

Maps Showing Thermal Maturity of Upper Cretaceous Marine Shales in the Wind River Basin, Wyoming



Pamphlet to accompany
Scientific Investigations Map 3266

COVER. Upper part of the Cody Shale at Eagle Point, Wyo., in the northwestern part of the Wind River Basin. Photograph by R.C. Johnson, 1994.

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By Thomas M. Finn and Mark J. Pawlewicz

Scientific Investigations Map 3266

**U.S. Department of the Interior
U.S. Geological Survey**

U.S. Department of the Interior
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U.S. Geological Survey
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U.S. Geological Survey, Reston, Virginia: 2013

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Suggested citation:

Finn, T.M., and Pawlewicz, M.J., 2013, Maps showing thermal maturity of Upper Cretaceous marine shales in the Wind River Basin, Wyoming: U.S. Geological Survey Scientific Investigations Map 3266, 13 p., 1 sheet, scale 1:500,000, <http://dx.doi.org/10.3133/sim3266>.

ISSN 2329-132X (online)

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Sheet

Sheet with maps showing thermal maturity of upper Cretaceous marine shales in the Wind River Basin, Wyoming.....	link
1. Map showing vitrinite reflectance at the top of the Cody Shale, Bighorn Basin, Wyoming and Montana.	
2. Map showing vitrinite reflectance at the top of the lower shaly member of the Cody Shale.	
3. Map showing vitrinite reflectance at the top of the Mowry Shale.	

Conversion Factors

Inch/Pound to SI

Multiply	By	To obtain
	Volume	
cubic foot (ft ³)	0.02832	cubic meter (m ³)
	Area	
square mile (mi ²)	2.590	square kilometer (km ²)

SI to Inch/Pound

Multiply	By	To obtain
	Length	
millimeter (mm)	0.03937	inch (in.)
	Area	
square meter (m ²)	0.0002471	acre
square kilometer (km ²)	0.3861	square mile (mi ²)

Altitude, as used in this report, refers to distance above Sea Level.

Maps Showing Thermal Maturity of Upper Cretaceous Marine Shales in the Wind River Basin, Wyoming

By Thomas M. Finn and Mark J. Pawlewicz

Introduction

The Wind River Basin is a large Laramide (Late Cretaceous through Eocene) structural and sedimentary basin that encompasses about 7,400 mi² in central Wyoming (fig. 1). The basin is bounded by the Washakie Range, Owl Creek, and southern Bighorn Mountains on the north, the Casper arch on the east and northeast, the Granite Mountains on the south, and the Wind River Range on the west (figs. 1, 2).

Important conventional and unconventional oil and gas resources have been discovered and produced from reservoirs ranging in age from Mississippian through Tertiary (Keefer, 1969; Fox and Dolton, 1989, 1996; De Bruin, 1993; Johnson and others, 1996, 2007). It has been suggested that various Upper Cretaceous marine shales are the principal hydrocarbon source rocks for many of these accumulations (Keefer, 1969; Meissner and others, 1984; Fox and Dolton, 1989, 1996; Johnson and Rice, 1993; Nuccio and others, 1996; and Schelling and Wavrek, 1999, 2001). Numerous source rock studies of various Upper Cretaceous marine shales throughout the Rocky Mountain region by various authors, including Schrayner and Zarrella (1963, 1966, 1968), Nixon (1973), Clayton and Swetland (1977), Merewether and Claypool (1980), Burtner and Warner (1984), Hagen and Surdam (1984), Momper and Williams (1984), Davis (1986), Davis and others (1989), Nuccio (1990a,b), Longman and others (1998), Landon and others (2001), Roberts and others (2004, 2007, 2008), and Finn (2007a, 2010), have led to the conclusion that these rocks have generated, or are capable of generating, oil and (or) gas. With recent advances and success in horizontal drilling and multistage fracture stimulation there has been an increase in exploration and completion of wells in these marine shales in other Rocky Mountain Laramide basins that were traditionally thought of only as hydrocarbon source rocks (Sterling and others, 2009; Sonnenberg, 2011). Important parameters that control hydrocarbon production from shales include: reservoir thickness, amount and type of organic matter, and thermal maturity (Milici, 1993; Curtis, 2002; Passey and others, 2010). The purpose of this report is to present maps and a structural cross section showing levels of thermal maturity, based on vitrinite reflectance (R_o), for Upper Cretaceous marine shales in the Wind River Basin.

Stratigraphy

Figure 3 is a correlation chart showing the stratigraphic nomenclature for Upper Cretaceous marine shales and associated rocks in the Wind River Basin (Finn, 2007b) and correlative units in various localities in the Powder River Basin to the east (Merewether, 1996). The stratigraphic relationships and nomenclature for the Wind River Basin are also illustrated on the regional stratigraphic cross section in figure 4.

Mowry Shale

According to Keefer and Johnson (1997), the Mowry Shale in the Wind River Basin is characterized by two distinct units (fig. 5). The lower part consists of 50 to 150 ft of soft fissile clay-rich shale that has been referred to as the Shell Creek Shale in the adjacent Bighorn and Powder River Basins (Eicher, 1962; Burtner and Warner, 1984). The upper part is composed of 200 to 350 ft of hard brittle siliceous shale, with numerous bentonite beds (Keefer and Johnson, 1997). The shales are dark brown to black, organic-rich, and contain an abundance of fish scales (Burtner and Warner, 1984). The bentonites are gray to tan and range in thickness from less than an inch to as much as 6.5 ft (Byers and Larson, 1979). The top of the Mowry is considered to be the Clay Spur Bentonite Bed; the base is the contact with the underlying Muddy Sandstone (fig. 5). The combined thickness of the lower and upper units ranges from about 220 ft in the southeastern part of the basin to more than 500 ft in the western part. Paleontologic evidence and radiometric dating indicates that the formation is mainly early Late Cretaceous (Cenomanian) in age (Cobban and Kennedy, 1989; Merewether and others, 1997; Obradovich and others, 1996), but the lower part (Shell Creek Shale equivalent) has been dated radiometrically as Early Cretaceous (Albian) by Obradovich and others (1996). For convenience, all of the formation will be discussed with the Upper Cretaceous in this report.

Frontier Formation

The Frontier Formation consists of alternating sandstone, shale, and bentonite that accumulated in marine and marginal marine environments (Keefer, 1972). In the western part of

2 Maps Showing Thermal Maturity of Upper Cretaceous Marine Shales in the Wind River Basin, Wyoming

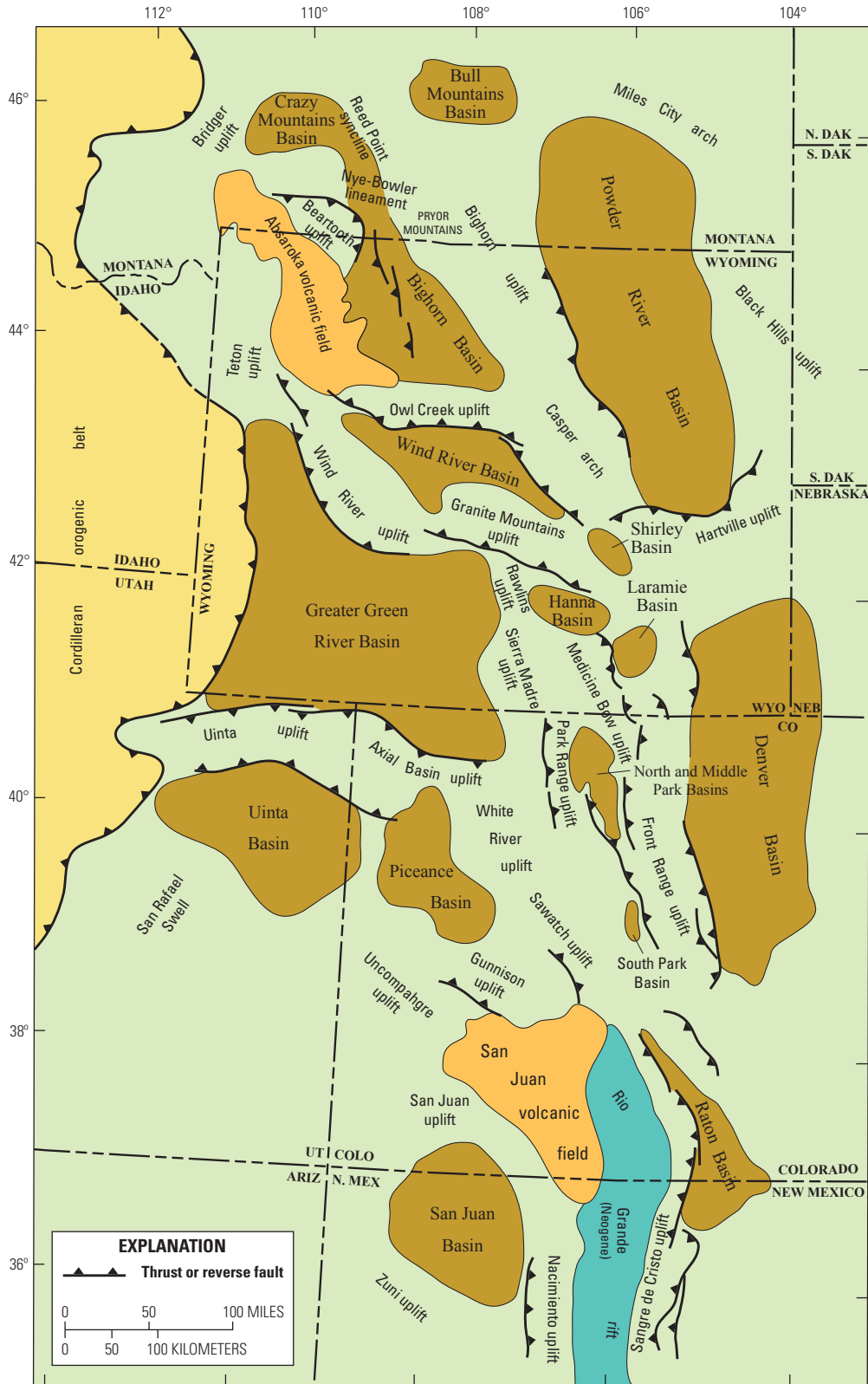


Figure 1. Map of Rocky Mountain region extending from southern Montana to northern New Mexico showing locations of Laramide sedimentary and structural basins (in brown) and intervening uplifts. Modified from Dickinson and others (1988).

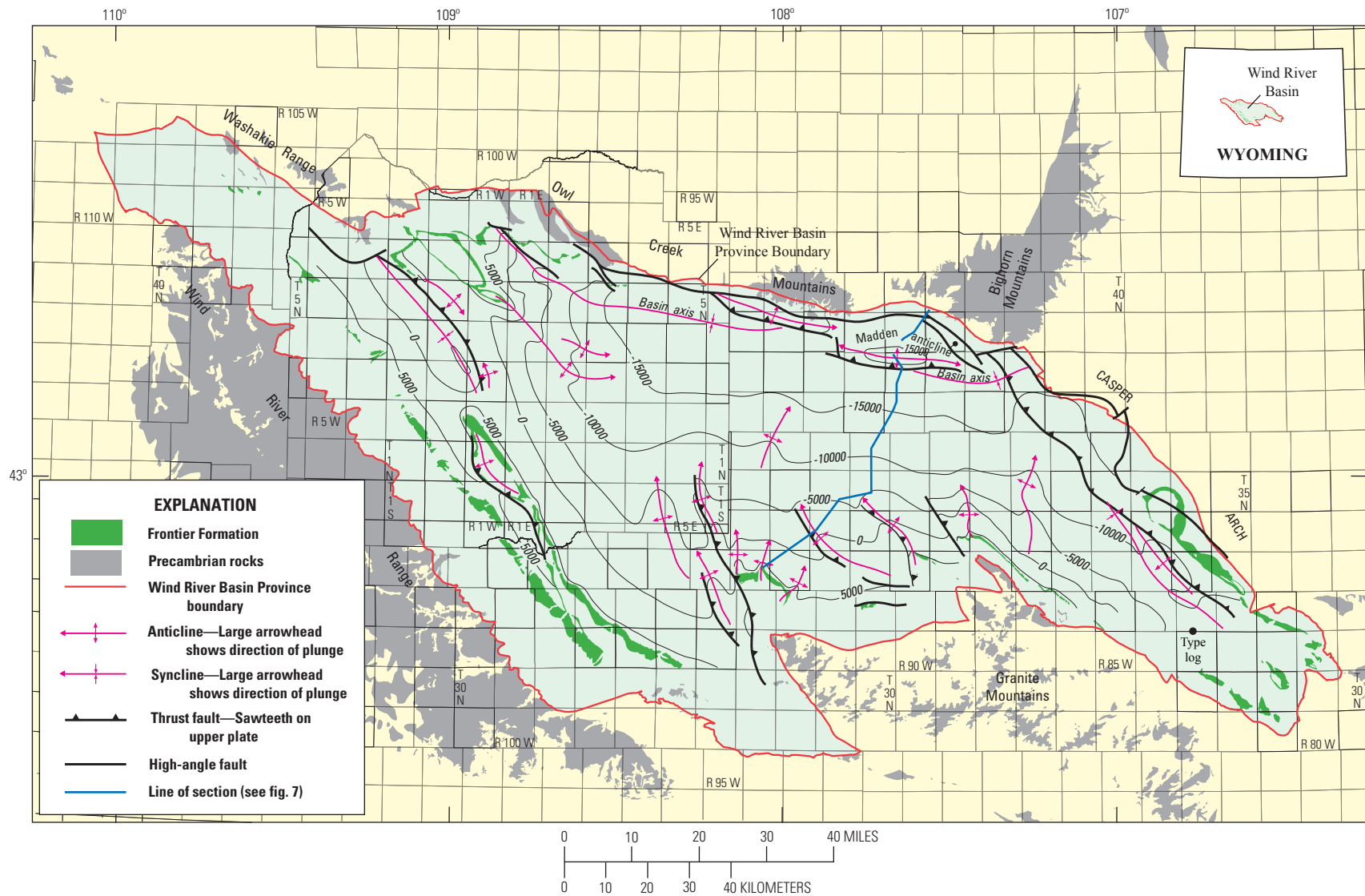


Figure 2. Index map of the Wind River Basin in central Wyoming showing major structural and physiographic features. Structure contours are drawn on top of the Frontier Formation. Contour interval 5,000 ft.

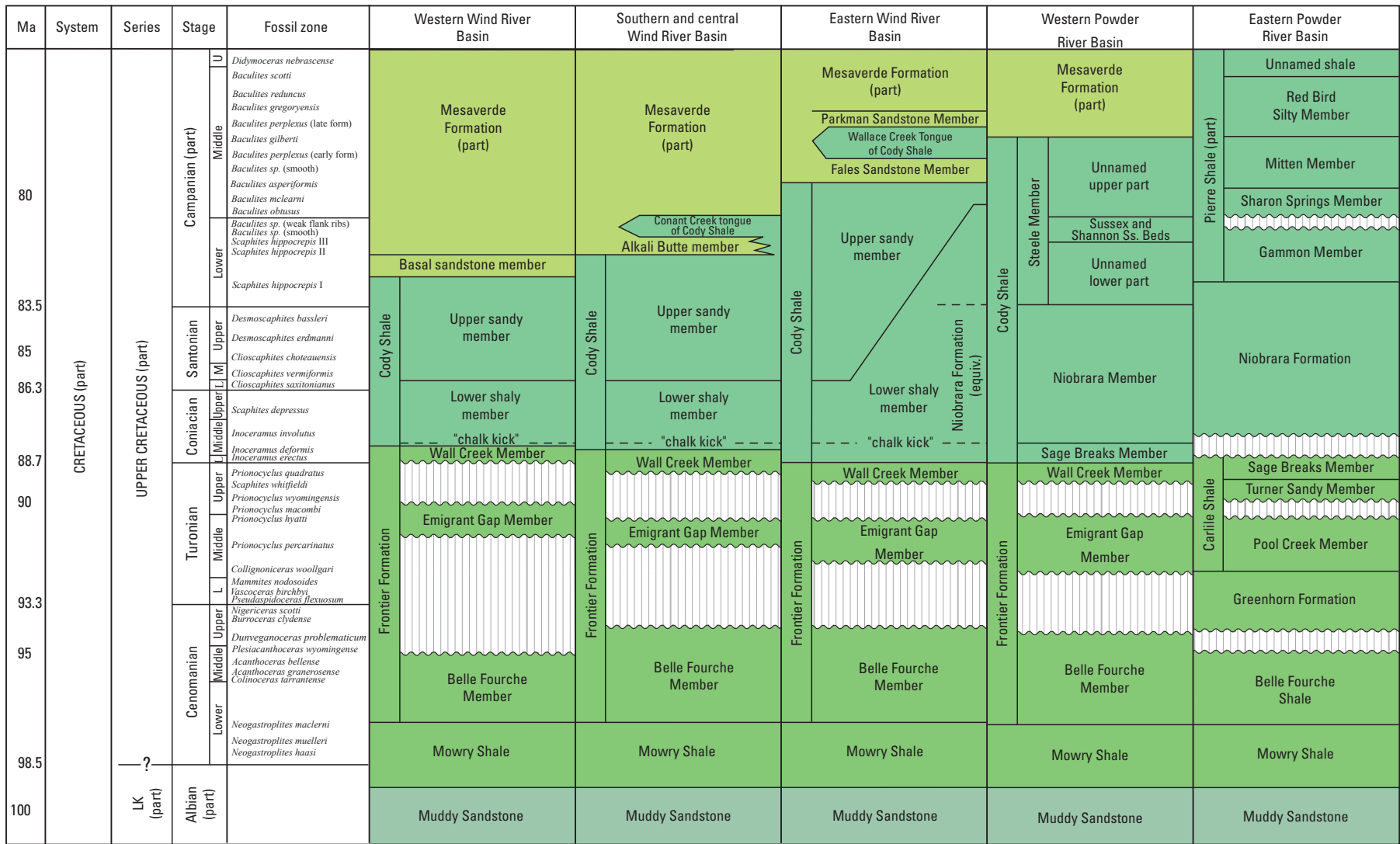


Figure 3. Correlation chart showing stratigraphic relations of lowermost Upper Cretaceous rocks in the Wind River Basin and correlation with equivalent rocks in the Powder River Basin. Modified from Merewether (1996) and Finn (2007b). Radiometric ages and fossil zones are from Obradovich (1993), and Merewether and others (1997). LK, Lower Cretaceous.

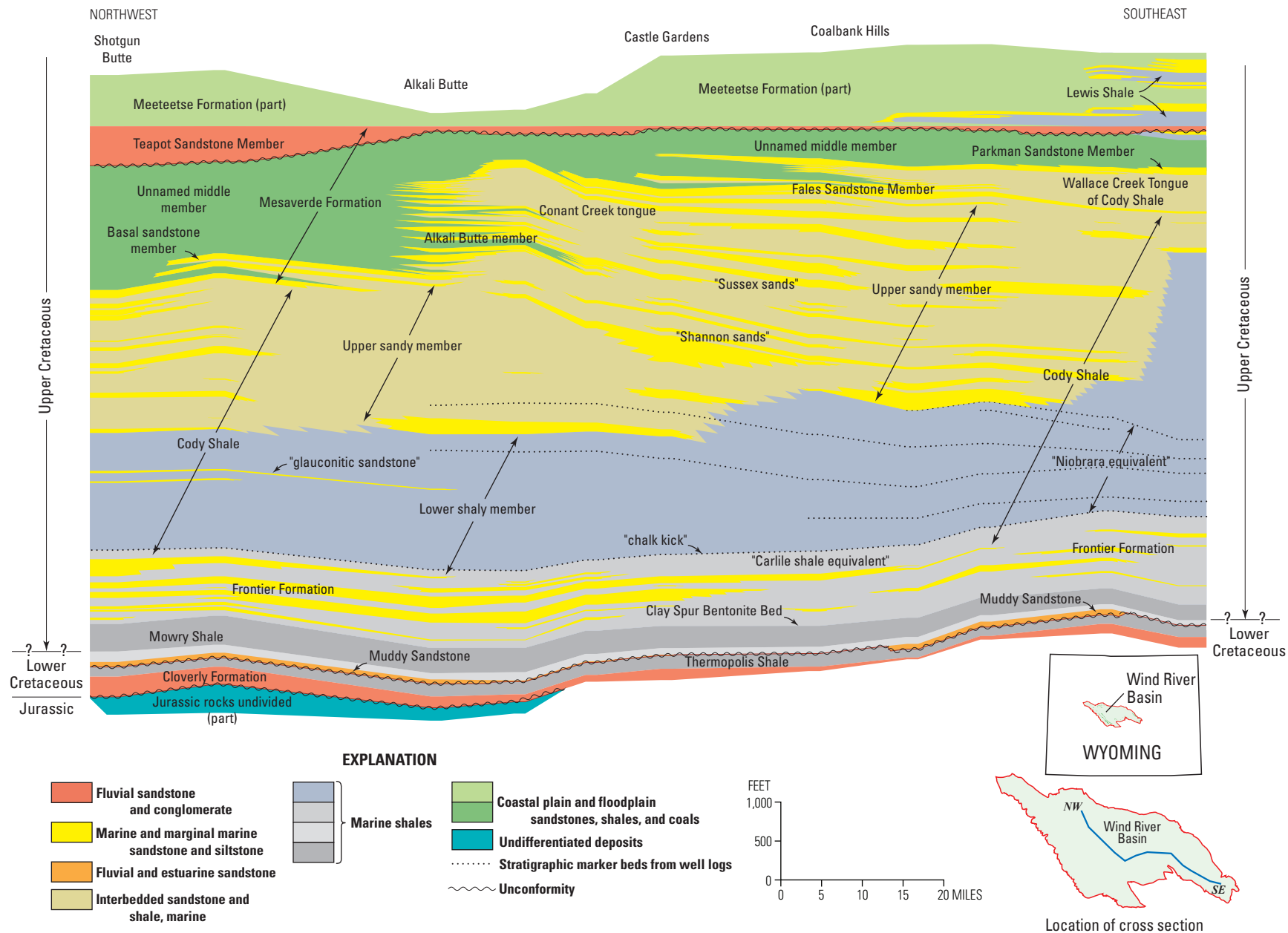


Figure 4. Regional northwest-southeast stratigraphic cross section of Cretaceous rocks (part) in the Wind River Basin. Modified from Finn (2007b).

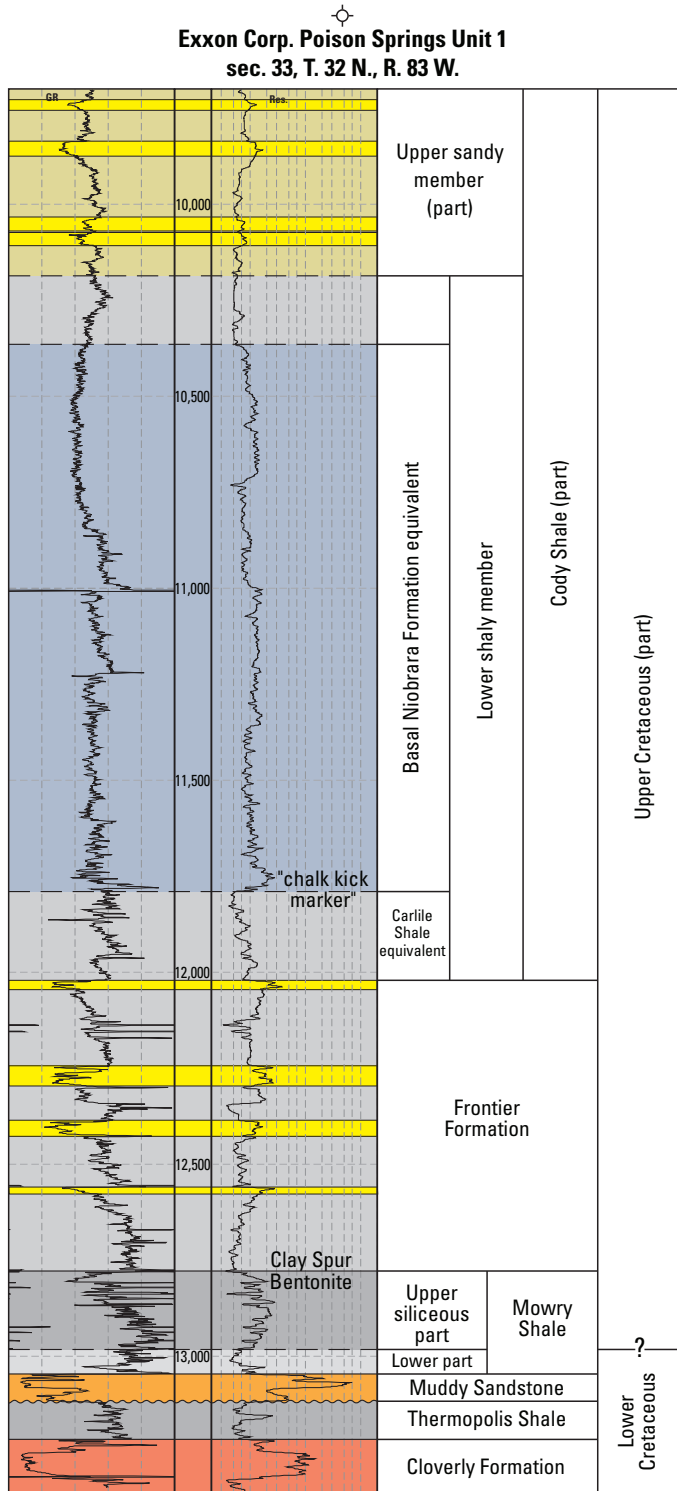


Figure 5. Type log of Exxon Corp. Poison Springs Unit 1, of the lower and lowermost Upper Cretaceous rocks in the southeastern part of the Wind River Basin. Sandstones and conglomerates of predominantly fluvial origin are shown in red; marine and marginal marine sandstones and siltstones, in yellow; marine shales, in various shades of gray; estuarine and fluvial sandstones in orange; interbedded sandstone and shale in tan. Location shown in figure 2.

the basin, the Frontier also contains some nonmarine rocks, including minor coal deposits. Sandstones generally dominate in the upper part of the formation, whereas the lower part is mostly marine shale. Many individual sandstones are blanket-like and can be traced over several miles, but locally may be discontinuous and pinch out into marine shale. The shales are generally sandy or silty and vary from gray to black in color. The formation is Cenomanian to Coniacian in age (Merewether and others, 1997) and ranges in thickness from 400 to 800 ft.

Cody Shale

The Cody Shale consists of marine shale, siltstone, and sandstone that were deposited during a major transgressive-regressive cycle that Kauffman (1977) referred to as the “Niobrara Cyclothem.” According to Johnson and others (2007) the formation ranges in thickness from about 3,250 ft in the western part of the basin to more than 5,500 ft in the eastern part. The lower and upper contacts are conformable and interfinger with the underlying Frontier and overlying Mesaverde Formations (fig. 4). Four members are recognized in ascending order: (1) the lower shaly member (Thompson and White, 1954; Yenne and Pippingos, 1954; and Keefer and Troyer, 1964), (2) the upper sandy member (Thompson and White, 1954; Yenne and Pippingos, 1954; Keefer and Troyer, 1964), (3) the informally named “Conant Creek tongue” (Szmajter, 1993), and (4) the Wallace Creek Tongue (Barwin, 1961) (fig. 4). The age of the Cody ranges from Coniacian to middle Campanian (Keefer, 1972).

The unnamed lower shaly member is about 1,750 ft thick in the western part of the Wind River Basin and thickens to the east to more than 2,250 ft (Johnson and others, 2007). It is comprised of gray to black shale and bentonite deposited in an offshore marine environment. A persistent geophysical-log marker in the lower 100–300 ft of the lower member (fig. 5), referred to as the “chalk kick” by Keefer (1972), can be traced in the subsurface throughout most of the basin; it separates noncalcareous shales in the lower part of the shaly member from overlying calcareous shales. Based on sample descriptions from well cuttings, the calcareous interval is about 1,000 ft thick in the eastern part of the basin where Merewether and others (1977a, b) correlated it with the Niobrara Formation in the Powder River Basin (fig. 3). They (Merewether and others, 1977a, b) correlated the interval below the “chalk kick” marker and the top of the uppermost sandstone in the Frontier Formation with the Carlile Shale in the Powder River Basin (figs. 4, 5). In the eastern part of the Wind River Basin the Niobrara-equivalent rocks are of Coniacian to latest Santonian or earliest Campanian age. To the west, the upper part of the interval grades into the sandy and shaly facies of the upper sandy member of the Cody Shale and only the basal Niobrara strata of the Powder River Basin are represented by the lower shaly member (fig. 4). In the western part of the Wind River Basin, rocks assigned to the shaly member are Coniacian to early Santonian in age (Yenne and Pippingos, 1954; Keefer, 1972; Finn, 2007b).

The upper sandy member ranges in thickness from about 1,800 ft in the western part of the basin to about 3,500 ft in the southeastern part (Johnson and others, 2007). It consists of light to medium gray sandstones and tan and gray shales. Dunleavy and Gilbertson (1986) referred to sandstones in the upper part of the member in the northern part of the basin as the “Sussex” and “Shannon” sandstone beds, which, according to them were deposited “as a near-shore bar complex along the edge of a delta.” Like sandstones in the underlying Frontier Formation, many individual sandstones in the upper part of the Cody are blanket-like and can be traced over several miles before pinching out into marine shale. The upper sandy member becomes less distinct in the southeastern part of the basin where it grades laterally into more shaly facies (Finn, 2007b) (fig. 4).

The Conant Creek tongue of the Cody Shale, informally named by Szmajter (1993), trends north-south across the central part of the basin, and is separated from the upper sandy member by an eastward thinning clastic wedge of marginal marine and nonmarine rocks informally referred to as the Alkali Butte member of the Mesaverde Formation by Hogle and Jones (1991) (fig. 4). The Conant Creek tongue is generally 400 to 900 ft thick, but thins to zero where it grades westward into nonmarine rocks of the Mesaverde Formation (fig. 4).

The Wallace Creek Tongue of the Cody Shale occupies the eastern and southeastern parts of the Wind River Basin and is stratigraphically higher and younger than the Conant Creek tongue to the west. It is a westward-thinning tongue of marine shale that separates the Fales Sandstone Member at the base of the Mesaverde Formation from the upper part of the formation (Barwin, 1961) (fig. 4). The Wallace Creek Tongue is nearly 500 ft thick in the southeastern corner of the basin and thins to zero in the northern part of the Coalbank Hills, where it grades into the main part of the Mesaverde Formation (fig. 4).

Thermal Maturity Maps

Maps showing the levels of thermal maturity based on vitrinite reflectance (R_o) for Cretaceous marine shales in the Wind River Basin are shown as three separate maps on the map sheet. They were constructed using R_o data published by Johnson and others (1991), Barker and Crysdale (1993), Nuccio and others (1993, 1996), Pawlewicz (1993), Finn and others (2006), and Pawlewicz and Finn (2013), from samples collected from wells drilled for oil and gas exploration and production, and from outcrops. For wells that reported R_o measurements over a range of depths, these data were plotted on a log (R_o) versus depth plot for each well and a visual best-fit line was drawn through the data (for example, see fig. 6). Data points that appeared to be anomalously high or low with respect to the main trend were disregarded. In some cases, data from closely associated wells were combined into a composite log (R_o) versus depth plot. For a single well or outcrop location where few R_o measurements were available over a short depth range or stratigraphic interval, the available data were averaged and the calculated R_o was plotted for that location.

The three maps presented on the map sheet were constructed to show the thermal maturity of Upper Cretaceous marine shales in the Wind River Basin at: (1) top of the Cody Shale, (2) top of the lower shaly member of the Cody Shale (Niobrara equivalent), and (3) the top of the Mowry Shale.

- Because the contact between the Cody Shale and the overlying Mesaverde Formation rises stratigraphically to the east (figs. 3, 4), it was necessary to change the contoured interval from west to east across map 1 (top of Cody Shale). In the western part of the basin, the contours are drawn at the base of the basal sandstone member of the Mesaverde Formation. In the central part, they are drawn at the top of the Conant Creek tongue of the Cody Shale, and in the eastern and southeastern parts they are drawn at the top of the Wallace Creek Tongue (figs. 3 and 4; and map 1 on the sheet).
- The R_o map for the top of the lower shaly member (map 2 on the map sheet) is contoured at the contact between the lower shaly member and the overlying sandy member in the western and central parts of the basin. However, in the eastern and southeastern parts of the basin, these subdivisions are less distinct and the contours are drawn at the top of the Niobrara-equivalent strata based on well log correlations from the Powder River Basin by Merewether and others (1977a, b) (figs. 3 and 4; and map 2 on the map sheet).
- The R_o map for the Mowry Shale (map 3 on the map sheet) is drawn on the Clay Spur Bentonite Bed (fig. 4).

For all three contoured horizons, the maps show gradual increases in maturity from the margins of the Wind River Basin northward into the deeper parts of the basin, which reflect its structural configuration (fig. 2). In each case, R_o values range from less than 0.6 percent along the shallow margins of the basin increasing to greater than 3.0 to 4.0 percent in the deeper parts along the main basin trough and in the area of the Madden anticline near the northern margin of the basin (maps 1–3 on the map sheet). These thermal maturity trends show that large areas of the basin are thermally mature ($R_o >$ about 0.6 percent) to overmature ($R_o >$ about 1.35 percent) with respect to oil generation based on parameters defined by Peters and Cassa (1994). A structural cross section extending from the Madden anticline south to the southern margin of the basin illustrates the relation between present-day structural configuration of the basin and the levels of thermal maturity (fig. 7). The cross section shows that the isorefectance lines are nearly horizontal and cut across present-day structures, indicating that the thermal maturation developed mainly after basin formation and deformation during the Laramide orogeny and subsequent accumulation of basin fill (Bustin and others, 1983; Barker and Crysdale, 1993; Nuccio and others, 1996).

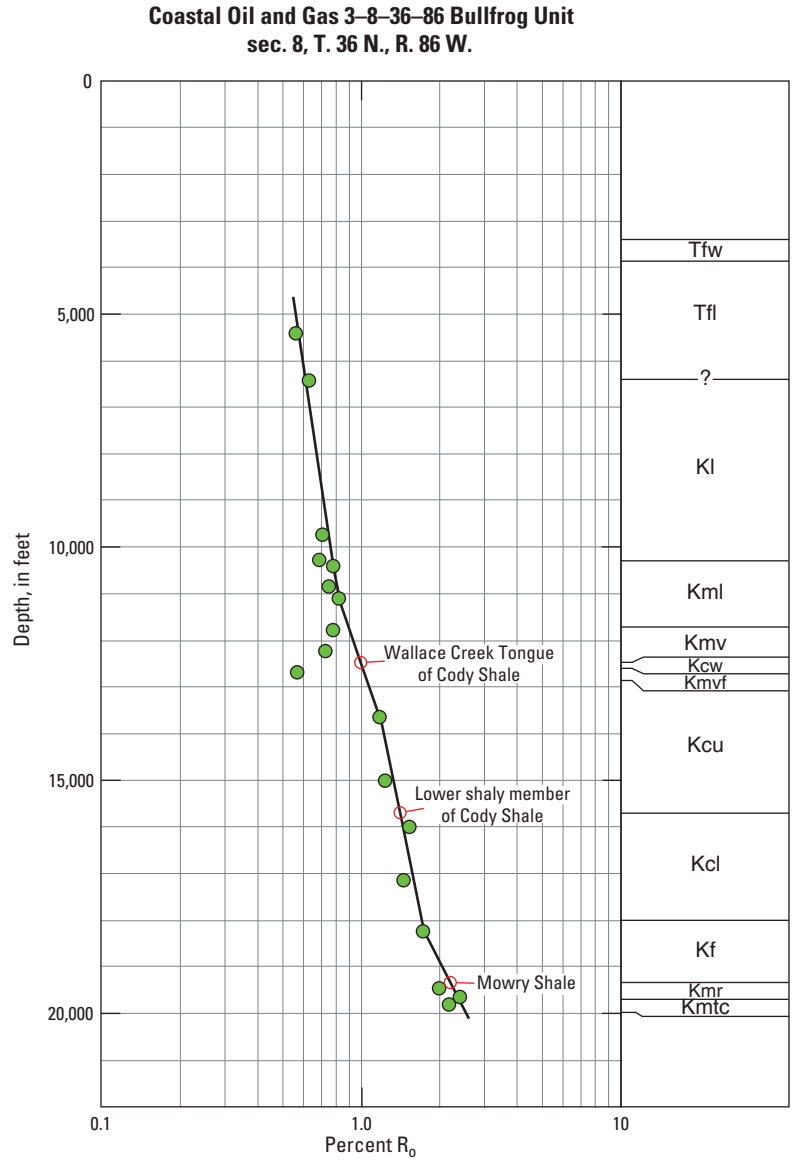
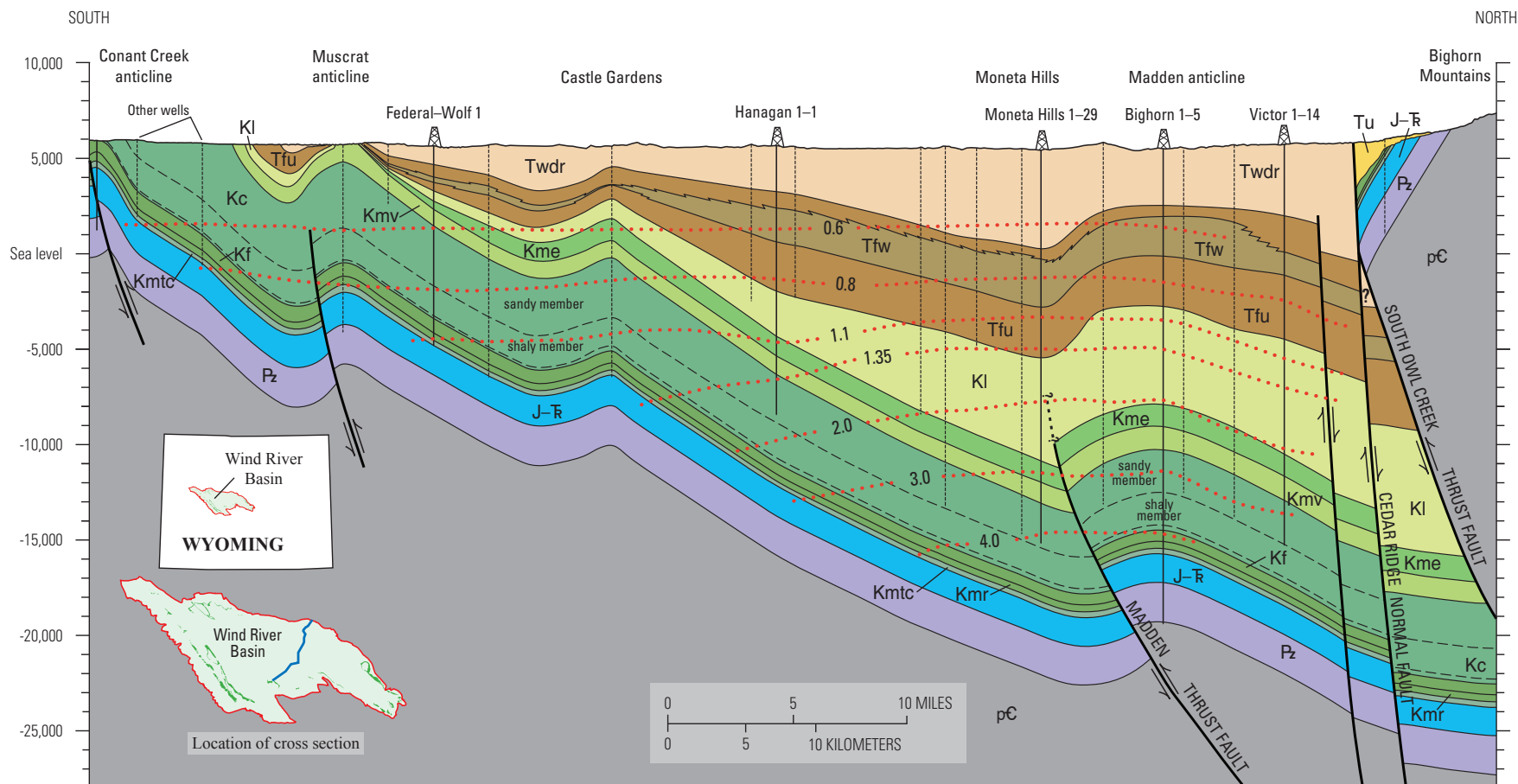


Figure 6. Log (R_0) versus depth plot for the Coastal Oil and Gas 3-8-36-86 Bullfrog Unit well. Tfw, Waltman Shale Member of Fort Union Formation; Tfl, lower unnamed member of Fort Union Formation; Kl, Lance Formation; Kml, Meeteetse Formation and Lewis Shale, undivided; Kmv, Mesaverde Formation; Kcw, Wallace Creek Tongue of Cody Shale; Kmvf, Fales Member of Mesaverde Formation; Kcu, upper sandy member of Cody Shale; Kcl, lower shaly member of Cody Shale; Kf, Frontier Formation; Kmr, Mowry Shale; Kmtc, Muddy Sandstone, Thermopolis Shale, and Cloverly Formation, undivided.



VERTICAL EXAGGERATION X 4

EXPLANATION

- | | | |
|---|---|---|
| Tu Post Lower Eocene rocks | Kme Upper Cretaceous—Meeteetse Formation | Kmtc Lower Cretaceous—Muddy Sandstone, Thermopolis Shale, and Cloverly Formation |
| Twdr Lower Eocene—Wind River Formation | Kmv Upper Cretaceous—Mesaverde Formation | J-R Jurassic to Triassic sedimentary rocks—Undivided |
| Tfu Paleocene—Fort Union Formation | Kc Upper Cretaceous—Cody Shale | Pz Paleozoic sedimentary rocks—Undivided |
| Tfw Paleocene—Fort Union Formation, Waltman Shale Member | Kf Upper and Lower Cretaceous—Frontier Formation and Mowry Shale | pC Precambrian crystalline rocks |
| Kl Upper Cretaceous—Lance Formation | Drilling well for hydrocarbon exploration | 0.60 Isorefectance line (in percent) |
| | | Contact—Dashed where inferred |
| | | Fault |

Figure 7. South-north structural cross section across the central part of the Wind River Basin, Wyoming, showing relation between present-day structure and levels of thermal maturity based on vitrinite reflectance. Line of section shown in fig. 2 and on map sheet. The part of the line of section north of the Madden anticline is modified from Ray and Keefer (1985).

Acknowledgments

The manuscript and plates benefited from reviews by Dick Keefer, Phil Nelson, Dave Ferderer, and Tom Judkins, and their suggestions and comments are greatly appreciated.

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Publishing support provided by:
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