial velocity of V_1 and vertical initial velocity of V_2 . The black solid line indicates spherical sand grains, S_1 and S_2 are two different ellipsoid-shaped sand particles, S_3 represents cube-shaped sand particles, S_4 and S_5 are two types of cylinder-shaped sand particles, and S_6 represents frustum-shaped sand particles.

doi:10.1371/journal.pone.0105208.g004

10. Computational steps

- 1) Substitute the parameters of non-spherical sand grains A_{surf}^* , A_{proj}^* and E into Eqs. (4)–(8) to calculate the corresponding shape parameters and calculate the resistance coefficients of different shape of particles C_D , then bring the coefficients into Eq. (10) to obtain the drag forces of different shapes of sand particles F_D .
- 2) Find the values of parameters τ_{ti} and u_{in} and calculate the initial friction wind speed u_* , then obtain the initial wind speed distribution and denote them $u^1(y)$ or $u^i(y)$.
- 3) Select the testing value of the sand liftoff rate from the sand bed *s*, calculate the co-initial speed distribution function of the saltation particles from formula (25), and denote the calculated results $f^{1}(v_{p})$ or $f^{i}(v_{p})$.
- 4) Substitute u¹(y) or uⁱ(y) into Eq. (23) to calculate the saltation trajectories of various sand groups and mark the obtained trajectories as x¹(t) or xⁱ(t), y¹(t) or yⁱ(t), x¹(t) or xⁱ(t) and y¹(t) or yⁱ(t). Put s, f¹(v_p) (or fⁱ(v_p)), x¹(t) (or xⁱ(t)), and y¹(t) (or yⁱ(t)) into formula (26) to find the reaction force of the particles at different heights of wind F_xⁱ(y).
- 5) Substitute $F_x^i(y)$ into Eq. (22) to find the wind speed profile modified by saltation particles $u^{i+1}(y)$, then find the shear stress of the wind acting on the bed surface τ^{i+1} .
- 6) If τ^{i+1} is greater than (smaller than) τ_{ti} , choose the greater (smaller) *s* and repeat Steps 3–6 until $|\tau^{i+1} \tau_{ti}| \leq 0.001$.
- 7) Substitute the values of s, $f^{i}(v_{p})$, $\dot{x}^{i}(t)$ and $\dot{y}^{i}(t)$ calculated according to Step 6 and satisfying $|\tau^{i+1} \tau_{ti}| \le 0.001$ into

formulas (27) and (28) to find the STR and the STR per width considering sand shape changes with height.

Results and Discussion

The relationships of the drag coefficients of six non-spherical sand granules to the Reynolds number are shown in Fig. 2. From the figure, it is obvious that the drag coefficients of non-spherical sand particles are greater than those of spherical granules and that the greater the Reynolds number R_{ep} , the more obvious the difference between the non-spherical and spherical granules.

Fig. 3 shows the relationship between the critical starting friction wind speed of non-spherical sand flow and the Reynolds number. It can been seen from this figure, the critical starting friction wind speed given by the Bagnold is a fixed value when the particle size is invariant, while the critical starting friction wind speed in this paper changes with the particle Reynolds number. And the critical starting friction wind speeds for different shapes of sand particles are different for the same Reynolds number. For example, when the Reynolds number $R_{ep} = 184$, the critical starting wind speed of spherical sand particles is $u_{*t} = 0.176 m/s$ (for D = 0.35 mm) (Fig. 3-b), while the critical starting wind speed of non-spherical particles S₂ is $u_{*t} = 0.273 m/s$ (for D = 0.35 mm) (Fig. 3-b), representing an increase of 35.53% compared to the spherical case. Thus the effect of the sand shape on the critical starting friction wind speed is evident. This is mainly due to the drag coefficient changes with the particle Reynolds number and it is related to the shapes of the particles.

Fig. 4 shows that when the friction velocity is 0.49 m/s and the equivalent diameter is 0.35 mm, the distributions of the drag force are affected by the shapes of the sand particles taking off at an



Figure 5. The sand transport rates of sand particles with different shapes as a function of height. (a) and (b) show the sand transport rates of sand particles with the same equivalent diameter of $d_v = 0.25mm$ as a function of height but at different friction wind velocities. (c), (d) and (e) represent the sand transport rates of sand particles with the same equivalent diameter of $d_v = 0.35mm$ as a function of height but at different friction wind velocities. (c), (d) and (e) wind velocities. Here, the black solid line indicates the spherical sand grains, S_1 and S_2 are two different ellipsoid-shaped sand particles, S_3 represents cube-shaped sand particles, S_4 and S_5 are two types of cylinder-shaped sand particles, and S_6 represents frustum-shaped sand particles. doi:10.1371/journal.pone.0105208.g005



Figure 6. The sand transport rate per width for sand particles of different shape. (a). Comparison of calculated sand transport rates per width for different shapes sand particles with equivalent diameters of $d_v = 0.25mm$ with experimental data. (b). Comparison of calculated sand transport rates per width for sand particles of different shape with equivalent diameters of $d_v = 0.35mm$ with experimental data. The black solid line indicates spherical sand grains, S_1 and S_2 are two different ellipsoid-shaped sand particles, S_3 represents cube-shaped sand particles, S_4 and S_5 are two types of cylinder-shaped sand particles, and S_6 represents frustum-shaped sand particles. doi:10.1371/journal.pone.0105208.q006

initial speed (such as at a horizontal initial velocity of 0.6 m/s and a vertical initial velocity of 0.735 m/s) along the elevation. It can be observed that during the saltation process, the changes in the drag forces of different shapes of sand particles are consistent, first increasing and then decreasing with increasing saltation height during their ascent, while during their descent, the drag forces first decrease and then increase with decreasing saltation height. Thus, the drag force for different shapes of sand particles at the same height can vary greatly. For example, the maximum drag force of spherical sand particles is 2.65×10^{-6} N, while that of non-spherical sand particles S_1 is 7.5×10^{-6} N; the latter is 2.72 times larger than the former.

It can also be observed from Fig. 4 that the drag forces applied to the non-spherical particles whose aspect ratios E are greater than 1.0 are smaller than those applied to the spherical particles, while the drag forces applied to the non-spherical particles whose aspect ratios E are smaller than or equal to 1.0 are larger than those applied to the spherical particles. In addition, the smaller the aspect ratio, the greater the drag force applied to the sand grains and the smaller the maximum height the grains can reach after taking off. Because the drag force suffered by sand grains during their ascent has a vertical downward component, the maximum height they can reach is less than 0.0276 m (the maximum height the freely lifting motion at a vertical initial velocity of 0.735 m/s can achieve is 0.0276 m). The greater the drag force, the greater the vertical downward component and the smaller the maximum height the sand particles can achieve. For example, the drag forces acting upon non-spherical sand particles S_2 and S_5 are smaller than those acting on spherical sand particles; the maximum height the spherical sand particles can reach is 2.15 cm, while the nonspherical particles S_1 can only reach a height of 1.45 cm. The shape of the sand grains therefore has significant impact on the drag force and the maximum height these particles can achieve. Thus, it also has a great impact on the structure of wind-blown sand.

Fig. 5 shows the STR per unit area of different shapes of nonspherical particles with equivalent diameters of 0.25 mm and 0.35 mm as a function of the saltation height. The solid line represents the distribution of the STR per unit area of the spherical sand particles as a function of height. As shown in Fig. 5, the variations in the STR per unit area for different shapes of sand particles with height are equivalent; that is, the STR per unit area decreases with increasing height. However, the STRs per unit area for different shapes of sand particles at the same height vary greatly. In general, the STRs per unit area of the sand particles S_2 and S_5 at various elevations are slightly larger than that of spherical sand particles of equivalent diameter. It can also be observed in Fig. 5 that the maximum heights of the wind sand flow formed by different shapes of sand particles at the same wind speed vary greatly, even several-fold. For example, when the equivalent diameter is 0.25 mm and the friction velocity is 0.49 m/s (Fig. 5-a), the sand transport rates per unit area for the non-spherical particles S_1 , S_2 , S_3 , S_4 , S_5 and S_6 at heights above 13.2 cm, 34.8 cm, 27.6 cm, 30 cm, 37.2 cm and 20.4 cm, respectively, are equal to zero, while that of the spherical sand particles at a height above 34.8 cm is zero. Comparing the nonspherical sand particles S_1 and S_5 , the maximum height of S_5 is 2.82 times larger than the maximum height of S_1 .

Fig. 6 shows the relationship of the STR per width of nonspherical sand particles to the friction wind velocity, among which Fig. 6-a and Fig. 6-b correspond to equivalent particle diameters of 0.25 mm and 0.35 mm, respectively. It is clear from the figures that the STRs per width for the different shapes of sand particles with friction wind speed exhibit similar variations; that is, they increase with increasing friction speed, in agreement with Dong's experimental results [31]. Carneiro et al. pointed out that the midair collisions exert an enormous effect on saltation flux and it depends strongly on restitution coefficient and wind speed [32]. As shown in Fig. 6, the impact of the sand shape on the STR per width is also significant. Under the same conditions, the STR per width for different shapes of sand particles can differ by severalfold. For example, when the equivalent particle size is 0.35 mm and the friction wind speed is 1.016 m/s (Fig. 6-b), the STRs per width of the non-spherical particles S_5 and S_6 are 0.056 kg/m/s and 0.015 kg/m/s, respectively, where the former is 3.73-fold larger than the latter.

Thus, in the discussion of the STR, the sand shape is again an important factor. The main reason is that the air drag forces acting on different shapes of sand particles taking off at the same velocity as well as their maximum heights are different, resulting in large variations in energy when they fall to the bed surface. Thus, the differences in the STR are also significant.

Conclusions

In this article, we established a relationship between the starting friction wind velocity and the sand shape, simulated the STR and STR per width of wind sand flow reaching the steady state considering the sand shape and the distribution of co-velocities at which the sand particles lift off, and calculated the drag forces and the starting friction wind velocities of different shapes of nonspherical sand particles in wind sand flow. The results show that:

(1) The drag coefficients of non-spherical sand particles are greater than those of spherical granules, The greater the Reynolds number is, the more obvious the difference between the nonspherical and spherical particles.

(2) The change in the drag force is consistent for different shapes and sizes of sand particles; i.e., during their ascent, the drag force first increases and then decreases with increasing saltation height, while during their descent, it first decreases and then increases with decreasing saltation height. The drag force to which the sand particles are subjected is inversely proportional to the maximum height they can reach after their takeoff and closely related to the aspect ratio.

References

- Kok JF, Parteli EJR, Michaels TI, Karam DB (2012) The physics of wind-blown sand and dust. Reports on Progress in Physics 75(10): 106901.
- Bagnold RA (1941) The Physics of Blown Sand and Desert Dunes. London: Methuen.
- Zheng XJ (2009) Mechanics of Wind-blown Sand Movements. German: Springer.
- McTainsh G, Livingstone I, Strong C (2013) Fundamentals of Aeolian sediment transport: Aeolian sediments. In: Shroder JF, editor. Treatise on Geomorphology. Oxford: Academic Press. pp. 23–42.
- Dupont S, Bergametti G, Marticorena B, Simoëns S (2013) Modelling saltation intermittency. Journal of Geophysical Research: Atmospheres 118(13): 7109– 7128.
- Kok JF, Renno NO (2009) A comprehensive numerical model of steady-state saltation. Journal of Geophysical Research: Atmospheres 114: D17204.
- Owen PR (1964) Saltation of uniform grains in air. Journal of Fluid Mechanics 20(2): 225–242.
- Ungar JE, Harf PK (1987) Steady state saltation in air. Sedimentology 34(2): 289–299.
- Duran O, Hermann H (2006) Modeling of saturated sand flux. Journal of Statistical Mechanics: Theory and Experiment 2006: P0711.
- Almeida MP, Andrade JS, Herrmann HJ (2006) Aeolian transport layer. Physical review letters 96: 018001.
- Anderson RS, Haff PK (1991) Wind modification and bed response during saltation of sand in air. In: Barndorff-Nielsen OE, Willetts BB, editors. Acta Mechanica Supplementum. German: Springer. pp. 21–51.
- Zhou YH, Guo X, Zheng XJ (2002) Experimental measurement of wind-sand flux and sand transport for naturally mixed sands. Physical Review E 66(2): 021305.
- Huang N, Ren S, Zheng XJ (2008) Effects of the Mid-air Collision on Sand Saltation. Science in China (Series G) 38(3): 260–269.
- Huang N, Wang C, Pan XY (2010) Simulation of acolian sand saltation with rotational motion. Journal of Geophysical Research 115: D22211.
- Dong QS (1997) Physical characteristics of the sand and dust in different deserts of China. Chinese Journal of Radio Science 12(1): 15–25.
- Sagga AMS (1993) Roundness of sand grains of longitudinal dunes in Saudi Arabia. Sedimentary Geology 87(1–2): 63–68.

(3) The critical starting friction wind speed of sand particles is not a constant. It changes with the particle Reynolds number. The critical starting friction wind speeds of different shapes of sand granules are different for the same particle Reynolds number.

(4) The STR per unit area as a function of height display similar variations for sand particles of different shapes that is, it decreases with increasing height. However, the STR per unit area for sand particles of different shapes at the same height varies greatly, as does the maximum height that the sand particles can reach at the same wind velocity, which varies even by several-fold.

(5) The STR per width as a function of the friction wind speed is similar for sand particles of different shapes; that is, it increases with increasing friction wind speed. The effect of the sand shape on the STR per width is significant; under the same conditions, the STR per width for sand particles of different shapes can vary by several-fold.

Author Contributions

Conceived and designed the experiments: ZW SR NH. Performed the experiments: ZW SR. Analyzed the data: ZW SR. Contributed reagents/ materials/analysis tools: ZW. Contributed to the writing of the manuscript: ZW NH.

- Juan JKZ (2009) Roundness in quartz grains from inland and coastal dune sands, Altar Desert, Sonora, Mexico. Boletín de la Sociedad Geológica Mexicana 61(1): 1–12.
- Cheng X, Qi HY, You CF, Zhou F, Xu XC (2003) Analyse of Aeolian Sandy Physical Characteristics in Desert. Journal of Engineering 24(3): 437–440.
- 19. Wu Z (1987) Aeolian Geomorphology (in Chinese). Beijing: Science Press.
- 20. Gao D (1983) Shape of dune sand in desert. World Desert Study 4: 47-49.
- Barton ED (1973) Slip correction factors for nonspherical bodies-I Introduction and continuum flow. Journal of Aerosol Science 4(2): 139–145.
- Gavze E, Shapiro M (1997) Particles in a shear flow near a solid wall: Effect of nonsphericity on forces and velocities. International Journal of Multiphase Flow 23(1): 155–182.
- Yin CG, Lasse R, Søren KK, Sørensen H (2003) Modelling the motion of cylindrical particles in a nonuniform flow. Chemical Engineering Science 58(15): 3489–3498.
- Xu WJ, Che DF, Xu TM (2006) Drag and Lift Forces Acting on Rotational Nonspherical Particles. Journal of Xi'an Jiaotong University 40(3): 298–301.
- Zastawny M, Mallouppas G, Zhao F, Wachem BV (2012) Derivation of drag and lift force and torque coefficients for non-spherical particles in flows. International Journal of Multiphase Flow 39:227–239.
- Blott SJ, Pye K (2006) Particle size distribution analysis of sand-sized particles by laser diffraction: an experimental investigation of instrument sensitivity and the effects of particle shape. Sedimentology 53 (3): 671–685.
- Loth E (2008) Drag of non-spherical solid particles of regular and irregular shape. Power Technology 182(3): 342–353.
- Clift R, Grace JR, Weber ME (1978) Bubbles, Drops and Particles. New York: Academic Press.
- Dong Z, Liu X, Li F (2002) Impact/entrainment relationship in a saltating cloud. Earth Surface Process and landforms 27(6): 641–658.
- Ren S, Huang N (2010) A numerical model of the evolution of sand saltation with consideration of two feedback mechanisms. European Physical Journal E 33(4): 351–358.
- Dong ZB, Liu XP (2004) Experimental investigation of the concentration profile of a blowing sand cloud. Geomorphology 60(3–4): 371–381.
- Carneiro MV, Araujo NAM, Pahtz T, Herrmann HJ (2013) Midair collisions enhance saltation. Phys. Rev. Lett 111: 058001.