

Prepared in cooperation with the Town of Danby and the Tompkins County Planning Department

Geohydrology and Water Quality of the Stratified-Drift Aquifers in Upper Buttermilk Creek and Danby Creek Valleys, Town of Danby, Tompkins County, New York



Scientific Investigations Report 2015–5138

U.S. Department of the Interior U.S. Geological Survey





**Cover.** *A*, Panorama of the valley from a location near Comfort Road, Danby, N.Y., looking southwest. *B*, Signage for the town offices and public library of the Town of Danby. *C*, U.S. Geological Survey (USGS) hydrologist holding soil sample from test well TM2806 (Whitehawk Lane, Danby). *D*, The Danby town hall. All photographs copyright Ted Crane (http://tedcrane.com/), used with permission.

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## **Conversion Factors**

Multiply	Ву	To obtain				
	Length					
foot (ft)	0.3048	meter (m)				
mile (mi)	1.609	kilometer (km)				
	Area					
square foot (ft <sup>2</sup> )	0.09290	square meter (m <sup>2</sup> )				
square foot (ft <sup>2</sup> )	0.00000036	square mile (mi <sup>2</sup> )				
square inch (in <sup>2</sup> )	6.452	square centimeter (cm <sup>2</sup> )				
square mile (mi <sup>2</sup> )	2.590	square kilometer (km <sup>2</sup> )				
square mile (mi <sup>2</sup> )	27,878,400	square feet (ft <sup>2</sup> )				
	Volume					
gallon (gal)	0.003785	cubic meter (m <sup>3</sup> )				
gallon (gal)	3.785	liter (L)				
million gallons (Mgal)	3,785	cubic meter (m <sup>3</sup> )				
cubic foot (ft <sup>3</sup> )	0.02832	cubic meter (m <sup>3</sup> )				
Flow rate						
foot per year (ft/yr)	0.3048	meter per year (m/yr)				
cubic foot per second (ft <sup>3</sup> /s)	0.02832	cubic meter per second (m <sup>3</sup> /s)				
cubic foot per second (ft <sup>3</sup> /s)	86,400	cubic foot per day (ft <sup>3</sup> /d)				
cubic foot per second (ft <sup>3</sup> /s)	31,536,000	cubic foot per day (ft <sup>3</sup> /yr)				
cubic foot per second (ft <sup>3</sup> /s)	646,317	million gallons per day (Mgal/d)				
cubic foot per second (ft <sup>3</sup> /s)	23,592,505	million gallons per year (Mgal/yr)				
cubic foot per second per square mile [(ft <sup>3</sup> /s)/mi <sup>2</sup> ]	0.01093	cubic meter per second per square kilometer [(m <sup>3</sup> /s)/km <sup>2</sup> ]				
cubic foot per day ( $ft^3/d$ )	0.02832	cubic meter per day $(m^3/d)$				
cubic foot per year (ft <sup>3</sup> /yr)	0.00000748	million gallons per year (Mgal/yr)				
gallon per minute (gal/min)	0.06309	liter per second (L/s)				
gallon per day (gal/d)	0.003785	cubic meter per day $(m^3/d)$				
million gallons per day (Mgal/d)	0.04381	cubic meter per second $(m^3/s)$				
million gallons per day (Mgal/yr)	1.547	cubic foot per second ( $ft^3/s$ )				
million gallons per day per square mile [(Mgal/d)/mi <sup>2</sup> ]	1,461	cubic meter per day per square kilometer [(m <sup>3</sup> /d)/km <sup>2</sup> ]				
inch per year (in/yr)	25.4	millimeter per year (mm/yr)				
	Radioactivity					
picocurie per liter (pCi/L)	0.037	becquerel per liter (Bg/L)				
	Hydraulic conductivity					
foot per day (ft/d)	0.3048	meter per day (m/d)				
	Hydraulic gradient	· · · /				
foot per mile (ft/mi)	0.1894	meter per kilometer (m/km)				

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit

(°F) as °F = 
$$(1.8 \times °C) + 32$$
.

Temperature in degrees Fahrenheit (°F) may be converted to degrees Celsius

## **Supplemental Information**

Specific conductance is given in microsiemens per centimeter at 25 degrees Celsius ( $\mu$ S/cm at 25 °C).

Concentrations of chemical constituents in water are given in either milligrams per liter (mg/L) or micrograms per liter ( $\mu$ g/L).

## Datum

Vertical coordinate information is referenced to the North American Vertical Datum of 1988 (NAVD 88) where light detection and ranging (lidar) and Digital Elevation Model (DEM) data were used, and to the National Geodetic Vertical Datum of 1929 (NGVD 29) elsewhere.

Horizontal coordinate information is referenced to North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above the vertical datum.

## **Abbreviations**

CFCL	USGS Chlorofluorocarbon Laboratory
EPA	U.S. Environmental Protection Agency
HA	EPA Health Advisory
HVSR	horizontal-to-vertical spectral ratio
lidar	light detection and ranging
MCL	EPA Maximum Contaminant Level
NAD 83	North American Datum of 1983
NAVD 88	North American Vertical Datum of 1988
NWQL	USGS National Water Quality Laboratory
NYSDEC	New York State Department of Environmental Conservation
NYSDOH	New York State Department of Health
SDWS	EPA Secondary Drinking-Water Standards
SMCL	Secondary Maximum Contaminant Level
USGS	U.S. Geological Survey

# Geohydrology and Water Quality of the Stratified-Drift Aquifers in Upper Buttermilk Creek and Danby Creek Valleys, Town of Danby, Tompkins County, New York

By Todd S. Miller

## Abstract

In 2006, the U.S. Geological Survey, in cooperation with the Town of Danby and the Tompkins County Planning Department, began a study of the stratified-drift aquifers in the upper Buttermilk Creek and Danby Creek valleys in the Town of Danby, Tompkins County, New York. In the northern part of the north-draining upper Buttermilk Creek valley, there is only one sand and gravel aquifer, a confined basal unit that overlies bedrock. In the southern part of upper Buttermilk Creek valley, there are as many as four sand and gravel aquifers, two are unconfined and two are confined. In the south-draining Danby Creek valley, there is an unconfined aquifer consisting of outwash and kame sand and gravel (deposited by glacial meltwaters during the late Pleistocene Epoch) and alluvial silt, sand, and gravel (deposited by streams during the Holocene Epoch). In addition, throughout the study area, there are several small local unconfined aquifers where large tributaries deposited alluvial fans in the valley.

The principal sources of recharge to the unconfined aguifers in the study area include direct infiltration of precipitation (rain and snowmelt) at land surface, unchanneled surface runoff from adjacent hillsides that seeps into the aquifer along the edges of the valley, groundwater inflow from adjacent till and bedrock that enters the aquifer along the sides of the valley, and seepage loss from uplandtributary streams where they flow over their alluvial fans in the valley. The percentages of all sources of recharge to the contiguous unconfined aquifer in Danby Creek valley include 16 percent from precipitation that falls directly over the aquifer, 55 percent from unchanneled surface runoff and groundwater inflow from hillsides, and 29 percent from losing tributary streams that cross the aquifer. The total annual recharge to the contiguous unconfined aquifer is 2.56 cubic feet per second (604 million gallons per year).

The principal sources of recharge to the confined aquifers include precipitation that falls directly on the surficial confining unit, which then slowly flows vertically downward through the fine-grained sediments and enters the confined aquifer, and groundwater inflow from till and bedrock that borders the aquifer along adjacent hillsides and at the bottom of the valley. In addition, there is substantial amounts of recharge to the confined aquifers where the confining units are locally absent (forming windows) and where parts of the confining units consist of sediments of low to moderate permeability (forming a semiconfining layer).

In the northern part of the study area (upper Buttermilk Creek valley), groundwater in the stratified-drift aquifers discharges to (1) domestic and commercial wells; (2) Buttermilk Creek in the area near the northern town border, and (3) and a small unnamed stream in a ravine in Buttermilk State Park just north of the town border. In the southern part of the study area (Danby Creek valley), groundwater discharges (1) to domestic, commercial, and farm wells; (2) to Danby Creek; (3) to a large wetland in the central parts of Danby Creek valley; and (4) as losses because of plant uptake and evaporation. About 300 people depend on groundwater from the upper Buttermilk Creek and Danby Creek stratifieddrift aquifer system.

An unconfined surficial aquifer about 8,000 feet (ft) long and as much as 800 ft wide, with a saturated thickness of about 20 ft, occupies the lower (southeastern most) 8,000 ft of Danby Creek valley within the Town of Danby. However, because the aquifer is thin, the volume of water stored in the aquifer is small and the potential for induced recharge from Danby Creek during summer periods of low flow is also small, an array of wells would probably be needed to provide sustainable continuous amount of water to large water users such as municipalities and industries. Additional data and a groundwater flow model would be required to estimate sustainable withdrawal from the confined aquifers in upper Buttermilk Creek valley.

Well data from water-well drillers through 2012 indicate that the confined aquifers in upper Buttermilk Creek valley are thin (typically about 10 feet thick) and the reported well-yield data suggest these aquifers may not be capable of supplying sufficient water to meet the needs of municipalities and

industries. However, additional geohydrologic data leading to calibration of a groundwater flow model would be needed to properly evaluate the long-term (multiple years) potential yield of the confined aquifer system in upper Buttermilk Creek valley and of the unconfined aquifer in Danby Creek valley.

During 2007–10, groundwater samples were collected from 13 wells including 7 wells that are completed in the confined sand and gravel aquifers, 1 well that is completed in the unconfined aquifer, and 5 wells that are completed in the bedrock aquifers. Calcium dominates the cation composition and bicarbonate dominates the anion composition in most groundwater. Water quality in the study area generally meets state and Federal drinking-water standards but concentrations of some constituents exceeded the standards. The standards that were exceeded include sodium (3 samples), dissolved solids (1 sample), iron (3 samples), manganese (8 samples), and arsenic (1 sample).

## Introduction

In 2000, the U.S. Geological Survey published a map depicting the extent of stratified-drift (that is, sand and gravel) aquifers in Tompkins County (Miller, 2000), which was a reconnaissance of the area. In 2000-02, the U.S. Geological Survey and the Tompkins County Planning Department used the map to plan studies of the stratified-drift aquifers in more detail. These detailed studies would provide geohydrologic data for town and county planners to develop a strategy to manage and protect their water resources. In developing the plan, a list of a 17 stratified-drift aquifers was compiled and a program was developed to study these aquifers for 20 years. The extent of the aquifers was based mostly on natural hydrologic boundaries, but in some cases, political boundaries were used as well; In order to have manageablesized study areas, the aquifer extents were about 3 to 5 miles (mi) each. During 2006–12, the upper Buttermilk Creek and Danby Creek valleys (fig. 1) were the third aquifer study to be investigated of the county-wide program. The aquifer was investigated by the U.S. Geological Survey, in cooperation with the Tompkins County Planning Department and the Town of Danby.

Evaluation, development, and protection of these aquifers require information on the aquifer framework (the threedimensional extent and distribution of aquifers and confining units), sources of recharge, discharge areas, direction of groundwater flow, and water quality. Samples were collected from wells to characterize the chemical quality of groundwater and to determine whether the water is potable. Groundwater samples were collected from wells that are finished in the confined and unconfined aquifers as well as several wells that are finished in bedrock. A comparison of water quality was made to determine whether there are differences between the sand and gravel aquifers and bedrock aquifers. In addition, stream samples were collected to characterize the chemical quality of surface water under base-flow conditions (when the flow is mostly from groundwater discharging into stream channels) and to determine whether there are similarities in water quality between surface water and groundwater.

### Purpose and Scope

This report describes the geohydrology of the stratifieddrift aquifers in upper Buttermilk Creek and Danby Creek valleys in the Town of Danby, Tompkins County. Specifically, the report contains information that details the study results: (1) the geology of the study area including the geologic framework of the aquifer system; (2) the groundwater-flow system, including information about water levels, groundwater and surface-water interaction, direction of groundwater flow, and recharge and discharge conditions; and (3) the water quality, including information about concentrations of common inorganic ions, nutrients, and trace elements. The report also includes geohydrologic sections; maps and diagrams depicting well locations, geology, groundwater levels, and direction of groundwater flow; and tables of well records and water-quality data.

### **Description of Study Area**

The study area is in the Appalachian Plateau physiographic province (fig. 2). The plateau is characterized by hills and valleys that resulted from millions of years of dissection by initially south-flowing streams that were captured by advancing northward-flowing streams. Those streams drained into Lake Ontario (fig. 2). The hills and valleys have been subsequently modified by several glaciations.

The stratified-drift aquifer system in the upper Buttermilk Creek and Danby Creek valleys is one of the 17 stratified-drift aquifers in Tompkins County that were mapped by the USGS (Miller, 2000). The stratified-drift aquifer system in the upper Buttermilk Creek and Danby Creek valleys is one of two stratified-drift aquifer systems in the Town of Danby—the other aquifer system is in the Cayuga Inlet valley in the southwest part of the town (figs. 1 and 3). However, the aquifer system in the upper Cayuga Inlet valley is not part of this investigation.

The study area is 7.5 mi long and encompasses 4.7 square miles (mi<sup>2</sup>) in the upper Buttermilk Creek and Danby Creek valleys (fig. 3). Buttermilk Creek flows northward in a valley until it cascades over several waterfalls through Buttermilk Creek gorge before it joins lower Cayuga Inlet trough that is 2.7 mi north of the Town of Danby (fig. 3). Cayuga Inlet then drains into Cayuga Lake at Ithaca (fig. 3). Southeastward flowing Danby Creek is a headwater tributary of the Susquehanna River Basin (fig. 3).

The study area has moderate relief with altitudes ranging from as low as 965 ft above North American Vertical Datum



- Salmon Creek/Myers Point/Locke Creek
- Lower Sixmile Creek and Willseyville Creek valleys (Town of Caroline)

Figure 1. Location of 17 stratified-drift (unconsolidated) aquifers in Tompkins County, New York.



**Figure 2.** Physiographic features of New York and location of the upper Buttermilk Creek and Danby Creek valleys study area in Tompkins County, New York.

of 1988 (NAVD 88) in the channel where Buttermilk Creek enters the Buttermilk Creek gorge in the north part of the study area to about 1,800 ft on the highest hilltops (fig. 4). Buttermilk Creek and Danby Creek valleys are relatively highaltitude valleys that are 100 to 500 ft higher than the Cayuga Inlet trough to the west and north, and the lower Sixmile Creek and Willseyville Creek trough to the east (figs. 3 and 4).

## **Methods of Investigation**

Existing data were compiled and new data were collected for this study. Existing data included well records, geologic reports, and soil maps. New data included geologic mapping, test drilling, seismic surveys, groundwater-level measurements, and surface-water and groundwater-quality sampling.

## Geologic Data

This report includes geologic mapping that modifies the geology published by Muller and Cadwell (1986). In addition to interpretation from topographic maps and orthophotographs, soil maps (Neeley, 1961) were used in geologic interpretation.

Geologic data were also collected by seismic-refraction and horizontal-to-vertical spectral ratio (HVSR) seismic surveys to supplement data from test drilling. The seismic surveys were used to determine the thickness of the unconsolidated deposits and depth to bedrock.

## Well Inventory, Test Drilling, Water-Level Measurements, and Levels

Well records (140 records, appendix 1) were collected and compiled for wells in the Town of Danby. Sources of



**Figure 3.** Location of the stratified-drift aquifers in the Town of Danby, the Susquehanna River and St. Lawrence River drainage divide, and the Buttermilk Creek and Danby Creek subbasins, Tompkins County, New York.



Base from U.S. Geological Survey, Seamless Data Distribution System, 1:24,000 accessed in 2008 at http://seamless.usgs.gov/ Universal Transverse Mercator projection, zone 18

**Figure 4.** Shaded relief and boundary of stratified-drift aquifer study area in the upper Buttermilk Creek and Danby Creek Valleys, Tompkins County, New York.

well data include previous U.S. Geological Survey (USGS) groundwater studies, the USGS National Water Information System (NWIS), and well records obtained from the New York State Department of Environmental Conservation (NYSDEC) Water Well Drillers Registration Program, and seven test wells that were drilled during this project.

At selected wells, the altitudes of water-level measuring points, which are typically the tops of the well casings, were determined to 0.01 feet (ft) using standard surveying techniques (Kenney, 2010). Elsewhere, altitudes of land surface at wells were estimated using light detection and ranging (lidar) technology and 1:24,000 scale topographic contour maps that were accurate to 0.5 ft and 5 ft, respectively. Depths to water below the measuring points were then converted to altitudes, which were used to construct maps of the water table and potentiometric surface of the unconfined and confined aquifers.

# Streamflow Measurements, Water Sampling, and Water Analysis

Streamflow was measured at 21 sites in the Buttermilk Creek and Danby Creek drainage basins from July 31 to August 4, 2008, using an USGS pygmy current meter. Samples were collected at 6 of the 21 streamflow sites. Streamflow measuring techniques are described by Rantz and others (1982).

Surface-water samples were collected at seven sites in the study area. The samples were analyzed for inorganic constituents and nutrients at the USGS National Water Quality Laboratory (NWQL) in Denver, Colorado. Field measurements of pH, specific conductance, dissolved oxygen, and water temperature were also recorded during sample collection. The samples were collected and processed by methods described in the USGS manual for the collection of water-quality data (U.S. Geological Survey, variously dated).

Groundwater-sample collection and processing followed standard USGS procedures as documented in the National Field Manual for the Collection of Water-Quality Data (U.S. Geological Survey, variously dated). Groundwater samples were collected from 12 wells in the study area; of these samples, 8 were collected during 2007-08 and were analyzed for inorganic constituents (common ions and trace elements) and nutrients at the NWQL. Field measurements of pH, specific conductance, dissolved oxygen, and water temperature were also measured during sample collection. Additionally, in response to concerns about natural-gas drilling, four groundwater samples were collected (background water-quality data) during 2010 and were analyzed for chemical constituents associated with gas drilling and (or) deep groundwater systems. Analytes included bromide, radiochemical constituents (gross alpha and gross beta), methane, and other dissolved gases, in addition to many of the constituents analyzed for in the 2007–08 samples.

Six groundwater samples were from test wells drilled for this study, five were from domestic wells, and one was from the well that supplies water to the Danby Town Hall. For the test wells, a stainless-steel submersible pump was used to purge the well of 5 to 10 casing volumes of water before collecting the water samples. The domestic wells and the town hall well were sampled after 20 minutes of pumping (until at least five casing volumes of well water had passed the sampling point); then, a raw-water spigot between the well and the pressure tank was opened, and the water was allowed to run for several more minutes to flush the spigot. Samples were collected from the raw-water spigot to avoid all watertreatment systems and to ensure that the water collected was representative of the water in the aquifer.

### Seismic-Refraction Surveys

Seismic-refraction surveys were conducted at three sites and HVSR ambient-noise seismic surveys were conducted at 5 sites in the study area. Seismic-refraction techniques used in this study are described by Haeni (1988). The seismicrefraction survey method measures the time it takes for a compressional sound wave to travel down through geologic units and back up to detectors (geophones) on land surface. For this investigation, a series of 12 geophones, spaced 50 or 100 ft apart, were inserted into the ground and the arrival times of compressional waves generated by explosives were recorded and plotted as a function of "source-togeophone" distances. A two-layer (saturated unconsolidated sediments and bedrock surface) or three-layer (unsaturated unconsolidated sediments, saturated unconsolidated sediments, and bedrock surface) boundary-formula computer analysis (Scott and others, 1972) was used to calculate depths to the bedrock surface.

The HVSR ambient-noise seismic method was also used to estimate unconsolidated sediment thickness and map the bedrock surface where the use of seismic refraction was not feasible. The HVSR ambient-noise seismic method uses a single, broad-band three-component seismometer to record ambient seismic noise. The ratio of the averaged horizontal-to-vertical frequency spectrum is used to determine the fundamental site resonance frequency, which can be interpreted using regression equations to estimate sediment thickness and depth to bedrock (Lane and others, 2008).

## Geology

Geologic materials in the Town of Danby consist of consolidated sediments (sedimentary bedrock) and unconsolidated sediments (unstratified and stratified glacial drift, alluvium, and swamp deposits). In most places, bedrock is covered by unconsolidated sediments. However, in some places, bedrock crops out at land surface, such as on hilltops, on upper parts of over-steepened valley walls, and in some stream channels that have incised through the unconsolidated sediments and flow on bedrock.

### **Bedrock**

Bedrock in the Town of Danby consists of sedimentary rocks (shale, siltstone, fine-grained sandstone, and limestone, fig. 5) of that were deposited in ancient seas during the Upper Devonian Period from 359.2 to 385.3 million years ago (U.S. Geological Survey Geologic Names Committee, 2010; Fisher and others, 1970; Isachsen and others, 1991; Zambito IV and others, 2007; and Karig and Elkins, 1986). The rocks were uplifted during the end of the Upper Devonian and the Mississippian Periods, which formed the Appalachian Plateau; the northern part of the plateau extends into southern and eastern New York State (fig. 2). About 300 million years ago (Pennsylvanian Period), the African plate began to collide with the North American plate (Isachsen and others, 1991) and the compressional stresses from the collision resulted in gentle folding of the rock units and faulting (both thrust and shear faults; fig. 5; Murphy, 1981; Podwysocki and others, 1982). The plateau was subsequently dissected by streams and eroded by glaciers, resulting in a region dominated by hills and valleys. At present, the rock units have a regional dip to the south at about 40 to 50 ft per mile (Williams and others, 1909); however, superimposed on the south-dipping rocks are gentle folds whose anticline and syncline axes trend roughly east to west (fig. 5). Because of the folds, there are local areas with reversal of dip direction (gentle dips to the north) and southward dipping slopes that have greater than the average regional dip of 40 foot per mile (ft/mi).

## **Glacial History**

The study area has undergone several glaciations during the Pleistocene Epoch, commonly referred to as the Ice Age, which began 2.6 million years ago and ended 11,850 years before present (uncalibrated radiocarbon years; 13,700 calibrated calendar years; Fullerton, 1980; U.S. Geological Survey Geologic Names Committee, 2010).

The last major glacial episode to affect the study area was during late Wisconsin time when the Laurentide Ice Sheet flowed from Canada into New York State. By about 23,000 years before present, the ice covered central New York and extended as far south as northern Pennsylvania (Muller and Calkin, 1993). The last ice to affect the study area was the Valley Heads Moraine re-advance into central New York (Muller and Calkin, 1993). At the terminus of the Valley Heads ice, an end moraine was deposited that loops across central and western New York (fig. 2). A segment of the end moraine crosses the study area in the Danby Creek valley (fig. 6). Valley Heads ice left the study area about 14,400 years ago (Cadwell and Muller, 2004).

## High-Altitude Valleys Little Affected by Ice Scour

Relatively low-altitude valleys that were generally aligned in the direction of glacial flow, such as the lower Sixmile Creek and Willseyville Creek valleys and the Cayuga Inlet valley (figs. 3 and 4), were preferential avenues for ice flow and were severely eroded. These severely eroded valleys were transformed into troughs. These troughs are wider and much deeper than other valleys, with U-shaped transverse valley profiles and smooth, planar bedrock walls. Each trough formerly contained a bedrock col (a high, narrow pass or saddle between watersheds) that was lowered and displaced by ice erosion (Miller and Karig, 2010; Kappel and Miller, 2003). Later, these troughs became filled with 200 to more than 300 ft of glacial and postglacial sediments.

In contrast, relatively high-altitude valleys, such as upper Buttermilk Creek and Danby Creek valleys, underwent much less scour and deposition of sediment than relatively loweraltitude valleys such as in the Cayuga Inlet trough and the lower Sixmile Creek and Willseyville Creek trough, which surround the study area to the west, north, and east (figs. 3 and 4). Scouring by ice was less extensive in high-altitude valleys because the ice was thinner and forces of erosion were less than in lower-altitude valleys. Relief between valley floor and adjacent hilltops in the upper Buttermilk Creek and Danby Creek valleys typically ranges from 400 to 500 ft whereas relief in the Cayuga Inlet trough and the lower Sixmile Creek and Willseyville Creek trough typically ranges from 800 to 1,000 ft. In the upper Buttermilk Creek and Danby Creek valleys, well records and seismic data indicate that the thickness of stratified-drift deposits generally ranges from 85 to 150 ft, which is one-half as much as that in the loweraltitude valleys.

## Wisconsin Deposits in the Study Area

The unconsolidated deposits in the upper Buttermilk and Danby Creek Basins contain glacial drift that consists of unstratified drift (till) and stratified drift (glaciolacustrine and glaciofluvial) sediments that were deposited during the Wisconsin Stage of the Pleistocene Epoch; postglacial alluvium and swamp deposits that were deposited during the Holocene Epoch. The areal distribution of surficial deposits in the study area is shown in figure 6 and the subsurface stratigraphy along four geohydrologic-section lines is shown in figures 7 through 10.

In the study area, till composes the largest portion of the drift in the uplands and in the valleys (fig. 6). Till is a poorly sorted, unstratified mixture of rocks embedded in a matrix consisting chiefly of clay, silt, and fine sand units (geologic units tl, ground-moraine deposits in the uplands and tv, ground-moraine deposits in the valley fill, fig. 6), and was deposited directly from glacial ice rather than transported by meltwater. In the uplands, the rocks in till range in size from



Figure 5. Bedrock and structural geology of the Town of Danby, Tompkins County, New York.



**Figure 6.** Surficial geology of upper Buttermilk Creek, Danby Creek, and Michigan Creek Basins, Town of Danby, Tompkins County, New York.

fine pebbles to boulders and are predominantly subangular to angular, flat platy clasts of local siltstone and shale. In the valleys, the rocks in till are subangular to rounded and range in size from fine pebbles to cobbles, consisting mostly of local siltstone and shale but also includes many glacial erratics (rocks such as sandstone, limestone, igneous, and metamorphic rocks). In most places in the uplands, till is the only deposit that overlies the bedrock (fig. 6). However, on some hilltops and over steepened hillsides in the Danby Creek valley, till may be absent and bedrock crops out at land surface (fig. 6). Till has low permeability (hydraulic conductivity is low).

Stratified drift (glaciolacustrine and glaciofluvial deposits) in the study area is present chiefly in valleys and in some places along the lower flanks of valley walls (figs. 6 through 10). Stratified drift consists of layered and sorted sediments. In the study area, fine-to-medium grained stratified drift consists of clay, silt, and fine-to-medium-grained sand of glaciolacustrine origin (lake-bottom deposits) that were laid down where heavily sediment-laden meltwaters and upland streams emerged into former small lakes such as proglacial lakes or lakes that formed in small basins and depressions in the upper Buttermilk and Danby Creek valleys. The mediumto-coarse grained stratified drift near the Hamlet of Danby that consists of pebbly sand with some silt was deposited at a delta along the edge of a former shallow glacial lake that once extended from the moraine in the northern part of upper Buttermilk valley to the Hamlet of Danby (figs. 6 and 8).

Coarse-grained stratified drift (sand and gravel) consists of layered and sorted glaciofluvial deposits (outwash sand and gravel, kames, and kame-end moraine) that were deposited by glacial meltwater streams that flowed on top, beneath, or within a glacier. Holocene fluvial deposits that consist of channel and flood-plain alluvium, and alluvial-fan deposits (fig. 6) were deposited by post-glacial streams. Kame endmoraine deposits tend to be poorly sorted and consist of silt, sand, and gravel and, in some places, of till. The kame-end moraine in the northern part of the study area forms a 30 to 40 ft high arcuate ridge where it crosses the upper Buttermilk valley and a 10 to 20 ft high ridge where it extends up both sides of the valley walls (fig. 6).

In the northern part of the study area (upper Buttermilk Creek valley), the glaciofluvial deposits in the valley are typically overlain by till (figs. 7 and 8), but in some places, the glaciofluvial deposits crop out at land surface, such as where a kame-end moraine crosses the valley and where several icecontact kame deposits crop out along the edges of the valley and on the lower flanks of hillsides (fig. 6). In the northern part of the study area, there may be one or two undifferentiated sand and gravel units (figs. 7 and 8). Near the northern boundary of the Town of Danby, well records indicate that there is only one sand and gravel unit—this unit is a basal unit 10 to 20 ft thick that overlies bedrock in the upper Buttermilk Creek buried gorge (fig. 7). Across the upper Buttermilk Creek valley and near the Hamlet of Danby, there are two undifferentiated sand and gravel units and one glaciolacustrine (delta) pebbly sand unit. The undifferentiated sand and gravel units include a basal sand and gravel unit on top of bedrock in the eastern part of the valley and an upper buried sand and gravel unit; each of these units range in thickness from 5 to 10 ft (fig. 8). The surficial glaciolacustrine unit is composed of deltaic-pebbly sand with some silt that ranges in thickness from 10 to 25 ft (fig. 8).

Holocene-age alluvial deposits in the northern part of the study area are typically thin, ranging from 1 to 10 ft thick in the stream channels and flood plains. However, the alluvial deposits are locally thicker (10 to 25 ft thick) at alluvial fans, such as the large fan deposit at Hornbrook Road east of the Hamlet of Danby (fig. 6). Large alluvial fans can form local small unconfined aquifers.

In the southern part of the study area (Danby Creek valley), the coarse-grained stratified drift (sand and gravel) is composed of glaciofluvial deposits (outwash and isolated kames along the lower flanks of hillsides) and Holocene alluvial deposits (fluvial channel, floodplain, and fan deposits; figs. 6, 9, and 10). However, it is unknown whether coarsegrained stratified drift underlies the large wetland that occupies a large part of the valley (fig. 6) because there are no subsurface data in this area. Where Steam Mill Road crosses Danby Creek valley in the southern part of the study area, the alluvial fan deposits ranges in thickness from 20 to 35 ft and forms a local unconfined aquifer (figs. 6 and 9). Where Durfee Hill Road crosses Danby Creek valley in the southern part of the study area, the outwash and alluvial channel and floodplain sand and gravel deposits range in thickness from 30 to 40 ft (figs. 6 and 10).

## Geohydrology of the Stratified-Drift Aquifer System

Characterization of the stratified-drift aquifer system in the Buttermilk and Danby Creek valleys in the Town of Danby, Tompkins County included describing the (1) aquifer type (confined or unconfined), (2) aquifer framework, (3) groundwater-flow system, including water levels and recharge and discharge conditions, and (4) water quality of surface water and groundwater. Data used to characterize the aquifer type and the geologic framework of the aquifers were obtained mostly from well records and geologist's logs of eight test wells drilled at five sites during this study (appendix 2). The locations of 140 wells in the Town of Danby are shown in figure 11. Records of these wells and construction details for the test wells drilled for this study are in appendix 1 and 2, and the range, mean, and median of well depths and reported well yields are summarized in table 1.



**Figure 7.** Geohydrologic section *A*–*A*′ across the northern part of upper Buttermilk Creek valley, Town of Danby, Tompkins County, New York.



**Figure 8.** Geohydrologic section *B–B*' across the headwaters of upper Buttermilk Creek and Danby Creek valleys near the hamlet of Danby, Town of Danby, Tompkins County, New York.



**Figure 9.** Geohydrologic section *C–C*′ across Danby Creek valley, along Steam Mill Road, Town of Danby, Tompkins County, New York.



#### **EXPLANATION**

#### Alluvial deposits of Holocene age

Channel and flood-plain alluvium—Gravel, sand, and silt deposited by recent streams in the valleys. Deposits are found beneath flood plains, low terraces, and in stream channels

#### Glaciofluvial deposits of Wisconsin age

Outwash sand and gravel—Chiefly sand and gravel beneath terrace remnants; as valley trains and beneath outwash plains; as fans and aprons; as delta topset beds. Forms the main part of the unconfined aquifer in the valley

#### Glaciolacustrine deposits of Wisconsin age

lss

al

osq

Lacustrine sand—Chiefly off-shore and near-shore deposits of former glacial and postglacial lakes; some

deposits are in small separate basins. Extent is unknown elsewhere

#### **Till of late Wisconsin age**

Ground-moraine deposits in the uplands (mostly lodgment till)tl

Consists of subangular to angular clasts (pebbles to boulders) embedded in a fine-grained matrix consisting of clay, silt and fine sand

#### **Devonian-age rocks**

Bedrock—Devonian-age shale, siltstone, and sandstone r

Geohydrologic symbols

035

TM1

Water table (unconfined aguifer)

÷ Water level—Static-water level in well (well may be finished in bedrock, or in the upper unconfined aquifer)

Well and site name-Site name is assigned by the U.S. Geological Survey. TM = Tompkins County, number following TM is sequential well number. CD = casing depth (in feet) and CD 81 TD = total depth (in feet) of an open-hole well finished in bedrock or of a test well (in some TD 88 cases, the casing of a test well may have been raised from TD to CD)

**Figure 10.** Geohydrologic section D-D' across Danby Creek valley, along Durfee Hill Road, Town of Danby, Tompkins County, New York.



Basemap created with TOPO!, scale 1:24:000 2003 National Geographic, URL www.nationalgeographic.com/topo

Figure 11. Location of wells in the Town of Danby, Tompkins County, New York.

## **Aquifer Types**

The unconsolidated sediments in the upper Buttermilk and Danby Creek valleys contain confined and unconfined aquifers and confining layers. The major stratified-drift aquifers in the study area are confined sand and gravel units in the upper Buttermilk Creek valley and an unconfined aquifer in the southern part in the Danby Creek valley (fig. 12). In addition to these major aquifers, alluvial fans (fig. 6) form thin and discontinuous unconfined sand and gravel aquifers in several localities (fig. 12). Aquifers are composed of sediments of sufficient permeability (moderate to high hydraulic conductivity) that can store and yield usable amounts of water to wells. Moderate hydraulic conductivity is defined from 10 to 100 feet per day (ft/d) and high hydraulic conductivity is greater than 100 ft/d. In an unconfined aquifer, the water table is at or near atmospheric pressure and is the upper boundary of the aquifer. There were only two wells (TM1654 and TM2590) in the study area that were finished in unconfined aquifers (fig. 12). The two wells were 27 and 41 ft deep and reported well yields were 10 and 20 gallons per minute (gal/min).

Confining layers are made up of poorly permeable sediments (low hydraulic conductivity) that prevent groundwater from rapidly moving through these layers. In the study area, the confining layers consist of mostly of till and some fine-grained glaciolacustrine sediments (figs. 7 through 10) clay, silt and fine sand.

A confined aquifer (also known as an artesian aquifer) is between confining layers and is composed of sediments and bedrock of sufficient permeability that it is able to store and yield usable amounts of water to wells. In a confined aquifer, the potentiometric surface is, by definition, always above the top of the aquifer. When a confined aquifer is tapped by a well, water in the casing is forced up to its potentiometric altitude. If the hydraulic pressure in the confined aquifer is

#### **EXPLANATION**



great enough, it may cause the well to flow above land surface, in which case, the well is referred to as a flowing artesian well. Wells that tap confined stratified-drift aquifers (glaciofluvial unit ud, figs. 7 through 9) in the Town of Danby ranged from 23 to 160 ft deep and mean and median well depths were 74 ft and 70 ft, respectively (table 1). Reported yields for wells finished in confined stratified-drift aquifers ranged from 3 to 50 gal/min and mean and median yields were 16 and 15, respectively (table 1).

In the Town of Danby, wells that tap the Devonian-age bedrock (mostly shale, siltstone, and sandstone) ranged from 50 to 420 ft deep and mean and median well depths were 202 ft and 183 ft, respectively (table 1). Reported yields for wells finished in bedrock aquifers ranged from 0.1 to 30 gal/min with mean and median yields of 6 and 4 gal/min, respectively (table 1).

### Upper Buttermilk Creek Valley

In the northern part of upper Buttermilk Creek valley, near the northern boundary of the Town of Danby, well records indicate that only one sand and gravel aquifer, a confined basal unit that is 10 to 20 ft thick and overlies bedrock at the valley bottom (fig. 7). Several domestic wells (wells ranged from 109 to 160 ft deep) in the central part of the valley tap this aquifer. Water-well drillers' logs indicate that well yields range from 8 to 50 gal/min, and averaged 20 gal/min for the six open-ended wells (6-in. diameter) that tap the confined aquifer in this area. These six wells are under confined conditions (but do not flow above land surface).

In the southern part of upper Buttermilk Creek valley, near the Hamlet of Danby, there are as many as four coarsegrained, stratified-drift geohydrologic units:

1. A thin surficial (unconfined) deltaic pebbly sand and silt unit that ranges from 5 to 10 ft thick (figs. 6 and 8) that was determined from test well TM2588 (well construction log in appendix 2, fig. 2–1).

2. A surficial (unconfined) alluvial fan deposit that is on the east side of the valley (figs. 6 and 8).

3. An upper sand and gravel unit (confined above and below by till) that ranges from 5 to 10 ft thick (fig. 8).

4. A basal confined sand gravel unit (confined above by till and below by bedrock) that was determined at test well TM2467 (well construction log in appendix 2, fig. 2–2) at the surface-water divide between Buttermilk Creek and Danby Creek Basins in the central part of the valley (figs. 7, 8, and 12).

The thin surficial (unconfined) deltaic pebbly sand unit had been tapped by shallow dug wells, but this unit is not considered to be a reliable aquifer because it is thin and typically becomes unsaturated during dry periods. Well data are not available at the alluvial fan in the eastern part of the valley (near Hornbrook Road); therefore, the geohydrologic condition is mostly unknown. The confined aquifers in the southern part of upper Buttermilk Creek valley do not flow above land surface in most areas except along a

Table 1.	Range, mean, and	l median of wel	l depths and	reported we	Il yields in the	Town of Danby, New York.
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[ft, feet; gal/min, gallons per minute. Data reported from water well drillers logs and stored in USGS National Water Information System computer data base accessible at http://waterdata.usgs.gov/ny/nwis/inventory.]

		Well dephs				Reported well yields			
Aquifer type	Number of wells	Range (ft)	Mean (ft)	Median (ft)	Number of wells	Range (gal/min)	Mean (gal/min)	Median (gal/min)	artesian flowing wells
Unconfined sand and gravel	2	27–41	34	34	2	10-20	15	15	0
Confined sand and gravel	32	23-160	74	70	30	3-50	16	15	5
Bedrock	99	50-420	202	183	95	0.1-30	6	4	0

low-altitude corridor adjacent to Buttermilk Creek north of the Hamlet of Danby (figs. 11 and 12). Most domestic wells in the upper Buttermilk Creek valley are completed in these confined aquifers.

## Danby Creek Valley

The aquifer types in Danby Creek valley are variable and depend on location (figs. 3 and 12). The lower (southeastern most) 8,000 ft of Danby Creek valley within the Town of Danby is occupied by a contiguous unconfined surficial aquifer (fig. 12) composed chiefly of sand and gravel (kames, outwash, and alluvium, figs. 6 and 10). The aquifer averages about 800 ft wide. Several domestic wells use this aquifer for water supply. Where Durfee Hill Road crosses Danby Creek valley (fig. 12), test-well records for TM1035 and TM2590 (fig, 10; well construction logs in appendix 2, figs. 2–3) indicate several important features:

1. The surficial sand and gravel deposits are 32 ft thick at that site.

2. The saturated thickness is 20 ft.

3. The unconfined aquifer may contain fine-grained lenses, such as silt.

4. A basal confined aquifer was not present.

5. Depth to bedrock was 68 ft (fig. 10).

This aquifer extends southward beyond Tompkins County and into Tioga County where it becomes part of the large aquifer system in the Catatonk Creek drainage basin (fig. 2 of Miller and Pittman, 2012).

There are also several small unconfined units (alluvial fans) along the west and east sides of the valley near Steam Mill Road (figs. 6, 8, 9, and 12). However, subsurface data are not available in these areas to determine the saturated thickness of these alluvial fan deposits.

An upper confined aquifer (confined above and below by till) was determined between depths of 19 and 23 ft at test well TM2467 (fig. 8). This aquifer apparently grades to fine sand between 43 and 46 ft at test well TM2591 near Steam Mill Road (fig. 9; well construction log in appendix 2, fig. 2–4) and between 44 and 48 ft at test well TM1035 at Durfee Hill

Road (fig. 10; well construction log in appendix 2, fig. 2–3). This sand unit is not considered an important aquifer in Danby Creek valley because it is thin (3 to 4 ft thick). Water-well drillers in central New York do not typically finish an openend cased well in fine sand because the sand is likely to flow inside the casing and, in some cases, mix with pumped water.

At a nested, test-well site (wells TM2589 and TM2591) at Steam Mill Road, well TM2589 penetrated a shallow 7-ft-thick confined aquifer at depths from 24 to 31 ft (fig. 9). Test well TM2591 (well log in appendix 2, fig. 2–4), an opened-ended cased well finished at depth 27 ft, yielded 20 gal/min during development of the well. Because well data are not available where a large wetland occupies most of the valley floor to the south of the test well TM2591, the presence of aquifers is unknown in that area (fig. 11).

The basal confined aquifer (confined above by till and below by shale) that was identified at test well TM2467 (at the surface-water divide between Buttermilk Creek and Danby Creek Basins, figs. 8 and 12), probably extends southward into Danby Creek valley. However, the basal confined aquifer was absent at test well TM2589 at Steam Mill Road (fig. 9, and well construction log in appendix 2, fig. 2–4), which indicates that the unit pinched out between the divide and Steam Mill Road (fig. 12).

### Michigan Creek Valley

Well data are not available in Michigan Creek valley (fig. 11); therefore, aquifer conditions are unknown. Based on geologic information, the ice-contact sand and gravel deposits (kames) that are on the lower flanks of the west valley wall of Michigan Creek valley (fig. 6) may extend below the wetland in the valley and form an aquifer; however, drilling a well would be difficult because of the wetland (fig. 12).

#### Bedrock

Bedrock aquifers (Devonian-age sedimentary rocks, fig. 6) that underlie the Town of Danby are the important sources of drinking water for homeowners in the uplands.



**Figure 12.** The types of stratified-drift aquifers in the upper Buttermilk Creek, Danby Creek, and Michigan Creek valleys in the Town of Danby, Tompkins County, New York.

However, the yields available from bedrock wells are generally much lower than from wells finished in sand and gravel. In the uplands, bedrock is both the only dependable source of water as well as the most economical source of potable water for homes, farms, and small commercial facilities.

Generally, water-well drillers do not drill wells more than 300 to 400 ft deep into bedrock in the study area because well yields typically decline and the water typically becomes increasingly more mineralized and less potable with increasing depth. Well yields generally decline with increasing depth because the density, width, and connectivity of water-bearing fractures decrease with increasing depth. The deepest domestic well completed in bedrock (TM2519) in the Town of Danby was 420 ft deep (reported yield was 0.1 gal/min before hydraulic fracturing [fracking] and 1.5 gal/min after fracking).

In the southern part of Tompkins County, records from water-well drillers indicate that 2 to 5 gal/min can usually be obtained from wells that tap bedrock at most locations. Locally, yields as large as 10 to 30 gal/min or as small as less than 1 gal/min are obtained (Miller, 2009).

## **Aquifer Geometry**

The stratified-drift aquifer system in the study area is contained in narrow, high-altitude valleys. These valleys contain relatively thin stratified- and unstratified-drift deposits. In most places in the study area, the thickness of these deposits are less than 100 ft, except in the northern part of upper Buttermilk Creek valley, where it is as much as 160 ft thick. About two-thirds of the drift is composed of fine-grained deposits (till or glaciolacustrine fine sand, silt, and clay) that form confining units. The remaining one-third of the drift is composed of coarse-grained deposits (sand and gravel) that form the unconfined and confined aquifers.

## **Unconfined Aquifers**

In the southern part of Danby Creek valley, an unconfined aquifer (green aquifer unit shown in fig. 12) that is 0.1 to 0.2-mi wide extends from the kame-end moraine at the southern end of the large wetland to the southern border of the town. The aquifer is composed of kame-end moraine, outwash, and Holocene alluvium (fig. 6). The unconfined aquifer was 32 ft thick at test-well sites TM1035 and TM2590 (well construction logs in appendix 2, fig. 2–3), Durfee Hill Road (fig. 10) and at least 41 ft thick at domestic well TM1654 (fig. 12). This aquifer is underlain by till (fig. 10; appendix 2, fig. 2–3).

Several of the large alluvial fans that extend throughout the study area (fig. 6) may form local, discontinuous unconfined aquifers where they have a sufficient saturated thickness (fig. 12). These formations may be as thick as 25 to 35 ft at the apex of the fan and become thinner toward the lower edges of the fans where they eventually pinch out. Although these aquifers are small, they typically receive high rates of recharge. In addition to receiving recharge from precipitation that directly falls on the fan, these fans are also recharged from tributary streams that typically lose water as they flow over the fans.

## **Confined Aquifers and Confining Units**

The confined aquifers in upper Buttermilk Creek valley are as much as 0.65 mi wide in the southern part and narrow to 0.3 mi wide in northern part of the valley (fig. 12). The upper confined aquifer (fig. 8) is south of the kame-end moraine (fig. 6) that crosses the valley in the northern part of the study area. This aquifer is typically 5 to 10 ft thick and is 20 to 40 ft below land surface (fig. 8). Many domestic wells in the valley in this area are completed in this aquifer. This aquifer is confined above and below by till that typically ranges in thickness from 15 to 30 ft (fig. 8).

A basal confined aguifer is present in the upper Buttermilk Creek interstadial buried gorge in the northern part of upper Buttermilk Creek valley (fig. 7) and in the eastern part of the southern part of the valley (fig. 8). Although the basal confined aquifer was absent at several well sites in the western part of the valley, it may be present in other areas. For example, at test-well site TM2806, well depth 47 ft, (fig. 11; well log in appendix 2, fig. 2-5) a second confined aquifer (probably the basal aquifer) was penetrated at 40 to at least 47 ft. Although the well was not drilled deep enough to penetrate bedrock, the results of a passive seismic survey at the well site indicated that bedrock was about 53 ft below land surface. The basal confined aquifer ranges in thickness from 3 to 24 ft and lies 58 to 74 ft below land surface in the central part of the valley, but locally it can be shallower (20 to 40 ft below land surface) such as along the edges of the valley (fig. 8). This aguifer is confined above by 15 to 30 ft of till in the southern part of the valley and by as much as 157 ft of mostly till in the northern part of the valley and is confined from below by bedrock (figs. 7 and 8).

In the northern part of Danby Creek valley, underlying an unconfined aquifer, there is an upper confined aquifer in the bottom of a shallow channel that had contained a shallow proglacial lake (figs. 9 and 12). In this shallow proglacial lake, a thin layer of glaciolacustrine fine sand and clay (confining unit) was deposited on top of an undifferentiated sand and gravel deposit (fig. 9). The glaciolacustrine sediments were subsequently overlain by Holocene alluvial-fan deposits (sand and gravel), which might form a local unconfined aquifer in the eastern part of the valley (fig. 9). The confined aquifer was 7-ft thick and was at depths from 24 to 31 ft at nested, test-well site TM2589 and TM2591 (appendix 2, fig. 2–4), Steam Mill Road (fig. 9). Elsewhere, the extent of this aquifer is unknown.

### **Groundwater Recharge**

Groundwater is recharged ultimately by infiltration of precipitation (rain and snowmelt) on the land surface. The amount of recharge to an aquifer is needed to determine a groundwater budget and the long-term availability of groundwater. The distribution and amount of recharge in the study area varies spatially, seasonally, and by degree of aquifer confinement. Most recharge is during the plant dormant season, typically March through April; and mid-October through mid-December. During this period, groundwater levels generally rise in the aquifers (fig. 13), indicating that aquifer storage is increasing. During the growing season (May through mid-October), evapotranspiration (the process by which water is changed from the liquid state into the vapor state and transpiration is the process by which water vapor escapes from living plants into the atmosphere) typically exceeds precipitation, which results in decreasing water levels and indicates that aquifer storage is decreasing. However, during large storms, the rate of recharge can exceed the rate of evapotranspiration, which results in a rise of groundwater levels even during the growing season (fig. 13).

The sources of recharge can be more easily estimated for unconfined aquifers than for confined aquifers. Large amounts of recharge are readily available to unconfined aquifers because these aquifers are exposed at land surface. Water from various sources, such as direct precipitation or loss from tributary streams contributes most of the recharge to unconfined aquifers. Recharge to confined aquifers is limited and generally poorly known because these aquifers are buried and typically only a small part of the aquifer is open to the atmosphere that can be recharged from direct precipitation. In the parts of the study area that contain confined aquifers, the extent, thickness, and type of sediments that overlie the aquifers affect the flux of water from the surface to the aquifers. The sources of recharge to and the discharge from the aquifers are depicted in figure 14.

#### **Unconfined Aquifer**

The principal sources of recharge to the unconfined aquifers in the study area include direct infiltration of precipitation (rain and snowmelt) at land surface, unchanneled surface runoff from adjacent hillsides that seeps into the aquifer along the edges of the valley, groundwater inflow from till and bedrock that enters the aquifer along the sides of the valley, seepage loss from upland-tributary streams where they flow over alluvial fans in the valley, and upward leakage from the underlying geologic units in the valley where the hydraulic head in these units are higher than the water table in the unconfined aquifer.

Direct infiltration of precipitation seeps to the water table and recharges the unconfined aquifer where sand and gravel is at land surface and when this water is not lost to evapotranspiration (the sum of evaporation and transpiration). The mean annual precipitation in the study area is about

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36.7 inches per year (in/yr) or 3.1 feet per year (ft/yr), based on precipitation data for 1971–2000 for Ithaca, N.Y. (Northeast Regional Climate Center, 2012), which is 5.0 mi northeast of the study area (fig. 4). In the southern part of Tompkins County, the mean annual evapotranspiration is estimated to be 18.5 in/yr (1.5 ft/yr; Kontis and others, 2004, plate 1). The amount of annual recharge from precipitation that falls directly over the unconfined aquifer in Danby Creek valley at the southern part of the study area can be estimated using equation 1:

$$R_p = (P - ET)A, \qquad (1)$$

where

$R_{n}$	is annual recharge from
P	precipitation that falls directly over
	the aquifer, in cubic feet per year,
Р	is mean annual precipitation, in feet
	per year,
ET	is mean annual evapotranspiration,
	in feet per year, and

*A* is area of the top of the unconfined aquifer, in square feet.

Using equation 1, the estimated annual recharge (*Rp*) from precipitation that falls directly over the unconfined aquifer in Danby Creek valley (the area represented by the contiguous surficial geologic deposits kame moraine, outwash sand and gravel, and alluvial fan deposits in fig. 6; and shown as a green unconfined aquifer unit in fig. 12) is computed to be  $1.25 \times 10^7$  cubic feet per year (ft<sup>3</sup>/yr) or 93.5 million gallons per year (Mgal/yr), where *P* is equal to 3.1 feet per year (ft/yr), ET = 1.5 feet per year (ft/yr), and  $A = 7.81 \times 10^6$  cubic feet ft<sup>3</sup> (0.28 square mile).

Another source of recharge to the unconfined aquifer is from adjacent unchanneled surface runoff and groundwater inflow that seeps into the aquifer from adjacent hillsides along the edges of the valleys (figs. 14 and 15). Runoff and groundwater inflow from these hillsides infiltrate into the unconfined aquifer along the edges of the valley (fig. 14*B* and 14*C*). The annual amount of recharge from these unchanneled hillsides can be estimated by using equation 2:

$$R_{u} = (P - ET)A, \qquad (2)$$

where

- $R_u$  is annual recharge from unchanneled hillsides, in cubic feet per year,
- *P* is mean annual precipitation, in feet per year,
- *ET* is mean annual evapotranspiration, in feet per year, and
- *A* is area of unchanneled hillsides that are adjacent to the unconfined aquifer, in square feet.



**Figure 13.** Altitude of groundwater levels in test wells TM2467, (near Town of Danby Highway Department) from June 1, 2008, to June 26, 2012, TM2588 (near the Hamlet of Danby) from December 17, 2008, to March 30, 2011, and TM2591, (Steam Mill Road) from February 11, 2009, to June 26, 2012, Town of Danby, New York.

#### A. Northern part of study area—Isolated basal confined aquifer



**B**. Central part of study area—Unconfined and confined aquifers



C. Southern part of study area—Unconfined aquifer



#### EXPLANATION

#### Hydrologic and recharge components

- ① Drainage divide: a boundary between adjacent drainage basins. Ground water and surface water divides in uplands are generally coincident
- (2) Edge of valley-fill deposits
- 3 Lateral contact between bedrock and either finegrained stratified drift and till. Flow across the contact is small
- Some water may recharge the buried aquifer as subsurface flow from underlying till or bedrock. Flow across the contact is small
- (5) Little or no water recharges the buried aquifer from streams flowing across the fine-grained deposits in the main valley. The main stem stream in the valley is typically a discharge zone (gains water)
- 6 Some water may recharge the confined aquifer as subsurface flow from adjacent fine-grained sediments, especially when the aquifer is pumped
- Precipitation recharges permeable surficial alluvial fans, kame moraines, outwash, and deltaic deposits, some of which constitute unconfined aquifers but some may be too thin to be of practical value
- 8 Precipitation recharges the confined aquifer where it crops out at land surface or underlies an alluvial fan along the edges of the valley
- Unchanneled and channeled runoff; and some subsurface flow from upland areas recharge surficial sand and gravel deposits along the edges of the valley and where tributary streams cross their alluvial fans. These flows may also be a source of recharge to buried aquifers (as well as precipitaion, see number 3) where confined sand and gravel units in central parts of the valley extend to the valley walls where confining units are absent and are directly overlain by surficial sand and gravel

#### Geohydrologic units

Water

Alluvial fan or channel sand and gravel Deltaic pebbly sand and gravel with some silt Outwash sand and gravel Confined sand and gravel aquifer Mostly till that is interlayered with some lacustrine fine sand, silt and clay Till Shale and siltstone Water table (unconfined aquifer)

**Figure 14.** Schematic geohydrologic sections showing sources of recharge to aquifers in the *A*, northern part of the study area (upper Buttermilk Creek valley); *B*, central part of the study area (upper Buttermilk Creek valley); and *C*, southern part of the study area (Danby Creek valley), Tompkins County, New York.

Using equation 2, the estimated annual recharge (Ru) to the unconfined aquifers in Danby Creek valley from precipitation that falls on adjacent unchanneled hillsides is estimated to be  $4.46 \times 10^7$  ft<sup>3</sup>/yr (334 Mgal/yr), where P = 3.1 ft/yr, ET = 1.5 ft/yr, and  $A = 2.79 \times 107$  ft<sup>2</sup> (1.00 mi<sup>2</sup>). Because the Danby Creek valley is narrow and the areas of the adjacent unchanneled hillsides are much larger (3.4 times larger) than the valley floor: the unconfined aquifer receives proportionally 3.4 times more recharge from the hillsides (334 Mgal/yr) than from precipitation that falls directly over the aquifer (93.5 Mgal/yr).

Additional sources of recharge to the unconfined aquifer in Danby Creek valley are tributaries that drain channelized upland basins (fig. 15). As upland tributary streams enter larger valleys floored with permeable sand and gravel, they typically loose water by infiltration through the streambed (Randall, 1978). These streams typically lose water where they flow over the alluvial fans in the large valleys. However, recharge from these tributary streams is difficult to estimate because streambed leakances are generally poorly known (Kontis and others, 2004). Paired streamflow measurements at the upstream (contact of the stratified drift and till-mantled bedrock valley wall) and downstream (mouth of stream) reaches can yield a rate of loss for that reach at that time and flow condition. However, paired streamflow measurements were not conducted for this study because most measuring sites were inaccessible. In addition to the unconfined aquifer in the southern part of Danby Creek valley, upland tributaries that flow over their isolated alluvial fans in the northern part of Danby Creek valley also provide recharge to these small unconfined aquifers. However, these small isolated alluvial fans may not be reliable aquifers because generally the deposits are thin; thus, the saturated thickness is thin and may be unsaturated during dry periods.

In the southern part of the study area, the recharge from six upland tributaries (losing streams) that flow over the contiguous unconfined aquifer in the Danby Creek valley (the green unconfined aquifer unit in fig. 15) was estimated using the following steps:

- Flow-duration at 10-percent intervals was computed for the six upland tributaries that flow over alluvial fans on the contiguous unconfined aquifer. These flow-duration intervals were computed by using the 1931–60 flowduration curve developed for Sage Brook near South New Berlin, N.Y. (streamgage 01501500) by Ku and others (1975). Sage Brook drains a small upland basin (0.61 mi<sup>2</sup>) that is composed mostly of till over bedrock, which is similar to the upland drainage basins in this study area.
- 2. The mean runoff (in cubic feet per square mile) was used from the Sage Brook flow-duration curve for each 10-percent flow-duration interval.
- 3. The upland drainage areas of the six upland tributaries that lose water to the contiguous unconfined aquifer in

Danby Creek valley were computed using a geographic information system (GIS).

- 4. The mean runoff values from step 2 were multiplied by the areas of the six upland drainage basins (fig. 15) from step 3 to obtain runoff (in ft<sup>3</sup>/s) in each stream in Danby Creek valley for each 10-percent flow-duration interval.
- 5. The length of each tributary channel that crosses the aquifer in Danby Creek valley was measured (in ft). These channel lengths were multiplied by 1.0 ft<sup>3</sup>/s per 1,000 ft to determine the upper limit of the rate of loss of flow that was most likely lost during high-flow periods. For tributary streams in the Susquehanna River Basin, Randall (1978) determined that a rate of loss of 1.0 ft<sup>3</sup>/s per 1,000 ft of channel was the typical maximum amount of loss for periods when streamflow is available.
- 6. For each tributary channel, the recharge for each 10-percent flow-duration interval were computed using the following criteria: For flow-duration intervals when the rate of flows that were determined in step 5 exceeded those in step 4, the flow in step 4 were used (which assumes 100 percent of streamflow infiltrates into the unconfined aquifer). For intervals where flows in step 4 exceeded those in step 5, the flows in step 5 were used (which assumes infiltration is a fraction of streamflow).
- 7. To determine the total recharge from each tributary, the results of step 6 were summed. The results of the above calculations indicated that the total drainage area of the six upland basins is 0.58 mi<sup>2</sup> and the sum of all recharge from losing streams that cross the unconfined aquifer was 0.75 ft<sup>3</sup>/s (table 2).

Some groundwater in the unconfined aquifer in Danby Creek valley is also recharged from bedrock. Some groundwater moves upward from the underlying bedrock at the bottom of the valley to the stratified-drift aquifer in the valley as indicated by the upward vertical gradient determined at the test wells TM1035 and TM2590 (fig. 11, well logs shown in appendix 2, fig. 2-3) in the southern part of the study area. For example, on November 5, 2008, the hydraulic head in test well TM1035 (finished in bedrock) was 6.4 ft higher than the water level in test well TM2590 (finished in the unconfined aquifer; appendix 2, fig. 2–3). However, the groundwater flow is small when compared to the horizontal flow in the unconfined aquifer because the bedrock is fairly impermeable (typically having a hydraulic conductivity less than 1 foot per day, ft/d) as well as is the till that overlies bedrock.

The estimated average-annual amounts of recharge to the unconfined aquifer in Danby Creek valley are summarized in table 2. The percentages of all sources of recharge to the contiguous unconfined aquifer in Danby Creek valley include 16 percent from precipitation that falls directly over the aquifer, 55 percent from runoff and groundwater inflow from adjacent unchanneled hillsides, and 29 percent





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Figure 15. Channeled and unchanneled upland drainage basins that provide recharge to the unconfined aquifers in the upper Buttermilk Creek, Danby Creek, and Michigan Creek valleys in the Town of Danby, Tompkins County, New York.

Table 2. Estimated average annual recharge to the contiguous unconfined aquifer in Danby Creek valley, Town of Danby, New York.

[ft<sup>3</sup>/s, cubic feet per second; ft<sup>3</sup>/d, cubic feet per day; ft<sup>3</sup>/yr, cubic feet per year; Mgal/d, million gallons per day; and Mgal/yr, million gallons per year]

Sources of recharge	(ft³/s)	(ft³/d)	(ft³/yr)	(Mgal/d)	(Mgal/yr)	Percent of total
Precipitation that falls directly on entire extent of aquifer	0.40	34,236	12,496,000	0.26	93	16
	Upl	and sources				
Runoff and groundwater inflow from adjacent unchannelized hillsides	1.41	122,300	44,640,000	0.91	334	55
Seepage from losing upland tributaries that flows over the aquifer (channeled flow)	0.75	64,800	23,652,000	0.48	177	29
Total average annual recharge to aquifer	2.56	221,336	80,788,000	1.65	604	100

from losing streams that cross the aquifer (table 2). The total annual recharge to the contiguous unconfined aquifer is 2.56 ft<sup>3</sup>/s (604 Mgal/yr)(table 2). The relatively small portion of recharge that comes from precipitation that falls directly on the aquifer in Danby Creek valley is a result of the relatively narrow valley (typically 600 to 1,000 ft wide), and thus, a relatively small part of the total drainage area. Hence, a proportionally larger amount of precipitation falls in the uplands in Danby Creek valley than in wider valleys. In general, where the unconfined aquifer in a valley is wider, more recharge is derived from precipitation as it falls directly on the aquifer when compared to that from other upland sources of recharge. For example, in two other central New York valleys in which recharge was estimated, Meads Creek valley is moderately wide (1,500 to 3,000 ft) and Otter Creek/Dry Creek valley is wide (6,000 to 8,000 ft; valleys not shown in this report); the portions of recharge that were derived from precipitation that falls directly on these aquifer valleys constituted 28 and 38 percent, respectively (Miller and others, 2008; and Miller and others, 1998), of the total recharge to these aquifers (compared to 16 percent in this study area).

## **Confined Aquifers**

The buried aquifers in the study area are confined by relatively impermeable fine-grained unconsolidated sediments (till or glaciolacustrine fine sand, silt, and clay) that impede or slow infiltration of water (recharge) from surface sources to the aquifer. However, parts of some confined aquifers appear to be semiconfined where they are bounded by relatively lowto-moderate permeable sediments (glaciolacustrine fine sand). In the part of the study area that contains confined aquifers (fig. 12), the sources and amounts of recharge are difficult to identify because the aquifers are concealed from view at land surface by overlying deposits (recent alluvium and confining units) and, in most areas, there is insufficient subsurface data (such as, well records, water-level measurements, and geophysical data). Therefore, the amount of recharge to the confined aquifers was not estimated in this study. In addition to more groundwater data, a numerical groundwater flow model would be the best tool to estimate the amount of water that recharges the confined aquifers. However, construction of a numerical groundwater flow model was beyond the scope of this study.

Although there were insufficient data to quantify the amount of recharge to the confined aquifers, some data are available to conceptualize how and where confined aquifers are recharged. Potential sources of recharge can be inferred by examining trends of groundwater levels. The trend of groundwater levels can offer clues to the degree of confinement of the aquifers.

In this study area, groundwater levels were monitored continuously for 2 to 3 years in test wells TM2467, TM2588, and TM2591, (fig. 13, location of wells shown in fig. 11). The degree of confinement is indicated by the response (or lack thereof) in water levels to phenomena, such as relatively rapid increases in water levels (spikes) during and just following individual rain-storm events, and from fluctuations caused by changes in stream and lake stages (Heath, 1976). Aquifers of low-to-moderate confinement have confining units that are discontinuous ("windows" of highly permeable sediments in the confining layer), thin, or consist of leaky semipermeable sediments such as fine-grained sand. Aquifers of low-to-moderate confinement receive moderate amounts of recharge from one or more of the following sources of water: (1) precipitation that falls directly on the windows of permeable sediments within the confining unit (the confining unit is discontinuous), (2) precipitation that falls directly on semipermeable parts of the confining unit and slowly infiltrates downward to the confined aquifer, and (3) groundwater inflow from till and bedrock from adjacent hillsides that seeps into the aquifer along the edges of the valley. Hydrographs of wells finished in aquifers of low-to-moderate degree of confinement are spiky and display relatively large fluctuations during the storm events, as determined for TM2591 (fig. 13).

Aquifers with a high degree of confinement (highly confined aquifers): (1) are distal from land-surface sources of recharge, (2) are confined by thick, extensive sediments with low hydraulic conductivities (low permeability), such as
till, silt, clay and shale, (3) receive only minor leakage from confining units, and (4) are insensitive to fluctuation in water levels in response to precipitation events. For highly confined aquifers, hydraulic conductivities are at least two orders of magnitude less than that of the aquifer. Groundwater level fluctuations in aquifers under high degree of confinement show gradual seasonal changes but show little or no changes due to individual storms (spikes). Wells TM1008 and TM1009 (well depths 187 and 130 ft, respectively) in lower Sixmile Creek valley, 4 mi east of Danby (locations shown in fig. 11), are examples of wells that are finished in confined aquifers (Miller and Karig, 2010, appendix 1) under a high degree of confinement (fig. 16).

The groundwater levels in test wells TM2467, TM2588, and TM2591 in this study area show that the degrees of confinement vary from low-to-moderate depending on location (fig. 13). Fluctuation of groundwater levels in response to storms is readily apparent (hydrograph with spikes and relatively large fluctuations during storm events) in well TM2591, somewhat less apparent in well TM2467, and least apparent in well TM2588 (fig. 13). The largest groundwater fluctuations were in well TM2591 (at the headwaters of Danby Creek valley, fig. 13), which suggests that the thin confining unit in that area (a 9-ft layer of fine sand and clay with some silt; fig. 9 and well-construction log in appendix 2, fig. 2-4) may be a discontinuous lens rather than an extensive layer. Also, the well site is on a large alluvial fan on the east side of Danby Creek valley (fig. 5, 6 and 12)-alluvial fans are areas where confined aquifers are commonly recharged (fig. 14B). In either case, the relatively large fluctuation in groundwater levels in response to storms indicates that the confined aquifer

is receiving relatively large amounts of recharge from a nearby source.

Groundwater fluctuations were smaller in well TM2467 (at the groundwater divide near Hornbrook Road) than in well TM2591 (fig. 13), which suggests that there is more confinement in the area of well TM2467 than in the area of well TM2591. The confining layer is relatively thick (35 ft) at well TM2467 (fig. 8 and appendix 2, fig. 2-2), and, thus, is likely to be more extensive than the confining layer near TM2591. However, the moderate spikes and fluctuations shown in this hydrograph for well TM2467 indicates that there is substantial hydraulic connection to a nearby source of recharge, even though the well is more distal to plausible sources of recharge than is well TM2591. The most likely source of nearby recharge to well TM2467 is the large alluvial fan 0.5 mi to the northeast on the east side of the valley near Hornbrook Road (figs. 6 and 8), which is where the basal confined aguifer in figure 8 is shown rising in altitude and meeting the unconconfined aquifer at the alluvial fan.

The abruptness in the spikes in the water level and the magnitude in the rise in the groundwater level were the least in well TM2588 (fig. 13, location shown in fig. 11) near the Hamlet of Danby in the western part of the headwaters of upper Buttermilk Creek valley. These two characteristics of the hydrograph indicate that the aquifer near well TM2588 is more confined and more distal from sources of recharge than aquifers near wells TM2591 and TM2467. The subtle spikes shown in the hydrograph for well TM2588 coincide with individual storms and also indicate that, although the aquifer is under a higher degree of confinement at this site than at the sites of wells TM2591 and TM2467 in this study



Figure 16. Altitude of groundwater levels in wells TM1008 and TM1009 finished in aquifers under high degree of confinement, lower Sixmile Creek valley, Town of Caroline, Tompkins County, New York. (Location of wells shown in fig. 11.)

area, the aquifer is not as confined as those aquifers that wells TM1008 and TM1009 are finished in the lower Sixmile Creek valley (fig. 16).

## **Groundwater Flowpaths and Discharge**

In the Danby Creek valley (southern part of the study area), groundwater in the unconfined aquifers flows to surface-water bodies (streams and wetlands) (fig. 17). The groundwater discharges (1) to domestic, commercial, and farm wells, (2) to Danby Creek, (3) to a large wetland in the central part of Danby Creek valley, and (4) as losses because of evapotranspiration (fig. 17). Where the unconfined aquifer is less than 15 ft from land surface, transpiration during the growing season reduces the amount of water that infiltrates to the water table and evaporation removes some water from the top of the water table. Mean annual evapotranspiration is estimated to range between 19 and 20 in/yr or about one-half of the annual precipitation in this area (Kontis and others, 2004; plate 1).

In the upper Buttermilk Creek valley, groundwater in the confined aquifers generally flows northwestward (downvalley) and discharges to (1) domestic and commercial wells, (2) Buttermilk Creek in the area near the northern town border, and (3) a small unnamed stream in a ravine at the south end of Buttermilk State Park just north of the town border (fig. 17). In the area of artesian-flowing conditions in the southwest part of the valley (shown by stippled area in fig. 17), the wells typically are not capped and continuously flow.

Although groundwater discharges continually throughout the year, the rate of discharge is typically less than recharge from late fall to mid spring. During this period, the rate of discharge is less than recharge because plants are dormant and the temperature is cool, which result in less loss as evapotranspiration and, therefore more precipitation is available as recharge. The greater rate of recharge is reflected by rising groundwater levels in the aquifer (fig. 13) indicating that aquifer storage is increasing. Conversely, during the growing season (from late spring to mid fall), the rate of discharge is generally greater than the rate of recharge and is reflected by groundwater levels that typically decrease in the aquifers during this period (fig. 13) indicating that aquifer storage is decreasing.

## **Groundwater Withdrawals**

An estimated 301 people depend on groundwater from the upper Buttermilk Creek and Danby Creek stratifieddrift aquifer system (table 3). The estimate was determined by a visual count of homes, farms, and businesses over the aquifer area on orthoimage maps (minus the number of households that were dependent on wells that tap bedrock) and multiplying that result by 2.3, the average number of persons per households in Tompkins County (table 3) from 2010 U.S. Census (U.S. Census Bureau, 2012). Because no people live over the Michigan Hollow Creek valley part of the aquifer system, no groundwater is withdrawn from wells. Another estimated 78 people reside in both valleys but depend on groundwater from wells that tap the bedrock (shale) aquifer (table 3).

The estimated total groundwater use from the stratifieddrift aquifer system is about 22,600 gallons per day (gal/d; 8.2 Mgal/yr) and about 5,860 gal/d (2.1 Mgal/yr) from the bedrock aquifers, based on an estimated average water use of 75 gal/d per person for self-supplied water systems in New York (Hutson and others, 2000) times the estimated 301 people that withdraw groundwater from the stratified-drift aquifers and the estimated 78 people that withdraw groundwater from the bedrock aquifers (table 3).

## Potential for Large-Capacity Municipal or Commercial Wells

The potential for large ground-water withdrawals is of interest to town and county planners as well as to potential commercial businesses that may wish to locate in the area. In addition, should hydraulic fracturing for unconventional development of wells drilled for natural gas be allowed in New York, the gas drilling companies may be seeking to withdraw large amounts of water from aquifers as well as surface water. The amount of recharge to stratified-drift aquifers is a key determinant of the long-term availability of water. In general, more recharge water is available to unconfined sand and gravel aquifers than to confined aquifers. Although several other large through valleys studied by the USGS contain stratified-drift aquifers capable of large yields (Miller, 2009; Miller and Karig, 2010; Miller and Bugliosi, 2013), the aquifers in this study area might not yield water in large quantities. Low yields of water are expected in this study area because the main confined aquifers in upper Buttermilk valley are thin and the main unconfined aquifer in Danby Creek valley is relatively small in areal extent and is thinly saturated (fig. 12).

At the time of this study (2006–12), all wells in the upper Buttermilk Creek and Danby Creek valleys were pumped intermittently at rates that typically range from 3 to 10 gallons per minute (gal/min) to adequately supply individual homes, small shops, or public buildings. The aquifers in these valleys could easily accommodate additional similar withdrawals. The question may arise in the future as to whether the aquifers could meet the water needs of a municipality or commercial facility which may need water in excess of 100 gal/min. To meet this large demand of water, it may be feasible to use an array of wells to tap the unconfined aquifer in lower Danby Creek valley.

## Unconfined Aquifer in Lower Danby Creek Valley

The lower (southeastern most) 8,000 ft of Danby Creek valley within the Town of Danby is occupied by a contiguous



**Figure 17.** Water table in unconfined aquifers and the potentiometric surface in the confined aquifer in the upper Buttermilk Creek, Danby Creek, and Michigan Creek valleys in the Town of Danby, Tompkins County, New York.

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 Table 3.
 Estimated groundwater withdrawals by users that reside over the upper Buttermilk Creek and Danby Creek valleys aquifer 

 system study area, 2012.
 System Study area, 2012.

[--, not applicable]

Users	Estimated number of wells that tap the stratified drift, sand and gravel aquifers <sup>a</sup>	Estimated number of wells that tap the bedrock aquifers <sup>b</sup>	Average people per household°	Estimated number of people using water from stratified drift, sand and gravel aquifers	Estmated number of people using water from bedrock aquifers	Average use per person, in gallons per day <sup>d</sup>	Estimated withdrawal, in gallons per day from stratified drift, sand and gravel aquifers	Estimated withdrawal, in gallons per day from bedrock aquifers
Private homes over the stratified-drift aquifers in upper Buttermilk Creek valley	117	26	2.3	269	60	75	20,183	4,485
Private homes over the stratified- drift aquifers in Danby Creek valley	14	8	2.3	32	18	75	2,415	1,380
Total	131	34		301	78		22,598	5,865

<sup>a</sup>The estimated number of wells that tap the stratified-drift, sand and gravel aquifers was determined by a visual count of homes, farms, and businesses over the aquifer area on orthoimage maps. Then the number of wells within the aquifer area that are known to tap bedrock were subtracted from the total number of wells.

<sup>b</sup>The estimated number of wells that tap the bedrock aquifers within the area that is overlain by the sand and gravel aquifers were estimated from well records. Since well records are not available for all wells within the area that is overlain by the sand and gravel aquifers, the number of bedrock wells may be somewhat higher than indicated.

°Source: U.S. Census Bureau, 2012, http://www.census.gov/prod/cen2010/, accessed August 15, 2015.

<sup>d</sup> From Hutson and others (2000).

unconfined surficial aguifer composed chiefly of sand and gravel (kames, outwash, and alluvium; fig. 6). The aquifer is relatively narrow-averaging in width of about 800 ft and has a small saturated thickness of about 20 ft (well TM1035, fig. 10). The total annual recharge to the contiguous unconfined aquifer is 2.56 ft<sup>3</sup>/s (1.65 Mgal/d; table 2). In order to provide sustainable continuous amount of water to large water users such as municipalities and industries, an array of wells spaced along this aquifer (for example, wells that are spaced 500 ft apart along the center of the aquifer, each screened in the bottom 5 ft of the aquifer) could probably be pumped collectively at a rate of several hundred gallons per minute during the winter and spring of most years. Water from Danby Creek would readily infiltrate and recharge the aquifer. However, the potential for induced recharge (when a pumping well creates a cone of depression that lowers an adjacent water table below the level of a stream causing the stream to lose water to the adjacent groundwater aquifer) from Danby Creek during summer periods and dry years of low flow is likely to be inadequate to sustain this rate of withdrawal. In addition, the amount of water stored in this thin, and limited in extent, aquifer would be small. If development of the unconfined

aquifer in the Danby Creek valley was contemplated, test wells would need to be drilled to collect more detailed data about the saturated thickness and aquifer transmissivity in order to determine the potential yield more precisely.

# Confined Aquifers in Upper Buttermilk Creek Valley

Records of several wells document the presence of a basal confined sand and gravel aquifer over bedrock and beneath thick till in the deeply incised bedrock valley beneath Buttermilk Creek in the northern part of the study area (figs. 7, 11, and 12; figs. 2–1, 2–2, and 2–5, in appendix 2). In addition, a basal confined aquifer over bedrock and a midsection confined sand and gravel aquifer are also documented near and somewhat north of the surface-water divide between Buttermilk Creek and Danby Creek (figs. 8 and 11; appendix 2, fig. 2–2) at higher altitude than in the northern part of the study area. These aquifers may be connected, although sufficient data is lacking to conclusively determine this connection. Well data from water-well drillers through 2012 indicate that these confined aquifers are thin (typically about

10 ft thick) and the reported yield data (ranged from 3 to 50 gal/min; table 1). These well yield data suggest that the confined aquifers meets the needs of homeowners and small commercial facilities, but these aquifers may not be capable of supplying sufficient water to meet the needs of municipalities and industries. However, additional geohydrologic data leading to calibration of a groundwater flow model would be needed to evaluate the long-term potential yield of the confined aquifer system in upper Buttermilk Creek valley.

# Water Quality

Water-quality samples were collected to characterize the chemical quality of surface water and groundwater in the study area. On July 31, 2008, seven surface-water samples were collected from seven sites, including three sites along upper Buttermilk Creek, two sites along Danby Creek, and two sites on tributaries to upper Buttermilk Creek (fig. 18). On various dates from 2007 through 2010, groundwater samples were collected from 12 wells (fig. 18). Field measurements were made for pH, specific conductance, and water temperature for surface-water samples and groundwater samples. The concentrations of 34 constituents, including inorganic major ions, nutrients, and trace elements were measured in surface water. In 9 of the 13 wells that were sampled from August 2007 through December 2008, the concentrations of 40 constituents, including inorganic major ions, nutrients, and trace elements were measured in groundwater samples. In 4 of the 13 wells that were sampled from June 2010 through October 2010, the samples were analyzed for 20 constituents associated with hydraulic fracturing, including inorganic major ions, trace elements, radiochemicals, and dissolved gases; as well as five nutrients.

The samples were shipped by overnight delivery to the (1) USGS National Water Quality Laboratory (NWQL) in Denver, Colo., for analysis of inorganic major ions, nutrients, inorganic trace elements, and (2) USGS Chlorofluorocarbon Laboratory (CFCL) in Reston, Virginia, for select dissolved gases. Analytical results for selected constituents were compared with U.S. Environmental Protection Agency (EPA) and New York State Department of Health (NYSDOH) drinking-water standards. The standards include Maximum Contaminant Levels (MCLs), Secondary Maximum Contaminant Levels (SMCLs), Secondary Drinking-Water Standard (SDWS), and Health Advisories (HAs) established by the EPA (2009) and the NYSDOH (2011).

## **Surface Water**

A total of 7 surface-water samples were collected by USGS personnel during base-flow conditions on July 31, 2008—2 sample sites were at upper Buttermilk Creek, 3 sites were at tributaries to Buttermilk Creek, and 2 sites were at Danby Creek (tables 4 and 5, fig. 18). Summary statistics for pH and several selected chemical constituents are in table 6 and summary statistics for trace elements are in table 7.

## Physiochemical Properties, Inorganic Major Ions, Nutrients, and Trace Elements

The pH of surface-water samples ranged from 6.8 to 8.0, with a median value of 7.6 (tables 4 and 6); all of the 7 samples were within the accepted SMCL range of 6.5 to 8.5 (U.S. Environmental Protection Agency, 2009). Specific conductance of the samples ranged from 308 to 1,360 microsiemens per centimeter at 25 degrees Celsius ( $\mu$ S/cm °C), with a median value of 341  $\mu$ S/cm °C (tables 4 and 6).

## Inorganic Major Ions

The cation detected in the greatest concentration was calcium, which ranged from 33.7 to 76.4 milligrams per liter (mg/L), with a median value of 44.0 mg/L (tables 4 and 6). Magnesium concentrations ranged from 6.65 to 15.4 mg/L, with a median of 8.63 mg/L. Calcium and magnesium contribute to water hardness, which ranged from 112 to 254 mg/L as calcium carbonate (CaCO<sub>3</sub>), with a median of 145 mg/L. Potassium concentrations ranged from 0.57 to 1.60 mg/L, with a median value of 1.15 mg/L (tables 4 and 6). Sodium concentrations ranged from 7.89 to 119 mg/L, with a median value of 16.0 mg/L.

The anion detected in the greatest concentration was bicarbonate, which ranged from 135 to 265 mg/L, with a median value of 174 mg/L (tables 4 and 6). Chloride concentrations ranged from 15.0 to 222 mg/L, with a median value of 25.0 mg/L. Silica concentrations ranged from 4.34 to 10.5 mg/L, with a median value of 5.28 mg/L (tables 4 and 6). Sulfate concentrations ranged from 1.62 to 11.9 mg/L, with a median value of 6.22 mg/L. None of the common major ions collected form surface-water sites in the study area exceeded any Federal or State water-quality standards (table 6) except for one sample, which exceeded the EPA drinking-water quality taste threshold of 60 mg/L for sodium.

## Nutrients

The nutrient detected in the greatest concentration was nitrate (as nitrite plus nitrate,  $NO_2+NO_3$ , tables 4 and 6), which ranged in concentrations from 0.022 to 0.454 mg/L as nitrogen, with a median value of 0.06 mg/L (tables 4 and 6). High concentrations of nitrogen can cause excessive plant and algal growth in streams, depleting oxygen and stressing organisms in their aquatic habitat; high concentrations also are a human health concern when the concentration is more than 10 mg/L in drinking water (U.S. Environmental Protection Agency, 2009). Nitrite concentrations were at or below the reporting limit of 0.002 mg/L as nitrogen. Orthophosphate concentrations ranged from an estimated low of 0.005 to



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**Figure 18.** The location of groundwater and surface-water sampling sites in upper Buttermilk Creek and Danby Creek valleys, Town of Danby, Tompkins County, New York.

in tables 8, 9, and 10

**Table 4.** Physiochemical properties of, and concentrations of inorganic major ions and nutrients in surface-water samples from

 Buttermilk Creek, selected tributaries to Buttermilk Creek, and Danby Creek, Town of Danby, New York, July 31, 2008.

[Trib, tributary; Parm code, USGS National Water Information System (NWIS) parameter code; mg/L, milligrams per liter;  $\mu$ S/cm, microsiemens per centimeter at 25 degrees Celcius; °C, degrees Celsius; CaCO<sub>3</sub>, calcium carbonate; <, less than; N, nitrogen; P, phosphorous; e, estimated value—constituent was detected in the sample but with low or inconsistent recovery; ft<sup>3</sup>/s, cubic foot per second. Sampling site locations are shown in figure 18.]

	Loca	al name	Buttermilk Creek at Danby	Buttermilk Creek South of Ithaca	Buttermilk Creek Trib 1 at SR-96B at Danby	Buttermilk Creek Trib 2 at Nelson Road near Danby	Buttermilk Creek Trib 2 at Danby	Danby Creek at Steam Mill Road near Danby	Danby Creek at SR-96B near Willseyville
	Date s	ampled	7/31/2008	7/31/2008	7/31/2008	7/31/2008	7/31/2008	7/31/2008	7/31/2008
USGS station ident	tification	number	04233210	04233220	04233212	04233214	04233216	01514666	01514672
Physiochemical properties	Parm code	Units			Values of	physiochemic	al properties		
Dissolved oxygen (field, standard units)	00300	mg/L	6.5	6.8	1.7	7.0	6.0	3.5	5.5
pH (field, standard units)	00400	рН	7.6	8.0	7.4	7.5	7.8	6.8	7.6
Specific conductance, at 25 °C (field)	00095	μS/ cm	341	457	1,360	326	413	325	308
Water temperature (field, standard units)	00010	°C	19.5	19.3	18.0	19.5	17.5	23.0	23.3
Stream discharge, instan- taneous	00061	ft³/s	0.05	0.88	0.001	0.08	0.34	0.03	0.72
Constituent					Concentrat	tions of inorga	nic major ions	S	
Cations:									
Calcium, filtered	00915	mg/L	38.0	57.6	76.4	44.0	57.3	39.5	33.7
Magnesium, filtered	00925	mg/L	7.93	10.6	15.4	8.63	9.76	7.4	6.65
Potassium, filtered	00935	mg/L	0.97	1.60	1.44	1.15	1.34	1.03	0.57
Sodium, filtered	00930	mg/L	16.2	17.5	119	7.89	10.5	12.1	16.0
Anions:									
Bicarbonate, filtered, as CaCO <sub>3</sub>	CALC <sup>a</sup>	mg/L	154	222	265	174	223	170	135
Chloride, filtered	00940	mg/L	25.4	28.5	222	15.0	19.1	17.7	25.0
Silica, filtered	00955	mg/L	4.36	4.34	10.5	5.95	5.28	7.38	4.71
Sulfate, filtered	00945	mg/L	6.22	11.9	7.95	3.88	11.0	1.62	4.81
Hardness, filtered	00900	mg/L	127	188	254	145	183	129	112
Alkalinity, filtered CaCO <sub>3</sub>	29801	mg/L	126	182	217	143	177	139	111
Dissolved solids, at 180 °C	70300	mg/L	210	271	654	210	247	224	195
Nutrients					Conc	entrations of ı	nutrients		
Ammonia, as N, filtered	00608	mg/L	< 0.02	< 0.02	0.030	< 0.02	< 0.02	0.214	0.021
Nitrate, as N, NO <sub>2</sub> +NO <sub>3</sub> , filtered as nitrogen	00631	mg/L	0.454	e 0.022	0.040	0.15	0.162	e 0.03	0.06
Nitrite, as N, filtered	00613	mg/L	e 0.002	< 0.002	e 0.002	e 0.002	< 0.002	0.002	e 0.002
Orthophosphate, as P, filtered	00671	mg/L	0.021	e 0.005	0.028	0.007	0.008	0.040	0.015
Phosphorus, unfiltered	00665	mg/L	0.031	e 0.004	0.105	e 0.006	e 0.006	0.117	0.067

<sup>a</sup> CALC = Bicarbonate values were calculated from alkalinity concentrations, which are given in milligrams per liter of CaCO<sub>4</sub>.

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 Table 5.
 Concentrations of trace elements in surface-water samples from Buttermilk Creek, selected tributaries to Buttermilk Creek, and Danby Creek, Town of Danby, New York, July 31, 2008.

[Trib, tributary; Parm code, USGS National Water Information System (NWIS) parameter code; µg/L, micrograms per liter; <, less than; e, estimated. Sampling site locations are shown in figure 18.]

	USGS sit	te name	Buttermilk Creek at Danby	Buttermilk Creek South of Ithaca	Buttermilk Creek Trib 1 at SR-96B at Danby	Buttermilk Creek Trib 2 at Nelson Rd near Danby	Buttermilk Creek Trib 2 at Danby	Danby Creek at Steam Mill Road near Danby	Danby Creek at SR-96B near Willseyville
	Date s	ampled	7/31/2008	7/31/2008	7/31/2008	7/31/2008	7/31/2008	7/31/2008	7/31/2008
USGS station ider	ntification	number	04233210	04233220	04233212	04233214	04233216	01514666	01514672
Constituent	Parm code	Units			Concent	trations of trac	e elements		
Aluminum, filtered	01106	μg/L	1.8	2.5	2.0	3.7	e 1.5	5.0	4.0
Antimony, filtered	01095	μg/L	< 0.140	e 0.07	e 0.080	< 0.140	< 0.140	< 0.140	< 0.140
Arsenic, filtered	01000	μg/L	0.75	0.5	1.3	0.42	0.27	2.7	0.94
Barium, filtered	01005	μg/L	28.4	65.8	137	63.0	82.5	87.8	42.0
Beryllium, filtered	01010	μg/L	< 0.008	< 0.008	< 0.008	< 0.008	< 0.008	e 0.006	< 0.008
Cadmium, filtered	01025	μg/L	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040
Chromium, filtered	01030	μg/L	< 0.12	< 0.12	< 0.12	< 0.12	< 0.12	e 0.10	< 0.12
Cobalt, filtered	01035	μg/L	0.056	0.050	0.151	0.053	0.049	0.984	0.092
Copper, filtered	01040	μg/L	e 0.60	1.35	e 0.65	e 0.54	e 0.59	< 1.0	< 1.0
Iron, filtered	01046	μg/L	21.8	<8.0	55.1	16.2	< 8.0	1,170	52.7
Lead, filtered	01049	μg/L	< 0.080	0.09	e 0.064	< 0.080	e 0.078	e 0.071	< 0.080
Manganese, filtered	01056	μg/L	23.5	1.8	157	25.0	4.44	3,550	180
Molybdenum, filtered	01060	μg/L	0.366	0.50	1.97	0.256	0.238	0.507	e 0.191
Nickel, filtered	01065	μg/L	0.52	0.45	0.88	0.38	0.33	1.2	0.41
Selenium, filtered	01145	μg/L	0.07	0.07	0.10	0.08	0.05	0.15	0.07
Silver, filtered	01075	μg/L	< 0.100	< 0.100	< 0.100	< 0.100	< 0.100	< 0.100	< 0.100
Uranium, natural, unfiltered	22703	μg/L	0.141	0.370	0.772	0.251	0.307	0.140	0.140
Zinc, filtered	01090	μg/L	e1.3	e 1.3	3.3	< 1.8	< 1.8	< 1.8	< 1.8

0.028 mg/L as phosphorous, with a median value of 0.015 (tables 4 and 6). None of the surface-water samples exceeded drinking-water standards for nitrate or nitrite.

## **Trace Elements**

The trace elements detected in the highest concentrations were barium, which ranged in concentrations from 24.8 to 137 micrograms per liter ( $\mu$ g/L), with a median of 65.8  $\mu$ g/L; iron, which ranged in concentrations from <8.0 to 1,170  $\mu$ g/L, with a median 21.8  $\mu$ g/L; and manganese, which ranged in concentrations from 1.8 to 3,550  $\mu$ g/L, with a median 25.0  $\mu$ g/L (table 7). The highest detected concentration of a trace element was 3,550  $\mu$ g/L of manganese in a sample from a stream (USGS streamgage 01514666, table 5) that drains a wetland area. Drinking-water standards were not exceeded for trace chemicals except for the sample collected at USGS streamgage 01514666 in which the concentration of iron was 1,170 µg/L, which exceeded the EPA SMCL of 300 µg/L, and the concentration of manganese was 3,550 µg/L, which exceeded the EPA SMCL of 300 µg/L (table 5).

#### Groundwater

Groundwater samples were collected from 13 wells that tap the stratified-drift and bedrock aquifers—seven samples were from the confined stratified-drift aquifers, one from an unconfined stratified-drift aquifer, and five were from bedrock aquifers (locations shown in fig. 18). Nine samples were analyzed for physiochemical properties and

		Number	Number	Surfa	ce-water s	amples	Number	Number	Grou	ndwater sai	nples
Constituent	Drinking- water standard	of surface- water samples	of samples exeed- ing limit	Minimum	Median	Maximum	of ground- water samples	of samples exeeding limit	Minimum	Median	Maximum
Physiochemical properties											
Dissolved oxygen, mg/L	None	7	×	1.7	6.0	7.0	12	x	0.1	0.4	2.8
pH (field)	$6.5 - 8.5^{a}$	7	0	6.8	7.6	8.0	13	0	7.4	7.8	8.4
Specific conductance, μS/cm at 25 °C	None	7	Х	308	341	1,360	13	Х	237	387	950
Constituent					Concentrat	ions of inorga	nic major io	us			
Cations											
Calcium, mg/L	None	7	Х	33.7	44.0	76.4	13	Х	8.52	38.6	102
Magnesium, mg/L	None	7	Х	6.65	8.63	15.4	13	Х	1.66	8.52	25.0
Potassium, mg/L	None	7	Х	0.57	1.15	1.60	13	Х	0.55	0.86	1.27
Sodium, mg/L	$60^{\mathrm{b}}$	7	1	7.89	16.0	119	13	С	4.83	16.3	159
Anions											
Bicarbonate, mg/L as CaCO <sub>3</sub>	None	7	Х	135	174	265	13	Х	134	240	362
Chloride, mg/L	$250^{a,b}$	7	0	15.0	25.0	222	13	0	0.85	6.12	179
Fluoride, mg/L	2.2°	7	0				13	0	0.10	0.28	0.43
Silica, mg/L	None	7	Х	4.34	5.28	10.5	13	Х	6.92	13.5	15.3
Sulfate, mg/L	250 <sup>a</sup>	7	0	1.62	6.22	11.9	13	0	0.44	13.3	35.4
Hardness, mg/L as CaCO <sub>3</sub>	None	7	Х	112	145	254	13	Х	28.6	143	353
Alkalinity, mg/L as CaCO <sub>3</sub>	250 <sup>a</sup>	7	0	111	143	217	13	2	110	197	267
Dissolved solids, dried at 180 °C	$500^{a}$	7	1	195	224	654	13	1	136	221	510
Nutrients					Conc	entrations of I	nutrients				
Ammonia, as N, filtered	None	7	Х	<0.02	<0.20	0.214	10	Х	<0.020	0.154	0.455
Nitrate, as N, NO <sub>2</sub> +NO <sub>3</sub> , filtered as nitrogen	10 <sup>a,c</sup>	7	0	0.022	0.06	0.454	10	0	<0.04	<0.04	0.28
Orthophosphate, as P, filtered	None	7	Х	e 0.005	0.015	0.040	10	Х	0.006	0.012	0.069

Table 6. Summary statistics for physiochemical properties, concentrations of inorganic major ions, and selected nutrients in surface-water and groundwater samples

°New York State Department of Health maximum contaminant level.

Summary statistics for concentrations of trace elements in surface-water and groundwater samples collected in the upper Buttermilk Creek and Danby Creek valleys, Town of Danby, New York, 2007–10. Table 7.

[All concentrations are in micrograms per liter (µg/L); —, not analyzed; X, not applicable because of no drinking-water standard; <, less than]

			Number of	Su	rface-water sa	mples		Number of	Gr	oundwater sar	nples
Constituent	Drinking- water standard	Number of samples	surface- water samples exceeding limit	Minimum	Median	Maximum	Number of samples	groundwater samples exceeding limit	Minimum	Median	Maximum
Aluminum	50 <sup>a</sup>	7	0	1.5	2.5	5.0	6	0	0.8	<4.0	256
Antimony	$6^{b,c}$	L	0	0.07	0.14	0.14	6	0	0.028	0.04	0.06
Arsenic	$10^{\rm b}$	L	0	0.27	0.75	2.7	13	1	0.18	2.2	12.2
Barium	2,000 <sup>b,c</sup>	L	0	28.4	65.8	137	13	0	36.1	189	1,290
Boron	None	0	Х				6	Х	11	45	346
Cadmium	$5^{\rm b,c}$	L	0	<0.040	<0.040	<0.040	6	0	<0.018	<0.02	0.026
Chromium	$100^{b,c}$	L	0	<0.12	<0.12	0.10	6	0	<0.12	<0.12	0.46
Cobalt	None	L	Х	0.049	0.056	0.984	6	Х	<0.02	0.038	0.24
Copper	$1,000^{a}$	7	0	0.54	0.65	1.35	6	0	<1.0	<1.0	180
Iron	$300^{\rm a,b}$	L	0	<8.0	21.8	1,170	13	4	<6.0	218	1,660
Lead	$15^{d}$	7	0	0.064	0.08	0.09	6	0	<0.06	<0.06	1.48
Lithium	None	0	Х				6	Х	2.40	6.21	155
Manganese	$50a-300^{\circ}$	L	$2^{-0}$	1.8	25.0	3,550	13	8–2	1.20	105	421
Molybdenum	None	7	Х	0.191	0.366	1.97	6	Х	0.12	1.97	7.39
Nickel	None	L	Х	0.33	0.45	1.2	6	Х	0.19	0.39	1.1
Selenium	$50^{\rm b,c}$	9	0	0.05	0.07	0.15	6	0	<0.06	<0.06	0.07
Strontium	None	٢	Х				6	Х	53.9	290	1,430
Uranium	$30^{\mathrm{b}}$	L	0	0.140	0.251	0.772	6	0	0.009	0.159	1.47
Zinc	$5,000^{a,b}$	L	0	1.3	1.8	3.3	6	0	<1.8	<2.0	24.1
<sup>a</sup> U.S. Environment	tal Protection A	vgency second:	ary maximum co	ontaminant leve	.le						
<sup>b</sup> U.S. Environment	tal Protection A	Agency maxim	um contaminant	t level.							

"New York State Department of Health (NYSDOH) maximum contaminant level.

<sup>d</sup>U.S. Environmental Protection Agency treatment technique.

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Table 8.         Physical and physical           in the Town of Danby, New York,	hemical pr , 2007–08.	operties ol	; and concer	ntrations of ir	norganic maj	or ions and r	utrients in gr	oundwater se	imples from s	tratified-drift	and bedrock	aquifers
[S&G, sand and gravel; conf, confined timeter; °C, degrees Celsius; N, nitrog Bold values exceed one or more drink	d; unconf, ur gen; P, phosp cing-water st	rconfined; P bhorous; <, l andards. Sai	arm code, USC ess than; e, esti npling site loca	3S National Wa imated value— ations are show	ter Informatior constituent was n in figure 18.]	I System (NWI s detected in the	S) parameter cc e sample but wi	de; mg/L, milli, th low or incons	grams per liter; istent recovery;	ft, feet; μS/cm, i .—, not analyze	microsiemens p d; NA, not app	er cen- icable.
	<b>USGS Stat</b>	ion Name	TM1037	TM2467	TM2588	TM2591	TM1035	TM2590	TM1036	TM1038	TM1018	Drink-
	Date	e sampled	12/10/2008	4/24/2008	11/12/2008	11/12/2008	11/13/2008	11/13/2008	11/13/2008	12/18/2008	8/16/2007	-ing-
	Station II	0 Number	422125 076290001	422119 076282401	422116 076284701	422052 076273101	421934 076251002	421934 076251001	422308 076300401	422132 076283601	422116 076285601	water stan- dard
Physical properties	Parm code	Units				escription an	id values of ph	ysical propert	ies			
Aquifer type			S&G, conf	S&G, conf	S&G, conf	S&G, conf	Bedrock	S&G, unconf	S&G, conf	S&G, conf	Bedrock	NA
Well depth, below land surface	72008	ft	30	75	41	27	88	27	150	23	200	NA
Physiochemical properties						Values of	physiochemic	al properties				
Dissolved oxygen (field, stan- dard units)	00300	mg/L	< 0.2	0.1	0.1		< 0.2	2.8	0.3	0.5	< 0.3	NA
pH (field, standard units)	00400	Hq	7.6	7.6	7.4	7.8	8.4	8.0	7.7	7.4	7.8	6.5–8.5 <sup>a</sup>
Specific conductance, (field)	00095	μS/cm	324	303	519	312	384	237	387	843	950	NA
Water Temperature (field, stan- dard units)	00010	oC	15.5	10.6	9.5	10.8	9.2	9.0	12.0	8.4	21.5	NA
Constituent						Concentrat	ions of inorga	nic major ions				
Cations:												
Calcium, filtered	00915	mg/L	42.9	38.6	65.2	35.4	8.52	35.4	50.3	102	27.5	NA
Magnesium, filtered	00925	mg/L	10.3	7.51	20.1	7.38	1.66	6.41	12.1	23.6	7.64	NA
Potassium, filtered	00935	mg/L	0.86	0.92	1.27	0.70	0.59	0.55	06.0	0.86	0.73	NA
Sodium, filtered	00930	mg/L	10.9	10.4	15.1	20.2	80.9	4.83	16.3	36.9	159	$60^{\mathrm{b}}$
Anions:												
Bicarbonate, filtered, as CaCO <sub>3</sub>	CALC <sup>1</sup>	mg/L	192	167	266	185	240	134	244	295	246	NA
Chloride, filtered	00940	mg/L	3.07	5.48	15.7	7.74	6.55	2.91	0.85	113	179	250 <sup>a,c</sup>
Fluoride, filtered	00950	mg/L	0.31	0.20	0.28	0.21	0.35	0.11	0.32	0.10	0.43	2.2°
Silica, filtered	00955	mg/L	12.0	13.8	15.3	10.7	7.19	6.92	13.5	14.2	6.95	NA
Sulfate, filtered	00945	mg/L	13.3	14.3	35.4	4.01	4.75	11.1	12.6	25.7	0.44	250 <sup>a,c</sup>
Hardness, filtered, as CaCO <sub>3</sub>	00600	mg/L	150	128	246	119	28.6	115	176	353	100	NA
Alkalinity, filtered CaCO <sub>3</sub>	29801	mg/L	157	137	222	152	197	110	200	242	202	NA
Dissolved solids, at 180 °C	70300	mg/L	183	184	309	174	223	136	221	468	510	$500^{\rm b}$

Water Quality

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Physical and physiochemical properties of; and concentrations of inorganic major ions and nutrients in groundwater samples from stratified-drift and bedrock aquifers in the Town of Danby, New York, 2007–08.—Continued Table 8.

[S&G, sand and gravel; conf, confined; unconfined; Parm code, USGS National Water Information System (NWIS) parameter code; mg/L, milligrams per liter; ft, feet; μS/cm, microsiemens per cen-timeter; °C, degrees Celsius; N, nitrogen; P, phosphorous; <, less than; e, estimated value—constituent was detected in the sample but with low or inconsistent recovery; —, not analyzed; NA, not applicable. Bold values exceed one or more drinking-water standards. Sampling site locations are shown in figure 18.]

	<b>USGS Stati</b>	on Name	TM1037	TM2467	TM2588	TM2591	TM1035	TM2590	TM1036	TM1038	TM1018	Drink-
	Date	sampled	12/10/2008	4/24/2008	11/12/2008	11/12/2008	11/13/2008	11/13/2008	11/13/2008	12/18/2008	8/16/2007	ing-
	Station ID	Number	422125 076290001	422119 076282401	422116 076284701	422052 076273101	421934 076251002	421934 076251001	422308 076300401	422132 076283601	422116 076285601	water stan- dard
Nutrients	Parm code	Units				Сопсе	intrations of r	utrients				
Ammonia, as N, filtered	00608	mg/L	0.171	0.068	0.235	090.0	0.138	< 0.020	0.455	0.044	0.425	NA
Nitrate, as N, NO <sub>2</sub> +NO <sub>3</sub> , filtered as nitrogen	00631	mg/L	< 0.04	<0.04	<0.04	< 0.04	< 0.04	0.28	< 0.04	< 0.04	< 0.060	$10^{\rm d,c}$
Nitrite, as N, filtered	00613	mg/L	< 0.002	<0.002	<0.002	< 0.002	< 0.002	< 0.002	< 0.002	< 0.002	< 0.002	1 d,c
Ammonia plus organic-N, as N, filtered	00623	mg/L	0.20	e 0.10	0.26	e 0.08	0.02	< 0.10	0.45	0.11	0.48	NA
Orthophosphate, as P, filtered	00671	mg/L	e 0.007	0.014	e 0.006	0.021	0.069	e 0.006	0.017	e 0.006	0.012	NA
<sup>a</sup> U.S. Environmental Protection A <sub>i</sub>	gency second	lary drinkin	g water standa	rd.								
<sup>b</sup> U.S. Environmental Protection Ag	gency second	ary drinkin <sub>i</sub>	g water advisor	y taste thresho	ld.							
° New York State Department of H	ealth maxim	um contami	nant level.									

<sup>d</sup> U.S. Environmental Protection Agency maximum contaminant level.

 $^{1}$ CALC = Bicarbonate values were calculated from alkalinity concentrations, which are given in milligrams per liter of CaCO<sub>3</sub>.

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[S&G, sand and gravel; conf, confined; Parm code, USGS National Water Information System (NWIS) parameter code; µg/L, micrograms per liter; <, less than; —, not analyzed. Bold values exceed one or more drinking-water standards. Sampling site locations are shown in figure 18.]

	<b>USGS Sta</b>	ntion Name	TM1037	TM2467	TM2588	TM2591	TM1035	TM2590	TM1036	TM1038	TM1018		TM1018	_		
	Dai	te sampled	12/10/2008	4/24/2008	11/12/2008	11/12/2008	11/13/2008	11/13/2008	11/13/2008	12/18/2008	8/16/2007		8/16/200	7		
	Station	ID Number	422125 076290001	422119 076282401	422116 076284701	422052 076273101	421934 076251002	421934 076251001	422308 076300401	422132 076283601	422116 076285601	42	2116 0762	85601		
	A	quifer type	S&G, conf	S&G, conf	S&G, conf	S&G, conf	Bedrock	S&G, unconf	S&G, conf	S&G, conf	Bedrock		Bedroc	~		
Constitue	nt cod fil-	m e Units d)			Cor	ncentrations (	of trace elem	ents					Parm code (unfil- tered)	Units	Trace ele- ments	Drinking water standard
Aluminum, filtered	0110	16 μg/L	< 4.0	e 0.8	< 4.0	< 4.0	< 4.0	e 2.8	< 4.0	< 4.0		Aluminum, unfiltered	01105	µg/L	256	50 <sup>a</sup>
Antimony, filtered	0105	5 μg/L	e 0.028	< 0.140	0.051	e 0.030	0.060	0.040	e 0.03	< 0.040		Antimony, unfiltered	01097	μg/L	< 0.20	6 <sup>b,c</sup>
Arsenic, filtered	0100	00 µg/L	8.2	3.7	12.2	5.5	2.2	0.23	1.2	0.40		Arsenic, un- filtered	01002	μg/L	0.23	10°
Barium, filtered	0100	15 μg/L	189	199	116	581	424	36.1	65.6	384		Barium, un- filtered	01007	µg/L	1,290	2,000 <sup>b,c</sup>
Beryllium, filtered	0101	0 µg/L	< 0.020	< 0.008	< 0.020	< 0.020	< 0.020	< 0.020	< 0.020	< 0.020		Beryllium, unfiltered	01012	μg/L	< 0.06	4 <sup>b,c</sup>
Boron, filte	red 0102	0 μg/L	31	32	48	57	214	11	45	14	346	Boron, un- filtered	01022	µg/L		None
Cadmium, filtered	0102	5 μg/L	e 0.019	< 0.040	e 0.019	e 0.012	< 0.020	< 0.020	0.026	< 0.02		Cadmium, unfiltered	01027	µg/L	< 0.018	5 <sup>b,c</sup>
Chromium, filtered	0103	0 µg/L	< 0.12	< 0.12	< 0.12	< 0.12	< 0.12	0.20	< 0.12	e 0.06		Chromium, unfiltered	01034	μg/L	e 0.46	100 <sup>b,c</sup>
Cobalt, filte	red 0103	5 µg/L	0.038	0.03	0.077	0.025	< 0.020	0.036	0.047	0.126		Cobalt, un- filtered	01037	μg/L	0.24	None
Copper, filtered	0104	μ0 μg/L	< 1.0	< 1.0	< 1.0	< 1.0	< 1.0	< 1.0	< 1.0	< 1.0		Copper, un- filtered	01042	μg/L	180	$1,000^{a}$
Iron, filtere	d 0104	46 μg/L	247	163	196	324	89.2	13.2	255	381	238	Iron, unfil- tered	01045	μg/L	1,660	300 <sup>a,c</sup>
Lead, filtere	sd 0104	-9 μg/L	0.074	< 0.080	< 0.060	e 0.051	< 0.060	< 0.060	0.208	< 0.060		Lead, unfil- tered	01051	μg/L	1.48	15 <sup>d</sup>
Lithium, filtered	0113	0 µg/L	5.89	6.21	11.6	4.79	25.6	2.40	2.73	8.46		Lithium, unfiltered	01132	μg/L	155	None
Manganese	0105	6 μg/L	177	169	101	116	20.1	2.16	149	317	35.9	Manganese, unfiltered	01055	μg/L	43.4	50ª-300°

Concentrations of trace elements in groundwater samples from stratified-drift and bedrock aquifers in the upper Buttermilk Creek and Danby Creek valleys, Town of Danby, New York, 2007–08.—Continued Table 9.

[S&G, sand and gravel; conf. confined; Parm code, USGS National Water Information System (NWIS) parameter code; µg/L, micrograms per liter; </ less than; —, not analyzed. Bold values exceed one or

Ď	SGS Station	Name	TM1037	TM2467	TM2588	TM2591	TM1035	TM2590	TM1036	TM1038	TM1018		TM1018	~		
	Date sa	mpled	12/10/2008	4/24/2008	11/12/2008	11/12/2008	11/13/2008	11/13/2008	11/13/2008	12/18/2008	8/16/2007		8/16/200	2		
-	Station ID N	lumber	422125 076290001	422119 076282401	422116 076284701	422052 076273101	421934 076251002	421934 076251001	422308 076300401	422132 076283601	422116 076285601	422	2116 07628	85601		
	Aquif	er type	S&G, conf	S&G, conf	S&G, conf	S&G, conf	Bedrock	S&G, unconf	S&G, conf	S&G, conf	Bedrock		Bedroc	¥		
Molybdenum, filtered	01060	µg/L	7.39	2.53	5.86	1.97	1.72	0.178	6.44	0.825		Molybde- num, unfiltered	01062	µg/L	0.12	None
Nickel, filtered	01065	µg/L	0.37	e 0.19	0.60	0.39	0.31	0.39	0.45	1.1		Nickel, un- filtered	01067	µg/L	0.48	None
Selenium, filtered	01145	µg/L	< 0.06	< 0.06	< 0.06	< 0.06	< 0.06	0.07	e 0.03	< 0.06		Selenium, unfiltered	01147	µg/L	<0.080	50 <sup>b,c</sup>
Silver, filtered	01075	µg/L	<0.0008	< 0.100	<0.0008	<0.0008	<0.0008	< 0.008	< 0.008	< 0.008		Silver, unfil- tered	01077	µg/L	< 0.016	100 <sup>b,c</sup>
Strontium, filtered	01080	µg/L	324	290	424	184	176	53.9	480	189		Strontium, unfiltered	01082	µg/L	1,430	None
Thallium, filtered	01057	µg/L	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040	< 0.040		Thallium, unfiltered	01059	µg/L	<0.18	None
Uranium, natural, filtered	22703	µg/L	0.760	0.183	1.47	0.089	0.009	0.159	0.145	0.944		Uranium, natural, unfiltered	28011	µg/L	0.018	30 <sup>b</sup>
Zinc, filtered	01090	µg/L	4.3	< 1.8	< 2.0	< 2.0	< 2.0	< 2.0	e 1.7	24.1		Zinc, unfil- tered	01092	µg/L	4.6	5,000 <sup>a,c</sup>
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<sup>b</sup> U.S. Environmental Protection Agency maximum contaminant level.
<sup>c</sup> New York State Department of Health maximum contaminant level.
<sup>d</sup> U.S. Environmental Protection Agency treatment technique.

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chemical constituents (inorganic major ions, nutrients, and trace metals) that typically characterize the general water quality (tables 8 and 9). In addition, to obtain background data because of concerns about natural-gas drilling, four samples were analyzed for physiochemical properties and selected chemical constituents including inorganic major ions, nutrients, trace elements, bromide, radiochemical constituents [gross alpha and gross beta], methane, and other dissolved gases (table 10). Three groundwater samples were from the Town of Danby owned facilities (TM1037, Danby Fire Hall; TM2467, Highway Department; and TM1018, Town Hall), five were from private homes (TM1036, TM1038, TM2519, TM2760, and TM1491), and five were from test wells drilled for this study (TM2588, TM2591, TM1035, TM2590, and TM2806). Dissolved oxygen, pH, specific conductance, and water temperature were measured in the field. Samples were analyzed for inorganic major ions, nutrients, and trace elements by the USGS NWOL in Denver, Colo. Results of chemical analyses are tabulated in tables 8, 9, and 10, and summary statistics of selected chemical constituents are in tables 6 and 7.

## **Physical and Physiochemical Properties**

Wells that were sampled ranged from 23 to 420 ft deep and the pH ranged from 7.4 to 8.4, with a median value of 7.8 (tables 6, 8, and 10); no pH measurement was outside the accepted EPA SMCL range of 6.5 to 8.5. Specific conductance of the samples ranged from 237 to 950  $\mu$ S/cm, with a median value of 387  $\mu$ S/cm (tables 6, 8, and 10).

## Inorganic Major Ions

The cation that was detected with the largest concentration was calcium, which ranged from 8.52 to 102 mg/L, with a median value of 38.6 mg/L (tables 6, 8, and 10). Magnesium concentrations ranged from 1.66 to 25.0 mg/L, with a median value of 8.52 mg/L. Sodium concentrations ranged from 4.83 to 159 mg/L, with a median value of 16.3 mg/L. The concentrations of sodium in three samples from the bedrock aquifer exceeded the EPA Drinking Water Advisory Taste Threshold of 60 mg/L (tables 8 and 10) (U.S. Environmental Protection Agency, 2009). No samples from the sand and gravel aquifer exceeded the EPA Drinking Water Advisory Taste Threshold (U.S. Environmental Protection Agency, 2009). Potassium concentrations were detected in low concentrations in the study area (generally less than 1.0 mg/L). Potassium concentrations ranged from 0.55 to 1.27 mg/L, with a median value of 0.86 mg/L (tables 6, 8, and 10).

Calcium and magnesium contribute to water hardness as  $CaCO_3$ , which ranged from 28.6 to 353 mg/L, with a median value of 143 mg/L (tables 6, 8, and 10). Of the 13 samples, one was soft (0 to 60 mg/L as  $CaCO_3$ ), four were moderately hard (61 to 120 mg/L as  $CaCO_3$ ), and four were hard (121 to

180 mg/L as  $CaCO_3$ ), and four were very hard water (Hem, 1985). Hard and very hard water consumes excessive amounts of soap and detergents and forms an insoluble scum or scale (Hem, 1985).

Alkalinity, which results from dissolution of carbonate minerals such as those composing limestone and dolomite and is a measure of the capacity of water to neutralize acid, ranged from 110 to 267 mg/L as CaCO<sub>3</sub>, with a median value of 197 mg/L of CaCO<sub>3</sub> (tables 6, 8, and 10). Alkalinity concentrations lower than 100 mg/L can be corrosive when pH is low (below pH 6); when alkalinity is greater than 150 mg/L, it can cause scale (lime) buildup in plumbing pipes (Mechenich and Andrews, 2004). Dissolved solids concentrations ranged from 136 to 510 mg/L (tables 6, 8 and 10), with a median value of 221 mg/L; the dissolved solids concentration in one sample from a well that taps the bedrock aquifer (TM1018) exceeded the EPA SDWS for dissolved solids solids of 500 mg/L.

The anion detected with the largest concentration was bicarbonate, which ranged from 134 to 362 mg/L, with a median value of 240 mg/L (tables 6, 8, and 10). Bicarbonate values were calculated from alkalinity concentrations, which are given in milligrams per liter of CaCO<sub>3</sub>. Chloride concentrations ranged from and 0.85 to 179 mg/L, with median value of 6.12 (tables 6, 8, and 10). The concentrations of chloride did not exceed the NYSDOH MCL and EPA SDWS of 250 mg/L. Silica concentrations ranged from and 6.92 to 15.3 mg/L, with a median value of 13.5 (tables 6, 8, and 10). Sulfate concentrations ranged from 0.44 to 35.4 mg/L, with a median value of 13.3 mg/L (tables 6, 8, and 10). The concentration of sulfate did not exceed the NYSDOH MCL and EPA SDWS of 250 mg/L.

## Nutrients

The dominant nutrient detected in groundwater samples in the upper Buttermilk Creek and Danby Creek valleys was ammonia (table 8). The concentration of ammonia ranged from less than (<)0.020 to 0.455 mg/L as nitrogen (N), with a median value of 0.154 mg/L as N (tables 6, 8, and 10). In this report, the medians of data that included censored data (values of concentrations less than the reporting limit) were calculated by sorting those concentrations that had less than values at the smallest end of the range of values. Ammonia was detected in all samples from wells that were finished in confined sand and gravel aquifers and the bedrock aquifer, but was not detected in the sample from well TM2590; this well was finished in the unconfined aquifer. The concentrations of nitrate plus nitrite (NO<sub>2</sub>+NO<sub>3)</sub> and of nitrite did not exceed the NYSDOH and EPA MCLs of 10 mg/L as N and 1 mg/L as N, respectively, in any sample. Nitrate was in a concentration above the detection limit of 0.04 mg/L in only one sample-0.28 mg/L, which was from well TM2590 that was finished in the unconfined aquifer. Nitrite was not detected in any sample above the reporting limit of 0.002 mg/L. The concentrations of orthophosphate ranged from an estimated minimum of

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**Table 10.** Physical and physiochemical properties of; and concentrations of inorganic major ions, trace elements, nutrients, selected radiochemicals, and dissolved gases in groundwater samples from bedrock and stratified-drift aquifers in the Town of Danby, New York, 2010.

 $[S\&G, sand and gravel; conf, confined aquifer; Parm code, USGS National Water Information System (NWIS) parameter code; mg/L, milligrams per liter; <math>\mu g/L$ , micrograms per liter; °C, degrees Celsius; e, estimated; ft, feet; <, less than; —, not analyzed; picocurie per liter (pCi/L); R, radchem non-detect. Bold values exceed one or more drinking-water standards. Sampling site locations are shown in figure 18.]

	USGS Stat	ion Name	TM2519	TM2760	TM1491	TM2806	
	Date	sampled	6/2/2010	6/2/2010	6/2/2010	10/28/2010	Drinking
	Station II	D Number	422244 076310701	422254 076284801	422029 076255001	422143 076291101	standard
Physical properties	Parm code	Units	Descrip	otion and value	es of physical p	oroperties	
Aquifer type			Shale	Shale	Shale	S&G, conf.	None
Well depth, below land surface	72008	ft	420	177	300	47	None
Physiochemical properties			Va	lues of physio	chemical prop	erties	
Dissolved oxygen (field, standard units)	00300	mg/L	1.8	0.7	0.7	1.0	None
pH (field, standard units)	00400	рН	8.2	7.6	7.8	7.9	6.5-8.5 <sup>b</sup>
Specific conductance, at 25 °C (field)	00095	μS/cm	661	598	328	398	None
Water temperature (field, standard units)	00010	°C	13.5	12.0	11.6	10.8	None
Constituent			Con	centrations of	inorganic maj	or ions	
Cations:							
Calcium, filtered	00915	mg/L	34.4	71.5	43.3	52.4	None
Magnesium, filtered	00925	mg/L	7.7	25.0	8.52	14	None
Potassium, filtered	00935	mg/L	0.72	0.81	0.57	0.94	None
Sodium, filtered	00930	mg/L	104	26.0	14.5	10.0	60°
Anions:							
Bicarbonate, filtered, as CaCO <sub>3</sub>	CALC <sup>1</sup>	mg/L	315	362	183	231	None
Bromide, filtered	71870	mg/L	0.274	e 0.020	e 0.019	0.021	None
Chloride, filtered	00940	mg/L	51.3	1.59	4.10	6.12	250 <sup>b,d</sup>
Fluoride, filtered	00950	mg/L	0.72	0.21	0.19	0.29	2.2 <sup>d</sup>
Silica, filtered	00955	mg/L	10.9	13.6	14.4	15.3	None
Sulfate, filtered	00945	mg/L	5.82	32.3	17.2	19.2	250 <sup>b,d</sup>
Hardness, filtered, as CaCO <sub>3</sub>	00900	mg/L	118	282	143	189	None
Alkalinity, filtered, as CaCO <sub>3</sub>	29801	mg/L	258	267	150	189	None
Dissolved solids, at 180 °C	70300	mg/L	367	344	187	217	500°
Trace metals			C	oncentrations	of trace eleme	ents	
Arsenic, filtered	01000	μg/L	0.25	5.5	0.18	6.8	10 <sup>a,d</sup>
Barium, filtered	01005	μg/L	304	76.5	101	150	2,000 <sup>a,d</sup>
Iron, filtered	01046	μg/L	< 6.0	1,490	< 6.0	198	300 <sup>c,d</sup>
Manganese, filtered	01056	μg/L	41.7	421	1.20	105	50°
Nutrients				Concentratio	ons of nutrients	S	
Ammonia + organic N, as N, filtered	00623	mg/L	0.38	0.24	e0.08	0.20	None
Ammonia, as N, filtered	00608	mg/L		_		0.171	None
Nitrate, as N, NO <sub>2</sub> +NO <sub>3</sub> , filtered	00631	mg/L		_		< 0.020	10 <sup>a,d</sup>
Nitrite, as N, filtered	00613	mg/L		_		< 0.001	$1^{a,d}$
Orthophosphate, as P, filtered	00671	mg/L				0.011	None

Table 10.Physical and physiochemical properties of; and concentrations of inorganic major ions, trace elements, nutrients,<br/>selected radiochemicals, and dissolved gases in groundwater samples from bedrock and stratified-drift aquifers in the Town of<br/>Danby, New York, 2010. —Continued

[S&G, sand and gravel; conf, confined aquifer; Parm code, USGS National Water Information System (NWIS) parameter code; mg/L, milligrams per liter; µg/L, micrograms per liter; °C, degrees Celsius; e, estimated; ft, feet; <, less than; —, not analyzed; picocurie per liter (pCi/L); R, radchem non-detect. Bold values exceed one or more drinking-water standards. Sampling site locations are shown in figure 18.]

	USGS Stat	ion Name	TM2519	TM2760	TM1491	TM2806	
	Date	e sampled	6/2/2010	6/2/2010	6/2/2010	10/28/2010	Drinking
	Station I	D Number	422244 076310701	422254 076284801	422029 076255001	422143 076291101	standard
Radiochemicals	Parm code	Units	Сс	oncentrations	of radiochemi	cals	
Gross-alpha (Th-230 curve), unfiltered	01519	pCi/L	3.4	10.7	R 0.5	2.3	15 <sup>a,d</sup>
Gross-beta (Cs-137 curve), unfiltered	85817	pCi/L	0.9	19.1	R 0.3	R 0.7	$4^{a,d}$
Dissolved gases			Co	ncentrations	of dissolved ga	ises	
Methane, unfiltered	85574	mg/L	9.54	0.003	0.001	0.005	10—28°
							$>28^{\rm f}$
Nitrogen, unfiltered	00597	mg/L	19.7	27.6	25.3	25.1	None
Argon, unfiltered	82043	mg/L	0.745	0.919	0.850	0.839	None
Carbon dioxide, unfiltered	00405	mg/L	3.65	20.5	7.8	6.55	None
Oxygen, unfiltered	62971	mg/L	0.3	0.3	0.7	0.7	None

<sup>a</sup> U.S. Environmental Protection Agency maximum contaminant level.

<sup>b</sup> U.S. Environmental Protection Agency secondary drinking water standard.

<sup>c</sup> U.S. Environmental Protection Agency secondary drinking water advisory taste threshold.

<sup>d</sup> New York State Department of Health maximum contaminant level.

<sup>e</sup> Office of Surface Mining recommends that methane concentrations be closely monitored (Eltschlager and others, 2001).

<sup>f</sup> Office of Surface Mining recommends removing ignition sources and venting the gas away from confined spaces (Eltschlager and others, 2001).

<sup>1</sup>CALC = Bicarbonate values were calculated from alkalinity concentrations, which are given in milligrams per liter of CaCO<sub>3</sub>.

0.006 to a maximum of 0.069 mg/L, with a median of 0.012 (tables 6, 8, and 10).

## **Trace Elements**

Trace elements that were detected in every sample included arsenic, barium, boron, lithium, manganese, molybdenum, nickel, strontium, and uranium (tables 9 and 10). In general, concentrations of trace metals were higher in samples from wells that are finished in the bedrock aquifers than samples from wells finished in the confined and unconfined sand and gravel aquifers. The highest detected concentrations of trace elements were 1,660 and 1,490  $\mu$ g/L of iron; 1,430  $\mu$ g/L of strontium; and 1,290 of barium—all samples were from wells that are finished in the bedrock aquifer. Beryllium, silver and thallium were not detected in any samples.

Arsenic concentrations ranged from less than 0.18 to 12.2  $\mu$ g/L, with a median of 2.2  $\mu$ g/L (tables 7, 9, and 10). The concentration of arsenic in one sample (well TM2588, which

is finished in a confined sand and gravel aquifer) exceeded the EPA MCL of 10  $\mu$ g/L. The concentration of arsenic in another sample (well TM1037, which is also finished in a confined sand and gravel aquifer) was 8.2  $\mu$ g/L, which is slightly less than the MCL. Arsenic did not exceed the EPA MCL in samples from the bedrock aquifer or unconfined aquifers.

In the stratified-drift aquifer systems of the northern United States, elevated arsenic concentrations (greater than or equal to 10  $\mu$ g/L) were detected in 9 percent of samples (Thomas, 2007). In comparison, 8 of the 62 (13 percent) wells that were sampled by the USGS and the NYSDOH in the stratified-drift aquifers in Tompkins County from 2000 to 2010 had elevated arsenic concentrations. The distribution of arsenic concentrations in wells that were sampled in stratified-drift aquifers in Tompkins County are shown in figure 19. The predominant source of arsenic is iron oxides and the predominant mechanism for releasing arsenic to the ground water is reductive desorption or reductive dissolution (Thomas, 2007). Also, Thomas (2007) determined that elevated arsenic concentrations were associated with



**Figure 19.** The distribution of arsenic concentrations from groundwater samples in stratified-drift aquifers in Tompkins County from 2000–10 in Tompkins County, New York. (Location of valleys shown in fig. 1.)

strongly reducing conditions such as that in confined aquifers. A confining layer near land surface maintains reducing conditions by retarding the transport of oxygen (or other electron acceptor) from land surface. The slightly higher percentage of samples that had elevated arsenic concentrations in Tompkins County (13 percent) than in the stratified-drift aquifer systems of the northern United States (9 percent) is probably due to confined aquifers that are more prevalent in Tompkins County than in the study by Thomas (2007).

The concentration of aluminum in one unfiltered sample (256  $\mu$ g/L, from well TM1018 finished in the bedrock aquifer) exceeded the EPA MCL 50  $\mu$ g/L. Barium concentrations ranged 36.1 to 1,290  $\mu$ g/L, with a median of 189  $\mu$ g/L (tables 7, 9, and 10). The highest concentration of barium was 1,290  $\mu$ g/L (table 9), which was detected in an unfiltered sample from TM1018 (Danby Town Hall well). Boron concentrations ranged from 11 to 346  $\mu$ g/L, with a median of 45  $\mu$ g/L (tables 7, 9, and 10); MCLs have not been established for boron. Cadmium concentrations ranged <0.018 to 0.026  $\mu$ g/L, with a median of <0.02  $\mu$ g/L (tables 7, 9, and 10). Cadmium concentrations did not exceed the EPA MCL and NYSDOH MCL of 5  $\mu$ g/L. Chromium was detected;

concentrations were low (<0.05  $\mu$ g/L) and did not exceed the EPA and NYSDOH MCLs of 100  $\mu$ g/L. Cobalt concentrations ranged <0.020 to 0.24  $\mu$ g/L, with a median of 0.38  $\mu$ g/L (tables 7, 9, and 10).

Copper was detected at a concentration above the reporting limit of 1.0  $\mu$ g/L in only one unfiltered sample (well TM1018, finished in the bedrock aquifer; table 9). Copper was detected at a concentration 180  $\mu$ g/L, which did not exceed the EPA SDWS of 1,000  $\mu$ g/L.

The concentration of iron in three filtered samples and in one unfiltered sample exceeded the NYSDOH MCL and EPA SDWS of 300  $\mu$ g/L. Iron concentrations ranged from <6.0 to 1,660  $\mu$ g/L, with a median of 218  $\mu$ g/L (tables 7, 9, and 10). The concentration of manganese in seven filtered samples exceeded the EPA SDWS of 50  $\mu$ g/L; the concentration of manganese in two filtered samples exceeded the NYSDOH MCL of 300  $\mu$ g/L. Manganese concentrations ranged from 1.20 to 421  $\mu$ g/L, with a median of 105  $\mu$ g/L (tables 7, 9, and 10).

Lead was detected in four out of nine samples, but none exceeded EPA MCL of 15  $\mu$ g/L. Lead concentrations ranged from <0.06 to 1.48  $\mu$ g/L, with a median of <0.06  $\mu$ g/L (tables 7, 9, and 10). Lithium concentrations ranged from 2.40 to 155  $\mu$ g/L, with a median of 6.21  $\mu$ g/L (tables 7, 9, and 10)—drinking-water standards have not been established for lithium.

Molybdenum concentrations ranged from 0.12 to 7.39  $\mu$ g/L, with a median of 1.97  $\mu$ g/L (tables 7, 9, and 10). Nickel concentrations ranged from 0.19 to 1.1  $\mu$ g/L, with a median of 0.39  $\mu$ g/L (tables 7, 9, and 10). Strontium concentrations ranged from 53.9 to 1,430  $\mu$ g/L, with a median of 290  $\mu$ g/L (tables 7, 9, and 10). Drinking-water standards have not been established for molybdenum, nickel, and strontium.

Selenium concentrations ranged from <0.06 to 0.07  $\mu$ g/L, with a median of <0.06  $\mu$ g/L (tables 7, 9, and 10). Uranium concentrations ranged from 0.009 to 1.47  $\mu$ g/L, with a median of 0.159  $\mu$ g/L. Zinc concentrations ranged from <1.8 to 24.1  $\mu$ g/L, with a median of <2.0  $\mu$ g/L (tables 7, 9, and 10); the concentrations of selenium, uranium, and zinc did not exceed any Federal or State drinking-water standards.

## **Dissolved Gases**

Methane, nitrogen, argon, carbon dioxide, and dissolved oxygen concentrations were determined at four sites where baseline samples for hydraulic fracturing for unconventional development of wells drilled for natural gas were collected (table 10). Methane was detected at low concentrations (0.005 mg/L or less) in 3 of the 4 samples; the concentration in the fourth sample (TM2519; finished in the bedrock aquifer) was 9.54 mg/L, which is slightly lower than the concentration that the Office of Surface Mining recommends that methane concentrations be closely monitored (Eltschlager and others, 2001). Although the EPA and NYSDOH do not have MCLs for methane, dissolved methane concentrations greater than 28 mg/L can pose explosion hazards as a result of methane accumulation in confined spaces (Eltschlager and others, 2001). Nitrogen, argon, and carbon dioxide gases ranged in concentration from 19.7 to 27.6, 0.745 to 0.919, and 3.65 to 20.5 mg/L, respectively (table 10). Dissolved oxygen was determined at 12 sites and concentrations ranged from 0.1 to 2.8 mg/L, with a median value of 0.4 mg/L (tables 6, 8, and 10).

## Comparison to Other Stratified-Drift Aquifers in Tompkins County

The major-ion compositions of water from sampled wells from this study area and from three other stratified-drift aquifer systems in Tompkins County (lower Sixmile Creek and Willseyville Creek trough (Miller and Karig, 2010), upper Sixmile Creek and West Branch Owego Creek valleys (Miller and Sherwood, 2009), and Virgil Creek and Dryden Lake valleys (Miller and Bugliosi, 2013); fig. 1) are presented for comparison using a Piper (trilinear) diagram (fig. 20). A Piper diagram (Piper, 1944) shows the relative proportion of major cations and anions, on a charge-equivalent basis, to the total ion content of the water. Each major ion, in percent of milliequivalents per liter, is shown along the sides of the diagram. Cations are shown in the left triangle and anions are shown in the right triangle. The points are extended into the central diamond-shaped field. The intersection of the projections represents the composition of the water with respect to the combination of ions.

In all stratified-drift aquifers, the cation composition is dominated by calcium and, in a few cases by sodium (fig. 20). The anion composition is dominated by bicarbonate and, in a few samples, by chloride. Results of the groundwater sample analyses for this study indicate that the groundwater in the stratified-drift aguifers in the upper Buttermilk Creek and Danby Creek valleys is dominated by calcium bicarbonate ions. The exceptions are two samples that were collected in upper Buttermilk Creek and Danby Creek valleys and two samples in upper Sixmile Creek and West Branch Owego Creek valleys, which had ions that were sodium bicarbonate (fig. 20). These four samples with sodium bicarbonate to sodium chloride type waters are attributed to enrichment from brackish water in bedrock that locally discharges to the valley fill (Miller, 2009) or to road-salt contamination. Water chemistry in stratified-drift aquifers in all four study areas is similar because the geologic settings of all four areas are similar and are composed of Valley Heads glacial drift and Devonian shales and siltstones.

## Summary

In 2006, the U.S. Geological Survey, in cooperation with the Town of Danby and Tompkins County Planning Department, began a study of the stratified-drift aquifer system in the upper Buttermilk Creek and Danby Creek valleys in the Town of Danby, Tompkins County, New York. In the northern part of upper Buttermilk Creek valley, there is only one stratified-drift aquifer—a basal confined. In the southern part of upper Buttermilk Creek valley, near the Hamlet of Danby, there are as many as four stratified-drift aquifers—two are unconfined and two are confined. In the south-draining Danby Creek valley, an unconfined aquifer consists of outwash and kame sand and gravel—deposited by glacial meltwaters during the late Pleistocene Epoch and of alluvial sediments that were deposited by streams during the Holocene Epoch (post-glacial period).

The principal sources of recharge to the unconfined aquifers in the study area include direct infiltration of precipitation (rain and snowmelt) at land surface, unchanneled surface runoff from adjacent hillsides that seeps into the aquifer along the edges of the valley, groundwater inflow from till and bedrock that enters the aquifer along the sides of the valley, seepage loss from upland-tributary streams where they flow over alluvial fans in the valley, and upward leakage from the underlying geologic units in the valley where the hydraulic head in these units are higher than the water table in the



#### **EXPLANATION**

- imes Upper Buttermilk Creek and Danby Creek Valleys groundwater sample
- + Virgil Creek and Dryden Lake Valleys groundwater sample
- ▲ Lower Sixmile Creek and Willseyville Creek trough groundwater sample
- Upper Sixmile Creek and West Branch Owego Creek Valleys groundwater sample

**Figure 20.** Variability in major ion composition of groundwater in upper Buttermilk Creek and Danby Creek valleys, Virgil Creek and Dryden Lake valleys, Lower Sixmile Creek and Willseyville Creek trough, and upper Sixmile Creek and West Branch Owego Creek valleys, Tompkins County, New York.

unconfined aquifer. The percentages of all sources of recharge to the contiguous unconfined aquifer in Danby Creek valley are 16 percent from precipitation that falls directly over the aquifer, 55 percent from runoff and groundwater inflow from adjacent unchanneled hillsides, and 29 percent from losing streams that cross the aquifer. The total annual recharge to the contiguous unconfined aquifer is 2.56 cubic feet per second (604 million gallons per year).

The principal sources of recharge to the confined aquifers include precipitation that falls directly on the surficial confining unit, which then slowly flows vertically downward through the fine-grained sediments and enters the confined aquifer, and groundwater inflow from till and bedrock that borders the aquifer along adjacent hillsides and at the bottom of the valley. In addition, there is substantial amounts of recharge to the confined aquifers where the confining units are locally absent (forming windows) and where parts of the confining units consist of sediments of low to moderate permeability (forming a semiconfining layer).

Confined aquifers can be under a high degree of confinement; where this condition is present, the confined aquifers, are isolated from sources of recharge by thick, extensive, and low permeability confining units that impede seepage of water, and, therefore, receive small amounts of recharge. However, where the confined aquifer is under a low-to-moderate degree of confinement (semiconfined) or where the confining unit is locally absent, the confined aquifer receives moderate amounts of recharge from precipitation that falls directly over the windows in the confining unit or on the semi-permeable confining unit, which then slowly flows vertically downward and enters the confined aquifer.

In the northern part of the study area (upper Buttermilk Creek valley), groundwater in the stratified drift discharges to domestic and commercial wells, Buttermilk Creek in the area near the northern town border, and a small unnamed stream in a ravine in Buttermilk State Park just north of the town border. In the southern part of the study area (Danby Creek valley), groundwater discharges to domestic, commercial, and farm wells; to Danby Creek; to a large wetland in the central parts of Danby Creek valley; and as losses because of plant uptake and evaporation. About 300 people depend on groundwater from the Buttermilk Creek and Danby Creek stratified-drift aquifer system.

An unconfined surficial aquifer about 8,000 feet (ft) long and as much as 800 ft wide, with a saturated thickness of about 20 ft, occupies the lower (southeastern most) 8,000 ft of Danby Creek valley within the Town of Danby. The average annual recharge to the contiguous unconfined aquifer is 2.56 cubic feet per second (1.66 million gallons per day). However, because the volume of water stored in the aquifer is small and the potential for induced recharge from Danby Creek during summer periods of low flow is also small, an array of wells would probably be needed to provide a sustainable continuous amount of water to large water users such as municipalities and industries. Well data from water-well drillers through 2012 indicate that the confined aquifers in upper Buttermilk Creek valley are thin (typically about 10 ft thick) and the reported well-yield data ranged from 3 to 50 gallons per minute. These well yield data suggest that the aquifer meets the needs of homeowners and small commercial facilities, but these aquifers may not be capable of supplying sufficient water to meet the needs of municipalities and industries. Additional geohydrologic data leading to calibration of a groundwater flow model would be required to evaluate the long-term potential yield from the confined aquifers in upper Buttermilk Creek valley.

During 2007-10, groundwater samples were collected from 13 wells that tap the stratified-drift and bedrock aquifers-seven samples were from the confined stratifieddrift aquifers, one from an unconfined stratified-drift aquifer, and five were from bedrock aquifers. The samples were collected and processed using standard U.S. Geological Survey procedures and were analyzed for physiochemical properties and chemical constituents, including dissolved gases, major ions, nutrients, and trace elements. Calcium dominates the cation composition and bicarbonate dominates the anion composition in most groundwater. Water quality in the study area generally meets Federal and New York State drinking-water standards, but concentrations of some constituents exceeded the standards. The standards that were exceeded were sodium (3 samples), dissolved solids (1 sample), iron (3 samples), manganese (8 samples), and arsenic (1 sample).

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Appendix 1. Records of Wells, Town of Danby, Tompkins County, New York

Appendix 1. Records of wells, Town of Danby, Tompkins County, New York.

[At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table may be obtained from the U.S. Geological Survey's NWIS database. ft, feet; in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; ---, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Date drilled	Well depth (ft)	Depth of cas- ing	Casing diam- eter	Altitude land surface (NAVD 88,	Aquifer type	Water level below land sur-	Altitude water level (NAVD 88,	Date water level mea- sured	Depth to Bed- rock	Altitude top of bedrock (NAVD 88,	Re- ported yield (gal/	Remarks (Data from driller's logs except those that were drilled by the
		(H)	(III)	in ft)		faceb (ft)	in ft)		(Ht)	in ft)	min)	U.S. Geological Survey)
	180	63	5	1,240	Shale	15	1,225	9/15/1966	57	1,183	30	0-7 S&G, 7-57 till, 57-180 ft shale
	200	50	9	1,216	Shale	0	1,216	10/9/1997	40	1,176	0	WQ sample 8/16/2007. 0-40 overburden, 40-200 ft shale
	128	17	8	1,305	Shale	15	1,290	10/15/1971	17	1,288	15	0–17 till, 17–128 ft shale
	200	18	9	1,275	Shale	45	1,230	2/2/1994	15	1,260	8	0-15 till, 15-200 ft shale
	180	18	9	1,309	Shale	30	1,279	2/5/1994	5	1,304	9	0-5 till, 5-180 ft shale
	20	10	2	1,243	Shale	13	1,230	1/29/1990	20	1,223		0-15 S&G, 15-20 till, 20-204 ft shale
	12	$\mathfrak{c}$	7	1,233.2	Shale	4.2	1,229	1/29/1990	85	1,148		Screen 3–12 ft. 0–35 sand, 35–85 till, 85–120 ft shale
	12	ς	7	1,231.9	Shale	2.2	1,229.7	1/30/1990	8	1,224		Screen 3–12 ft. 0–4 S&G, 4–8 till, 8–12 ft shale
	88	81	9	1,161	Shale	5.8	1,155.2	7/7/2009	68	1,093	1	WQ sample 11/13/2008. 0–18 S&G, 18–22 silt with some gravel, 22–32 S&G, 32–44 till, 44–48 sand, 48–68 till, 68–80 ft weathered shale
	150	150	9	1,114.5	Confined S&G	73.9	1,040.6	7/7/2009			50	WQ sample 11/13/2008
	30	30	9	1,197.3	Confined S&G	3.8	1,193.5	7/7/2009				WQ sample 12/10/2008. Well is below land surface
	23	23	5	1,217.3	Confined S&G	2.4	1,214.9	7/7/2009				WQ sample 12/18/2008. Owner reports good yield
	140	94	9	1,120	Shale				06	1,030	15	0-25 gravel, 25-50 blue clay with shale (till), 50-90 S&G, 90-140 shale
	66	66	9	1,290	Confined S&G	25	1,265	4/27/2000			20	0–30 gravel, 30–90 fine sand, 90–100 ft gravel
	120	52	9	1,415	Shale	41	1,344	5/3/2000	48	1,367	9	0–30 till, 30–40 silty sand, 40–48 till, 48–120 ft shale

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[At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table may be obtained from the U.S. Geological Survey's NWIS database. ft, feet; in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; —, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Site name	Date drilled	Well depth (ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- faceb (ft)	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
TM1212	5/9/2000	358	25	9	1,522	Shale	80	1,442	5/9/2000	18	1,504	7	0-18 till, 18-358 ft grey shale
TM1214		138	54	9	1,230	Shale				54	1,176	4	0-54 till, 54-138 ft shale
TM1216	6/22/2000	100	20	9	1,500	Shale	22.5	1,477.5	6/22/2000	18	1,482	5	
TM1234	9/28/2000	28	28	S	1,220	Confined S&G	7	1,213	9/28/2000	I		28	0–26 clay and gravel (till), 26–28 ft $S\&G$
TM1235	7/11/2000	44	44	9	1,199	Confined S&G	6.0	1,193	7/11/2000			15	0–18 fine gravel, 20–42 till, 42–44 ft S&G
TM1267	10/6/2000	358	20	9	1,069	Shale	50	1,019	10/6/2000	8	1,061	0.5	0-8 till, 8-358 ft grey shale
TM1271	10/24/2000	360	160	9	1,143	Shale	130	1,013	10/24/2000	160	983	-	0–80 till, 80–97 till with thin layers S&G, 97–160 clay with thin layers S&G, 160–360 ft grey shale
TM1272	10/19/2000	120	30	9	1,583	Shale	26	1,557	10/19/2000	20	1,563	30	0-20 till, 20-120 ft grey shale
TM1273	10/26/2000	280	156	9	1,152	Shale	60	1,092	10/26/2000	156	966	-	0-50 till, 50-100 thin S&G layers and clay, 100-156 clay, 156-280 ft shale
TM1281	11/16/2000	338	26	9	1,075	Shale	б	1,072	11/16/2000	25	1,050	0.25	0-25 till, 25-338 ft shale
TM1283	12/13/2000	75	39	9	1,385	Shale	23	1,362	12/13/2000	39	1,346	7	0-30 S&G, 30-39 till, 39-75 ft shale
TM1295	1/29/2001	120	20	9	1,128	Shale	45	1,083	1/29/2001	10	1,118	S	0-10 till, $10-120$ ft shale DD = $30$ ft at 5 gal/min
TM1303	4/26/2001	73	73	9	1,490	Confined gravel	20	1,470	4/26/2001			9	0–20 till, 20–25 silty gravel, 25–70 till, 70–73 ft gravel
TM1317	5/15/2001	52	52	9	1,412	Confined gravel	18.5	1,393.5	5/15/2001				0–25 till, 25–52 ft gravel
TM1319	7/13/2001	200	55	9	1,038	Shale	43	995	7/13/2001	53	985	2	0–53 gravel, 53–200 ft shale
TM1320	5/25/2001	150	35	5	1,210	Shale	15	1,195	5/25/2001	27	1,183	4	0-17 till, 17-19 S&G, 19-27 till, 27-150 ft grey shale
TM1342	1/10/2001	150	105	9	1,245	Shale	25	1,220	1/10/2001	104	1,141	5	0-45 till, 45-104 cemented gravel, 104-150 ft shale

Appendix 1. Records of wells, Town of Danby, Tompkins County, New York.—Continued

[At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table may be obtained from the U.S. Geological Survey's NWIS database. ft, feet, in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; —, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Date M drilled de (	A b	/ell pth ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- faceb	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
46	11/14/2002	360	155	9	1,290	Shale	55	1,235	11/14/2002	153	1,137	5	0-153 till, 153-360 ft grey shale
- 1	6/6/2001	165	27	5	1,645	Shale	28	1,617	6/6/2001	20	1,625	4	0–19 till, 19–20 gravel (5 gal/min), 20–165 ft grey shale
	11/15/2001	300	18	5	1,425	Shale	15	1,410	11/15/2002	16	1,409	б	0-16 till, 16-300 ft grey shale
~	11/1/2001	179	20	9	1,676	Shale	99	1,610	11/1/2001	18	1,658	10	1-18 till, 18-180 ft shale
<del></del>	12/1/2001	358	118	9	1,175	Shale	50	1,125	12/1/2001	118	1,059	1	0-88 till, 88-91 gravel (2 gal/min), 91-101 S&G, 101-111 silty S&G, 111-118 till, 118-358 ft shale
	1/15/2002	202	103	9	1,220	Shale	45	1,175	1/15/2002	103	1,117	7	0-30 S&G, 30-103 cemented gravel, 103-202 ft shale
2	4/4/2002	160	61	9	1,495	Shale	70	1,425	4/4/2002	60	1,435	б	0–35 till, 35–40 gravel, 40–60 till, 60–160 ft shale
_	4/25/2002	180	47	0	1,232	Shale	80	1,152	4/25/2002	44	1,188	б	0–15 dirty gravel, 15–44 till, 44–180 ft grey shale
_	5/30/2002	300	22	9	1,742	Shale	49	1,693	5/30/2002	12	1,730	Ŷ	WQ sample 6/2/2010. 0–12 till, 12–300 ft shale
	5/23/2002	135	135	9	1,125	Confined S&G	82	1,043	5/23/2002			20	0–15 S&G, 15–80 cemented gravel, 80–134 till, 134–135 ft gravel
$\sim$	5/31/2002	45	45	S	1,265	Confined gravel	35	1,230	5/31/2002			15	0-20 till, 20-45 ft gravel
_	6/20/2002	69	48	9	1,490	Shale	25	1,475	6/20/2002	46	1,444	15	0–44 till, 44–46 dirty gravel, 46–69 ft shale
~	6/11/2002	183	23	9	1,730	Shale	60	1,670	6/11/2002	23	1,707	0	0-23 till, 23-183 ft shale
<u>,                                    </u>	6/24/2002	95	40	9	1,330	Shale	20	1,310	6/23/2002	40	1,290	ŝ	0-40 till, 40-95 ft shale
8	6/24/2002	43	43	9	1,160	Confined gravel	Flow- ing		6/24/2002			15	0-40 till, 40-43 ft gravel
	7/9/2002	54	54	9	1,413	Confined gravel	15	1,398	7/9/2002			12	0-54 till, 54 ft gravel and shale

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[At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table may be obtained from the U.S. Geological Survey's NWIS database. ft, feet; in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; —, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Site name	Date drilled	Well depth (ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- (ft)	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
TM1547	7/11/2002	57	38	9	1,598	Shale	21	1,577	7/11/2002	36	1,562	15	0–36 till, 36–57 ft shale
TM1578	8/20/2002	78	24	9	1,060	Shale	12	1,048	8/22/2002	20	1,040	9	0-20 till, 20-78 ft shale
TM1607	9/26/2002	40	40	9	1,174	Confined gravel						20	0–35 till, 35–40 ft gravel
TM1618	10/15/2002	109	109	9	1,134	Confined gravel	75	1,059	10/5/2002			10	0-10 top soil, 10-20 clay and gravel, 20-104 till, 104-109 ft S&G
TM1619	10/25/2002	83	83	9	1,293	Confined gravel	30	1,263	10/25/2002			б	0–76 till, 76–103 ft gravel
TM1622	10/30/2002	203	116	9	1,284	Shale				116	1,168	7	0–21 till, 21–116 gravel w/ clay, 116–203 ft shale
TM1654	2/20/2003	41	41	Ś	1,174	Uncon- fined S&G	26	1,148	2/20/2003			10	0–38 silty gravel (alluvium), 38–41 ft clean gravel
TM1659	3/6/2003	200	56	9	1,548	Shale	40	1,508	3/6/2003	56	1,492	4	0-56 till, 56-200 ft shale
TM1678	4/29/2003	240	45	9	1,235	Shale	50	1,185	4/29/2003	45	1,195	7	0-45 till, 45-240 ft grey shale
TM1681	5/2/2003	297	20	9	1,475	Shale	99	1,409	5/2/2003	С	1,472	5	0–3 till, 3–297 shale
TM1683	5/20/2003	180	59	9	1,155	Shale	80	1,075	5/20/2003	58	1,097	7	0–30 till, 30–40 dirty gravel, 40–58 till, 58–180 ft shale
TM1685	5/1/2003	360	226	9	1,490	Shale	60	1,430	5/1/2003	222	1,268	4	0-85 till, 85-87 gravel (2 gal/min), 87-222 till, 222-360 ft grey shale
TM1710	6/24/2003	220	119	9	1,244	Shale	25	1,219	6/24/2003	117	1,127	S	0-80 till, 80-83 S&G, 83-117 till, 117-220 ft grey shale
TM1711	7/10/2003	76	67	9	1,250	Confined gravel	Flow- ing		7/10/2003			10	0-50 till, 50-97 dirty gravel
TM1712	7/29/2003	95	95	9	1,255	Confined gravel	Flow- ing		7/29/2003			20	0–89 till, 89–95 ft gravel
TM1740	8/18/2003	183	83	5	1,266	Shale	66	1,200	8/18/2003	83	1,183	4	0-15 till,15-60 fine gravel, 60-83 silty clay (till?), 83-183 ft shale

Appendix 1. Records of wells, Town of Danby, Tompkins County, New York.—Continued

[At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table may be obtained from the U.S. Geological Survey's NWIS database. ft, feet; in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; ---, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Site name	Date drilled	Well depth (ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- (ft)	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
TM1743	8/14/2003	30	30	9	1,178	Uncon- fined S&G	0.1	1,178	8/14/2003			23	No log, water level at land surface
TM1748	9/16/2003	180	21	9	1,165	Shale	60	1,105	9/16/2003	4	1,161	1	0-4 till, 4-180 ft shale
TM1768	7/15/2005	220	26	9	1,083	Shale				18	1,065	0	1–18 till, 18–220 ft shale
TM1789	10/27/2003	360	110	9	1,232	Shale	40	1,192	10/27/2003	100	1,132	-	0-65 till and sand layers, 65-67 S&G, 67-100 till, 100-360 ft grey shale
TM1792	11/11/2003	300	46	9	1,500	Shale	37	1,463	11/12/2003	46	1,454		0-46 till, 46-300 ft shale
TM1793	11/11/2003	360	49	9	1,355	Shale	30	1,325	11/11/2003	47	1,308	1	0-47 till, 47-360 shale
TM1796	11/24/2003	145	132	9	1,124	Shale	30	1,094	11/24/2003	132	992	9	Bedrock at 132 ft
TM1815	2/18/2004	220	57	9	1,198	Shale	70	1,128	2/18/2004	55	1,143	12	0-50 till, 50-55 S&G, 55-220 ft shale
TM1830	4/14/2004	200	29	9	1,430	Shale	20	1,410	4/14/2004	25	1,405	7	0-25 till, 25-200 ft shale
TM1833	4/15/2004	260	31	9	1,300	Shale				20	1,280	9	0-20 till, 20-260 ft shale
TM1841	5/3/2004	260	39	9	1,190	Shale	10	1,180	5/3/2004	39	1,151	7	0–39 till, 39–260 ft shale
TM1842	5/4/2004	240	50	9	1,218	Shale	70	1,148	5/4/2004	48	1,170	5	0-48 till, 48-240 ft shale
TM1844	5/5/2004	280	55	9	1,160	Shale	10	1,150	5/6/2004	55	1,105	7	2–26 till, 26–28 S&G, 28–32 clay, 32–38 S&G, 38–55 till, 55–280 ft shale
TM1847	4/10/2004	100	17	9	1,121	Shale	1	1,120	5/10/2004	17	1,104	4	0–17 till, 17–100 ft shale
TM1883	7/15/2004	106	19	9	1,264	Shale	10	1,254	7/15/2004	19	1,245	9	0–19 S&G, 19–106 ft shale
TM1886	7/9/2004	148	20	9	1,415	Shale	09	1,355	7/9/2004	б	1,412	8	0–3 till, 3–148 ft shale
TM1893	7/23/2004	130	12	5	1,505	Shale	30	1,475	7/23/2004	8	1,497	3	0-8 till, 8-130 ft Shale
TM1894	7/24/2004	140	30	9	1,704	Shale	70	1,634	7/24/2004	25	1,779	10	0-25 till, 25-140 ft shale
TM1896	8/5/2004	99	66	9	1,070	Confined S&G						10	
TM1897	8/16/2004	140	20	9	1,255	Shale	30	1,225	8/16/2004	8	1,247	5	0-8 till, 8-140 ft shale

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	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- faceb (ft)	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
20		9	1,462	Shale	50	1,412	9/7/2004	16	1,446	9	0-16 till, 16-100 ft shale
22		9	1,498	Shale	77	1,421	9/22/2004	20	1,478	17	0-20 till, 20-154 ft shale, water supply for stock
110		9	1,590	Confined S&G	40	1,550	10/5/2004			8	Clay at land surface
156		9	770	Confined sand	130	640				2	0–3 gravel, 3–20 clay, 20–60 sandy clay, 60–80 sand, 80–90 gravel, 90–156 heaving fine sand
21		9	1,190	Shale	30	1,160	11/17/2004	9	1,184	4	0–6 till, 6–145 ft shale
20		9	1,430	Shale				4	1,426	3	0-4 till, 4-240 ft grey shale
50	•	9	1,425	Confined S&G	25	1,400	11/29/2004			10	0-45 till, 45-50 ft gravel
21		9	1,145	Shale	100	1,045	12/7/2004	15	1,130	4	0-15 till, 15-200 ft shale
22		9	1,070	Shale	100	970	12/9/2004	15	1,055	30	0-15 till, 15-220 ft shale
26	-	9	1,592	Shale				22	1,570	5	0-22 till, 22-240 ft grey shale
20	-	9	1,321	Shale	300	1,292	6/26/2005	18	1,303		0-18 till, 18-300 ft shale
23	Ũ	ý	1,385	Shale	80	1,305	7/20/2005	10	1,375	9	0-10 till, 10-180 shale
20	Ŭ	<i>.</i> 0	1,455	Shale	20	1,435	10/20/2005	10	1,445	7	0-10 till, 10-280 ft shale
21		9	1,275	Shale	50	1,225	8/12/2005	5	1,270	4	0-5 till, 5-120 ft grey shale
24		9	1,620	Shale	45	1,575	9/15/2005	24	1,596	б	0–24 till, 24–140 ft shale
26		9	1,465	Shale				25	1,440	4	0–25 till, 25–60 ft shale
33		9	1,655	Shale				32	1,623	5	0-32 till, 32-140 ft grey shale
19		9	1,485	Shale	20	1,466	1/25/2006	13	1,472	8	0-2 soil, 2-13 till, 13-105 shale, 105-115 sandstone, 115-127 ft shale
32		9	1,238.6	Shale	06	1,148	2/14/2006	30	1,209	10	0-30 till, 30-240 ft grey shale
20		9	1,300	Shale	25	1,275	5/19/2006	10	1,290	25	0-10 till, 10-208 ft shale

Appendix 1. Records of wells, Town of Danby, Tompkins County, New York.—Continued

[At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table may be obtained from the U.S. Geological Survey's NWIS database. ft, feet; in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; ---, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Site name	Date drilled	Well depth (ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- faceb (ft)	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
TM2253	8/31/2006	297	20	9	1,596	Shale	100	1,496	8/31/2006	10	1,586	12	0-10 till, 10-297 ft shale
TM2297	10/24/2006	120	120	9	1,565	Confined S&G	30	1,535	10/27/2006				0-110 till, 110-120 ft gravel
TM2300	10/26/2006	220	54	9	1,125	Shale	40	1,085	10/26/2006	50	1,075	-	0-20 till, 20-50 clay, 50-220 shale
TM2311	11/10/2006	300	20	9	1,225	Shale	24	1,201	11/13/2006	5	1,220	7	0-5 till, 5-300 ft shale
TM2333	1/22/2007	240	24	9	1,181	Shale	72	1,109	1/22/2007	24	1,157	4	0-24 till, 24-240 ft grey shale
TM2346	3/19/2007	50	24	9	1,328	Shale	14	1,314	3/19/2007	18	1,310	12	0–11 till, 11–18 S&G, 18–50 ft shale Specific capacity = 0.52 gal/ft
TM2369	5/25/2007	240	29	9	1,215	Shale	50	1,165	5/25/2007	28	1,187	З	0-28 till, 28-240 ft grey shale
TM2370	5/24/2007	360	69	9	1,154	Shale				68	1,086	0.5	0-68 till, 68-360 ft shale
TM2399	7/10/2007	68	68	9	1,045	Confined S&G	30	1,015	7/10/2007			30	0–30 till, 30–40 S&G, 40–52 fine to coarse sand, 52–60 S&G, 60–68 fi gravel
TM2459	10/31/2007	98	121	9	1,260	Confined, S&G and shale	31	1,229	10/31/2007	110	1,150	3	Perforated casing 95–98 ft; 0–59 till, 58–65 gravel, 65–97 till, 97–99 gravel, 99–110 till, 110–240 ft shale
TM2463	11/1/2007	140	48	9	1,261	Shale	40	1,221	11/7/2007	48	1,213	0	0–37 till, 37–40 S&G, 40–48 till, 48–140 ft shale
TM2467a	11/14/2007	75	6	9	1,224.2	Confined S&G	2.5	1,221.8	7/7/2009	82	1,142	10	WQ sample 4/24/2008. 0–19 till, 19–23 S&G, 23–58 till, 58–82 S&G, 82–89 ft shale
TM2483	1/10/2008	133	133	9	1,109	Confined S&G	101	1,008	1/10/2008			8	0-40 till with S&G layers, 40-130 till, 130-133 ft gravel
TM2519	5/8/2008	420	20	9	1,261	Shale		1,261		16	1,245	1.5	WQ sample 6/2/2010. Yield before hydraulic fracking was 0.1 gal/min, after hydraulic fracking it was 1.5 gal/min. 0–16 till, 16–420 ft shale

Appendix 1. Records of wells, Town of Danby, Tompkins County, New York.—Continued

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Site name	Date drilled	Well depth (ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- (ft)	Altitude water level in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
TM2523	5/20/2012	136	136	9	1,107	Confined S&G	110	266	5/20/2008			20	0-60 till, 60-80 sandy clay, 80-135 hard grey clay (till), 135-136 ft gravel
TM2542	6/3/2008	160	106	9	1,238	Shale	20	1,218	6/3/2008	103	1,135		0-103 gravelly sand, gray clay (till), 103-160 ft shale
TM2561	8/8/2008	160	30	9	1,645	Shale	I			4	1,641	7	0-4 till, 4-24 weathered shale, 24-160 ft grey shale
TM2588a	10/21/2008	41	41	9	1,220.6	Confined S&G	14.1	1,206.5	7/7/2009	55	1,166	4	WQ sample 11/12/2008. 0–10 pebbly fine sand, 10–18 silty fine sand, 18–32 silty clay, 32–38 till, 38–45 gravel, 45–55 till, 55–70 ft shale
TM2589a	10/24/2008	100	89	9	1,207.2	Shale	3.4	1,203.8	7/7/2009	87	1,120		0–15 silty gravel or till, 15–20 f sand, 20–24 clay or till, 24–31 S&G, 31–43 till, 43–46, fine sand, 46–87 till, 87–100 ft shale
TM2590a	10/28/2008	27	27	9	1,161	Uncon- fined S&G	10.1	1,150.9	7/7/2009			20	WQ sample 11/13/2008. 0-18 S&G, 18-22 silt with some gravel, 22-27 ft S&G
TM2591a	10/24/2008	27	27	9	1,207.2	Confined S&G	6.43	1,200.8	7/7/2009			20	WQ sample 11/12/2008. 0–15 silty gravel or till, 15–20 fine sand, 20–24 clay or till, 24–27 ft S&G
TM2667	7/13/2009	125	20	9	1,335	Shale	40	1,295	7/13/2009	8	1,327	8	0-8 till 8-125 ft shale
TM2679	8/17/2009	400	107	9	1,142	Shale	85	1,057	8/17/2009	107	1,035	15	0-10 till, 10-90 till with seams of S&G, 90-107 till with gravel lay- ers, 107-400 ft shale
TM2734	3/22/2010	94	24	9	1,081	Shale	9	1,075	3/22/2010	21	1,060	10	0-21 till, 21-94 ft shale

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Appendix

of land surface at wells were estimated using lidar and 1:24,000 scale topographic contour maps that were vertically accurate to 0.5 ft and 5 ft, respectively. The latitude and longitude of the wells in this table [At selected wells, the altitudes of water-level-measuring points, which are typically the tops of the well casings, were determined to within 0.01 ft using standard surveying techniques. Elsewhere, altitudes may be obtained from the U.S. Geological Survey's NWIS database. ft, feet; in, inches; NAVD 88 in ft, North American Vertical Datum of 1988, in feet: gal/min, gallons per minute; gal/d, gallons per day; S&G, sand and gravel; ---, no data; and WQ, water quality; =, equal to; DD, drawdown; HVSR, horizontal-to-vertical spectral ratio seismic survey site]

Site name	Date drilled	Well depth (ft)	Depth of cas- ing (ft)	Casing diam- eter (in)	Altitude land surface (NAVD 88, in ft)	Aquifer type	Water level below land sur- (ft)	Altitude water level (NAVD 88, in ft)	Date water level mea- sured	Depth to Bed- rock (ft)	Altitude top of bedrock (NAVD 88, in ft)	Re- ported yield (gal/ min)	Remarks (Data from driller's logs except those that were drilled by the U.S. Geological Survey)
TM2760		172	170	S	1,413	Shale	48	1,365	6/2/2010	170	1,243	9	WQ sample 6/2/2010. 0–103 till, 103– 114 unknown, 114–138 red clay, 138–163 dirty gravel, 163–170 till, 170–172 ft shale
TM2806a	10/14/2010	47	47	9	1,178.5	Confined S&G	+5.9	1,184.4	10/28/2010	58	1,120	20	WQ sample 10/28/2010. Flows, 0–5 fill, 5–19 till, 19–25 S&G, 25–40 till, 40–47 ft S&G. Bedrock at 58 ft (HVSR)
TM2808	10/11/2010	220	45	9	1,183	Shale	50	1,133	10/11/2010	45	1,138	12	0-45 till, 45-220 ft shale
TM2840	3/16/2011	160	160	9	1,132.0	Confined S&G	122	1,010	3/16/2011			20	0–15 till, 15–20 silty gravel, 20–157 till, 157–160 ft gravel
TM2882	7/28/2011	76	76	9	1,177	Confined S&G	20	1,157	7/28/2011	77	1,107	10	0-40 sandy clay and gravel layers, 40-74 till, 74-76 ft gravel, bedrock reported at 77 ft
TM2887	7/30/2011	105	27	9	1,155	Shale	20	1,135	7/30/2011	27	1,128	7	0-27 till, 27-105 ft shale
TM2905	8/10/2011	71	71	9	1,206	Confined S&G	14	1,192	8/10/2011				
TM2909	9/20/2011	360	88	9	1,159	Shale	09	1,099	9/20/2011	88	1,071	7	0–54 till, 54–88 till and clay layers, 88–360 ft shale
TM2923	11/16/2011	280	28	9	1,251	Shale	6	1,245	11/16/2011	24	1,227	5	0-15 till, 15-18 gravel, 18-24 till, 24-30 weathered shale, 30-280 ft shale
Gas well	6/1/1939	2,351			1,178					97	1,081		Bedrock at 97 ft. Hit water at 41 ft (S&G), fresh water noted at 545 ft (limestone)
<sup>a</sup> Well dril <sup>b</sup> Water lev	led by the U.S.	Geologica w land su	ll Survey. Irface exce	int those va	lues preceeded	hv a plus sign	which inc	licates the not	antiometric surf.	are above	and surface		

60 Geohydrology and Water Quality, Stratified-Drift Aquifers, Upper Buttermilk Creek and Danby Creek Valleys, N.Y.

Appendix 2. Well Logs and Construction Details of U.S. Geological Survey Test Wells

# **USGS TEST WELL TM2588**

Danby Park Association, Town of Danby, N.Y.

Site name: TM2588 (well depth = 41 ft) Site ID: 422116076284701 Latitude: 42° 21' 15.54'' Longitude: 076° 28' 46.62'' Date completed: 10/21/2008 Drilling contractor: Barber & DeLine, Tully, NY 6 inch diameter steel casing Casing above ground = 2.9 ft

Latitude and longitude measurement made by GPS (NAD83)



Pulled casing back from 46 ft (no water) to 41 ft and developed well for 1 hr at rate of 3-4 gal/min.

**Figure 2–1.** Well log and well construction details of U.S. Geological Survey test well TM2588 near the Hamlet of Danby, Town of Danby, Tompkins County, New York.
## **USGS TEST WELL TM2467**

Danby Highway Department, 93 Hornbrook Rd., Danby, N.Y.

Site ID: 422119076282401 Site name: TM2467 (75 ft deep well) Latitude: 42° 21' 19.16'' Longitude: 076° 28' 23.71'' Date completed: 11/14/2007 Drilling contractor: Barber & DeLine, Tully, NY 6 inch diameter steel casing Casing above ground = 1.4 ft

Latitude and longitude measurement made by GPS unit (NAD83)



11/14/2007- Developed well at 6 gal/min for 30 min. with casing at 80 ft (water was turbid at end of 30 minutes- pumping vf sand and silt)

Pulled casing back to 75 ft and developed well for 1 hr 10 min., started 30 gal/min but yield decreased to 10-15 gal/min after 30 min., and then to 5 - 10 gal/min after one hour. (still had some turbid water after 1 hr. 10 min.) May need to develope well for a longer time to determine whether water will clear up.

Developed well 5 hrs on 4/24/2008 by air lift. Started pumping at 2.2 gal/min (Drawdown = 7.5 ft), then increased pumping to 4.4 gal/min (DD = 14.5 ft). Water cleared up at end of pumping (5 hrs).

Collected water quality sample (common ions, trace metals, nutrients)

**Figure 2–2.** Well log and well construction details of U.S. Geological Survey test well TM2467 at the Danby Highway Department near Hornbrook Road, Town of Danby, Tompkins County, New York.

## USGS TEST WELLS TM2590 and TM1035

Durfee Hill Road, Town of Danby, N.Y.

Latitude and longitude measurement made by GPS (NAD83)

Site name: TM2590 (well depth = 27 ft) Site ID: 421934076251001Latitude:  $42^{\circ}$  19'  $34.01^{"}$ Longitude:  $076^{\circ}$  25'  $10.21^{"}$ Date completed: 10/28/20086 inch diameter steel casing Casing above ground = 3.0 ft Site name: TM1035 (well depth = 88 ft) Site ID: 421934076251002Latitude:  $42^{\circ}$  19' 34.06''Longitude:  $076^{\circ}$  25' 10.14''Date completed: 10/28/20086 inch diameter steel casing Casing above ground = 2.7 ft

Drilling contractor: Barber & DeLine, Tully, NY TM1035 TM2590 Altitude relative Elev. TOC (6 in.) Elev. TOC (6 in.) Altitude of land surface is 1161 feet to NGVD 88 above NGVD 88 = 1164 ft = 1162.7 ft 0 1161 Silty coarse gravel, brown, poorly sorted Water level Water level = 10.4 ft = 5.7 ft below land below land surface. surface. 11/5/2008 11/5/2008 18 Silt with trace gravel, soft, wet Depth below land surface, in feet 22 Flev. WI = Elev. WL = Fine to coarse sand and gravel with silt, subround 1150.6 ft 1157.0 ft to round clasts, makes water (unconfined aquifer) 32 Water from sand and gravel aquifer Till, grey, stones embedded in sand and silt matrix Δ enters open-end Δ 6-in casing at 27 ft 44 Well depth = 27 ft Sand, grey, soft, wet 48 (6" dia. casing) Till, grey, stones embedded in sand and silt matrix 68 09 Shale, dark grey, makes little water (1.5 gal/min) Bottom of casing = 81 ft Water from bedrock enters open-hole 88 073 interval from Bottom of hole = 88 ft Well depth = 88 ft 81 to 88 ft

Developed well TM2590 for 1 hr at rate of 20 gal/min. Developed well TM1035 for 0.5 hr at rate of 1.5 gal/min.

**Figure 2–3.** Well logs and well construction details of U.S. Geological Survey test wells TM2590 and TM1035, Durfee Hill Road, Town of Danby, Tompkins County, New York.

## USGS TEST WELLS TM2589 and TM2591

Steam Mill Road, Town of Danby, N.Y.

Site name: TM2591 (well depth = 27 ft) Site ID: 422052076273101 Latitude:  $42^{\circ}$  20' 51.60" Longitude:  $076^{\circ}$  27' 31.09" Date completed: 10/24/20086 inch diameter steel casing Casing above ground = 3.2 ft Site name: TM2589 (well depth = 100 ft) Site ID: 422052076273102 Latitude:  $42^{\circ}$  20' 51.57" Longitude:  $076^{\circ}$  27' 30.97" Date completed: 10/24/20086 inch diameter steel casing Casing above ground = 2.9 ft

Latitude and longitude measurement made by GPS (NAD83) Drilling contractor: Barber & DeLine, Tully, NY



Developed well TM2591 for 1 hr at rate of 20 gal/min. Developed well TM2589 for 0.5 hr at rate of 0.1 gal/min.

**Figure 2–4.** Well log and well construction details of U.S. Geological Survey test wells TM2589 and TM2591, Steam Mill Road, Town of Danby, Tompkins County, New York.



bedrock was at depth = 53 ft

**Figure 2–5.** Well log and well construction details of U.S. Geological Survey test well TM2806, Whitehawk Lane, Town of Danby, Tompkins County, New York.

For more information concerning this report, contact: Director, New York Water Science Center U.S. Geological Survey 30 Brown Road Ithaca, NY 14850

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